



Production of ^{40}Ar by an overlooked mode of ^{40}K decay with implications for K-Ar geochronology

Jack Carter¹, Ryan B. Ickert^{1,2}, Darren F. Mark^{1,3}, Marissa M. Tremblay², Alan J. Cresswell¹, and David C. W. Sanderson¹

¹SUERC, Rankine Avenue, Scottish Enterprise Technology Park, East Kilbride, Scotland, UK

²Department of Earth, Atmospheric, and Planetary Sciences, Purdue University, West Lafayette, IN 47901, USA

³Earth and Environmental Sciences, University of St Andrews, College Gate, St Andrews, Scotland, UK

Correspondence: Jack Carter (j.carter.1@research.gla.ac.uk)

Received: 31 March 2020 – Discussion started: 22 April 2020

Revised: 27 August 2020 – Accepted: 30 September 2020 – Published: 26 November 2020

Abstract. The decay of ^{40}K to the stable isotopes ^{40}Ca and ^{40}Ar is used as a measure of time for both the K-Ca and K-Ar geochronometers, the latter of which is most generally utilized by the variant $^{40}\text{Ar}/^{39}\text{Ar}$ system. The increasing precision of geochronology has forced practitioners to deal with the systematic uncertainties rooted in all radioisotope dating methods. A major component of these systematic uncertainties for the K-Ar and $^{40}\text{Ar}/^{39}\text{Ar}$ techniques is imprecisely determined decay constants and an incomplete knowledge of the decay scheme of ^{40}K . Recent geochronology studies question whether ^{40}K can decay to ^{40}Ar via an electron capture directly to ground state ($\text{EC}_{\text{ground}}$), citing the lack of experimental verification as reasoning for its omission. In this study, we (1) provide a theoretical argument in favor of the presence of this decay mode and (2) evaluate the magnitude of this decay mode by calculating the electron capture to positron ratio ($\text{EC}_{\text{ground}}/\beta^+$) and comparing calculated ratios to previously published calculations, which yield $\text{EC}_{\text{ground}}/\beta^+$ between 150–212. We provide support for this calculation through comparison of the experimentally verified $\text{EC}_{\text{ground}}/\beta^+$ ratio of ^{22}Na with our calculation using the theory of β decay. When combined with measured values of β^+ and β^- decay rates, the best estimate for the calculated $\text{EC}_{\text{ground}}/\beta^+$ for ^{40}K yields a partial decay constant for ^{40}K direct to ground-state ^{40}Ar of $11.6 \pm 1.5 \times 10^{-13} \text{ a}^{-1}$ (2σ). We calculate a partial decay constant of ^{40}K to ^{40}Ar of $0.592 \pm 0.014 \times 10^{-10} \text{ a}^{-1}$ and a total decay constant of $5.475 \pm 0.107 \times 10^{-10} \text{ a}^{-1}$ (2σ), and we conclude that although omission of this decay mode can be significant for K-Ar dating, it is minor for $^{40}\text{Ar}/^{39}\text{Ar}$ geochronology and

is therefore unlikely to have significantly biased published measurements.

1 Introduction

^{40}K is a naturally occurring radioisotope of K with atomic abundance of 0.0117 % (Garner et al., 1975). ^{40}K undergoes a branched decay to ^{40}Ar and ^{40}Ca with a total half-life of ca. 1.3 Ga and is the basis of the K-Ca and the K-Ar geochronometers (Aldrich and Nier, 1948; Wasserburg and Hayden, 1955; Marshall and DePaolo, 1982). The K-Ar system is most often exploited using the variant $^{40}\text{Ar}/^{39}\text{Ar}$ method, wherein some of the ^{39}K in the sample is transmuted to ^{39}Ar by irradiation with fast neutrons, thereby allowing both the parent and the daughter nuclides to be measured as isotopes of Ar (Merrihue and Turner, 1966). The latter is widely used to date geological events that span Earth's history, from volcanic eruptions recorded in historical texts (e.g., Preece et al., 2018; Renne et al., 1997) to the earliest events in the solar system (e.g., Renne, 2000).

Advances in analytical precision have forced practitioners in geochronology to address systematic uncertainties that are inherent in all radioisotope dating methods, such as uncertainties in the measurement apparatus, prior assumptions made by the observer, or interference from environmental factors. For the K-Ar system, these uncertainties also include those that arise from imprecisely known decay rates of ^{40}K . In the geological literature, there have been two influential reviews of measurements of the ^{40}K decay rate.

Beckinsale and Gale (1969) provided the first comprehensive review of measured and predicted decay rates, which became the basis of the convention adopted by Steiger and Jäger (1977) used by the geochronological community for the next 20 years. Subsequently, Min et al. (2000) provided a more lengthy, critical review of available specific activity data determined by direct measurements of decay and updated the derived decay rates for newer physical constants. More recently, the ^{40}K decay parameters were estimated by Renne et al. (2010, 2011), and although direct measurements of the ^{40}K decay were incorporated into the estimate, it was heavily weighted to an intercomparison with ^{238}U decay. The decay rate determined by Renne et al. (2010, 2011), and the Min et al. (2000) decay rates are the most frequently used in $^{40}\text{Ar}/^{39}\text{Ar}$ geochronology. These evaluations, along with those from the nuclear physics community, have been summarized recently by Cresswell et al. (2018, 2019).

Despite decades of work and long-standing interest in ^{40}K decay, there remains uncertainty over the nature of the decay scheme. There is consensus that most ^{40}K decays by β^- to ^{40}Ca or by electron capture to ^{40}Ar via an excited state and that a small amount ($\sim 0.001\%$) of ^{40}K decays to ^{40}Ar via β^+ . The early but influential review of ^{40}K decay by Beckinsale and Gale (1969) included these decay modes and also included a prediction of a second electron capture decay directly to the ground state of ^{40}Ar that would add an additional $\sim 2\%$ to the rate of decay from ^{40}K to ^{40}Ar . Many workers in geochronology (e.g., those who use the Min et al., 2000, decay constants) have ignored this prediction, and some nuclear physics tabulations do not clearly include it (Endt, 1990; Audi et al., 2003). The influential review by Min et al. (2000) described this decay mode as “unverified” and having a “questionable” existence.

However, the putative electron capture to ground-state decay-mode decay constant is of the same order of magnitude as the uncertainties in the decay rate of ^{40}K to ^{40}Ar and therefore may be a non-negligible and potentially important part of the geochronological system. Here, we describe the theoretical basis of this predicted decay mode and demonstrate the robust nature of the prediction via an analogous calculation of ^{22}Na decay. We describe experiments that could be made to measure this decay mode and also identify observations from nuclear physics experiments that offer evidence for its existence. We conclude that the evidence for this decay mode is strong and, despite the large uncertainty, should be considered in evaluations of the ^{40}K decay rate.

2 Historical overview

At present, ^{40}K has three experimentally verified decay modes (Fig. 1).

1. β^- decay to ^{40}Ca , this mode can be verified by direct measurement of the β^- emission.

2. Electron capture to an excited isomer of ^{40}Ar , followed by decay to the ground state of ^{40}Ar accompanied by emission of a 1.46 MeV γ -ray. Hereafter, we denote this decay mode as EC*. This mode can be verified by direct measurement of the γ emission.
3. β^+ decay from the ground state of ^{40}K to the ground state of ^{40}Ar (Engelkemeir et al., 1962), this is a very small component of the total decay rate and has been verified by direct measurement of the β^+ emission.

In their paper reporting the measurement of β^+/β^- , Engelkemeir et al. (1962), through private correspondence with Brosi and Kettle, proposed that an electron capture mode that goes directly to ground-state ^{40}Ar also exists, with an electron capture to positron ratio of 155. This decay mode is hereafter denoted EC_{ground}. This decay mode has not been experimentally detected, in part because the measurement is much more difficult to make than the others. If it exists, it would contribute about 0.2% to the total decay rate of ^{40}K or about 2% to the ^{40}Ar branch.

The EC_{ground} decay mode was included in the review by Beckinsale and Gale (1969) and then subsequently in Steiger and Jäger (1977). This decay mode is also included in the widely used Evaluated Nuclear Structure Data File (ENSDF) and Decay Data Evaluation Project (DDEP) evaluations (Chen, 2017; Mougeot and Helmer, 2009, respectively). However, evaluations by Endt and Van der Leun (1973, 1978), Endt (1990), and Audi et al. (2003) do not explicitly include this decay mode, with Audi et al. (2003) giving a transition intensity that is the combined EC* and β^+ intensities. Min et al. (2000) have questioned its validity because there is no experimental verification and therefore do not include EC_{ground} in their estimates.

3 Why there must be an EC_{ground} decay mode

In nuclides that are too proton-rich and therefore radioactive, protons decay to correct this imbalance via two mechanisms. Either, (1) the nucleus undergoes electron capture wherein an orbital electron is captured by the nucleus, or (2) the nucleus emits a positron (β^+). Both processes are types of β decay and result in the transformation of a proton to a neutron to conserve charge, and they both also emit a neutrino in order to conserve lepton number and energy. These two processes are typically paired: coupled electron capture– β^+ is the second most abundant decay type on the chart of the nuclides, after β^- decay (Audi et al., 2003). The electron capture (EC) and positron (β^+) decay modes are linked because both processes have the same initial and final nuclear states.

β^+ decay is always accompanied by EC, but the converse is not always true (Bambynek et al., 1977). This is because β^+ decay, unlike EC, requires a minimum amount of energy (~ 1022 keV, equivalent to the combined rest masses of both a positron and an electron) in order to produce the emitted

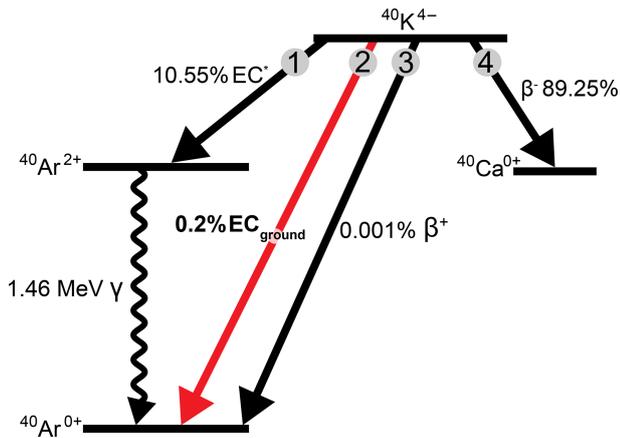


Figure 1. Decay scheme of ^{40}K following McDougall and Harrison (1999) and Leutz et al. (1965), where 1 is the electron capture branch to the excited state of ^{40}Ar with γ -ray emission (EC^*), 2 is the electron capture direct to the ground state of ^{40}Ar ($\text{EC}_{\text{ground}}$), 3 is the positron decay to ground state of ^{40}Ar , and 4 is the β^- decay to the ground state of ^{40}Ca . The disputed decay mode, $\text{EC}_{\text{ground}}$, is highlighted in red.

positron and an electron (the latter to satisfy charge conservation). The energy between initial and final states is the Q value, which is a function of the difference between the initial and final masses. This energy is shared between the outgoing neutrino, atomic excitation of the daughter system, recoil energy, and nuclear excitation of the daughter system. The energy that dictates if positron emission is possible is denoted Q_{EC} . In the decay of ^{40}K , the EC^* branch has an energy difference (Q_{EC}) between the initial and excited isomer state of only 44 keV. In contrast, the energy difference between ^{40}K and the ground state of ^{40}Ar is 1504.4 keV (Wang et al., 2017), an energy greater than the combined rest masses of the positron and electron. Therefore, the EC^* branch, with an energy difference of only 44 keV, cannot be the complement to the β^+ decay, and the $\text{EC}_{\text{ground}}$ must exist to provide the β^+ complement. The experimental observation of the β^+ decay mode comes from a single measurement by Engelkemeir et al. (1962). We rely on this measurement to make our argument for the existence of the $\text{EC}_{\text{ground}}$. It is possible that the positrons observed arise from the pair production of the ~ 1460 keV gamma; Engelkemeir et al. (1962) discuss this possibility in their experiment, calculating this as 55%–60% of the total positron detection rate. However, positrons produced by this mode of pair production would be monoenergetic at 440 keV, whereas the observed positron energy spectra exceed this value, with a maximum of 491 keV. A reasonable fit is also observed between the measured β^+ energy spectrum and the theoretical third forbidden unique energy spectrum, supporting the argument that these β^+ are from a decay mode rather than arising from pair production.

4 Theory and calculation of $\text{EC}_{\text{ground}}/\beta^+$

In the decay of ^{40}K , the nuclide can reach a more stable state (^{40}Ca or ^{40}Ar) only by violating quantum selection rules. Quantum selection rules place formal constraints of the possible transitions of a system from one quantum state to another. In this case it places constraints on the possible set of transitions from the parent ^{40}K state to the daughter ^{40}Ar . Decays which violate these selection rules undergo slow, so-called “forbidden” unique transitions, which give ^{40}K its long ~ 1.3 Ga half-life. The ^{40}K decay scheme itself is unusual because the coupled $\text{EC}_{\text{ground}}-\beta^+$ and β^- branches are the only third-order unique forbidden transitions known in nature. All ^{40}K decays undergo a parity reversal (where parity reversal is the change of sign in one of the spatial coordinates (x, y, z)) between the initial parent state and final daughter state. Therefore, we can define the selection rules as follows:

$$(|\Delta J - 1|)\text{st order unique forbidden decay,}$$

where $\Delta J = J_i - J_f$, is the change in spin from initial to final state following Krane and Halliday (1987). We can characterize each decay mode of ^{40}K by its degree of forbiddenness from the above selection rule. The EC^* mode undergoes a spin change of $\Delta J = 4 - 2 = 2$ and is classified as a first-order unique forbidden decay. The three other decay modes of ^{40}K , including $\text{EC}_{\text{ground}}$, all undergo a spin change of $\Delta J = 4 - 0 = 4$ and are classified as third-order unique forbidden decays.

The EC process occurs because the atomic electrons have a finite probability to be in the nucleus with the likelihood of being captured highest for those closest to the nucleus. A theoretical description of β emission was first given by Fermi (1934), while the possibility of electron capture was first recognized by Yukawa and Sakata (1935) and later developed by Bethe and Bacher (1936). Here we use Fermi theory of β^- decay to calculate the $\text{EC}_{\text{ground}}/\beta^+$ in the decay of ^{40}K .

We can use the ratio of orbital electron capture and positron emission to infer the existence of $\text{EC}_{\text{ground}}$. The ratio is defined as follows:

$$\text{br} = \frac{\lambda_{\text{ec}}}{\lambda_{\beta^+}}, \tag{1}$$

where λ_{ec} and λ_{β^+} are the probability per unit time of electron capture or β^+ emission. In electron capture, orbital electrons can be captured from any orbital shell of the atom. The EC/β^+ is therefore the summation of the individual capture ratios from each shell. Following Bambynek et al. (1977), the total electron capture-to-positron ratio is as follows:

$$\frac{\lambda_x}{\lambda_{\beta^+}} = \frac{\sum_x n_x C_x f_x}{f_{\beta^+} C(W)}, \tag{2}$$

where x is the shell, n_x is the relative occupation number, C_x contains the dependence of electron capture rates on nu-

clear structure giving the forbiddenness classification, similar to the shape factor in β decay (Emery, 1975), f_x is the integrated fermi function in β decay, f_{β^+} is the integrated positron spectrum, and $C(W)$ is the theoretical shape factor for allowed or forbidden transitions. A review of shape factors for ⁴⁰K transitions is provided by Cresswell et al. (2018, 2019). We initially simplify this equation to only consider the innermost K shell, the shell containing the electron with the highest probability to be captured by the nucleus:

$$\frac{\lambda_K}{\lambda_{\beta^+}} = \frac{n_K C_K f_K}{f_{\beta^+} C(W)}, \tag{3}$$

where λ_K is the probability of K shell capture. For this capture, f_K is defined as follows:

$$f_K = \frac{\pi}{2} q_K^2 \beta_K^2 B_K^K, \tag{4}$$

where q_K is the momentum of the neutrino particle, β_K is the Coulomb amplitude of the wave function, and B_K is the term for overlap and exchange corrections. Similarly, f_{β^+} is defined as follows:

$$f_{\beta^+} = \int_1^{W_0} F(-Z, W) W p(W_0 - W)^2 dW, \tag{5}$$

$$W = 1 + \frac{E_T}{m_e}, \tag{6}$$

$$W_0 = 1 + \frac{E_{\max}}{m_e}, \tag{7}$$

$$p = \sqrt{W^2 - 1}, \tag{8}$$

where W is the total energy of the positron given by its kinetic energy E_T and rest mass m_e , defined above, and the momentum of the positron is given by p (Eq. 8); W_0 is the total normalized energy defined above; E_{\max} is the upper limit of the positron energy (equal to the Q value of the decay); and $F(-Z, W)$ is the Fermi function. We follow Bambynek et al. (1977) in the formula for $\frac{C_K}{C(\bar{W})}$ which is given by

$$\frac{C_K}{C(\bar{W})} = [(2L - 1)!]^{-1} q_K^{2(L-1)} \left\{ \sum_{n=1}^L \lambda_n p^{2(n-1)} ((2n - 1)! \cdot [2(L - n) + 1]!)^{-1} \right\}^{-1}, \tag{9}$$

where $L = \Delta J$ and $L = 1$ for $\Delta J = 0$. The parameter λ_n cannot be calculated in a straightforward manner and therefore we follow a typical assumption that $\lambda_n = 1$ (Huber, 2011). This reduces the above expression to

$$\frac{C_K}{C(\bar{W})} = \frac{q_K^6}{p^6 + q^6 + 7p^2 q^2 (p^2 + q^2)}. \tag{10}$$

In a given decay, the change in charge from the initial to final state can lead to an imperfect overlap of the wave functions of these states. Furthermore, given the indistinguishability of electrons, there is the possibility of an exchange effect wherein an electron does not necessarily come from the orbital where the vacancy appears. For instance, it is possible that a vacancy may appear in the K-shell, but the captured electron from an outer shell is then subsequently filled by the inner shell electron (Bahcall, 1962; Bambynek et al., 1977). We follow Bahcall (1962) in implementing corrections for these effects, resulting in $B_K = 0.979$. Following this, using nuclear data given in Bambynek et al. (1977), we estimate an $EC_{\text{ground}}/\beta^+$ of 148.

We first note that this value is in approximate concordance with the private correspondence value in Engelkemeir et al. (1962). However, this is only the capture ratio from the K shell, so we extend our model to a total electron capture ratio from all orbitals following Bosch et al. (1977). The total electron capture to positron ratio, an extension of Eq. (1), is given by

$$\frac{EC}{\beta^+} = \frac{K}{\beta^+} \left(1 + \frac{L}{K} + \frac{M}{L} \frac{L}{K} + \dots \right). \tag{11}$$

We can simplify this equation by neglecting shells that make a negligible contribution. In ⁴⁰K the probability of capture is dominated by the two inner shells K and L1, with approximate probability of ca. $\sim 90\%$ and $\sim 10\%$ and a negligible contribution from the shells further out. We can therefore omit all shell captures except K and L1 to arrive at the total $EC_{\text{ground}}/\beta^+$ ratio:

$$\frac{EC}{\beta^+} = \frac{K}{\beta^+} \left(1 + \frac{L_1}{K} \right). \tag{12}$$

The ratio of each shell capture can be solved with the following equation:

$$\frac{x}{K} = \frac{\beta_x^2 (W_0 - W_x)^2 B_x}{\beta_K^2 (W_0 - W_x)^2 B_K}, \tag{13}$$

where $x = L1$ and the other symbols have the same definition as above. Using this equation we calculate a total $EC_{\text{ground}}/\beta^+$ of 164.

To further estimate the magnitude of the electron capture decay mode, we can perform another calculation of $EC_{\text{ground}}/\beta^+$ following Fireman (1949). This simplified form of the calculating $EC_{\text{ground}}/\beta^+$ is dependent only on the Q value (the difference between the initial and final state energies). This is given by

$$\frac{\lambda_{EC}}{\lambda_{\beta^+}} = \frac{(\eta + 2)^8}{0.450\eta^{6.5}} \cdot \frac{1}{0.0676 + 1.25\eta + 8.48\eta^2 + 12.5\eta^3 + 1.74\eta^4 + 0.079\eta^6}, \tag{14}$$

where $\eta = \frac{Q}{m_e} - 2$. We calculate an $\text{EC}_{\text{ground}}/\beta^+$ of 272 using this method and the updated Q value of Wang et al. (2017). We note that despite discrepancies in these values for each method of evaluation, they are of the same order of magnitude. The differences in the values in these evaluations highlight the need for experimental measurement of $\text{EC}_{\text{ground}}/\beta^+$.

5 Comparison with other evaluations

Other theoretical evaluations of $\text{EC}_{\text{ground}}/\beta^+$ for ^{40}K exist in the literature (Fig. 2). Pradler et al. (2013) and Mougeot (2018) report ratios of 150 ± 0.15 and 212 ± 0.15 , respectively (uncertainties are reported where they have been estimated). These workers use broadly similar methods to us. Mougeot (2018) uses higher-order corrections for both exchange and overlap and accounts for the dependence of λ_{K} , that we set equal to 1 in Eq. (9) using the energy of the decay. Pradler et al. (2013) use the Fermi method and data from Bambynek et al. (1977) but only perform the calculation for K-shell electrons, resulting in a slightly differently calculated value than we report. Notably, all estimated values are of the same order of magnitude, similar to the ratio 155 reported in Engelkemier et al. (1962) and our calculated value of 164. Currently, the most commonly used $\text{EC}_{\text{ground}}/\beta^+$ value is calculated via the LogFT program, a program used in nuclear data evaluations (ENSDF collaboration, LogFT). However, the program is capable of only calculating first and second unique forbidden decay ratios, so the $\text{EC}_{\text{ground}}/\beta^+$ value from LogFT of 200 ± 100 is an extrapolation, with the assumption that the increase in the ratio from second to third order is by the same factor as the increase from first to second order. Finally, Chen (2017) evaluates the ^{40}K decay data and reports a $\text{EC}_{\text{ground}}/\beta^+$ value of 45.2 ± 1.4 without elaboration.

The variability between the modern estimates are driven primarily by choices when making the approximations necessary for these calculations to be tractable. Uncertainties on individual estimates that could be derived by propagating the uncertainties in the underlying experimental data are small, and where uncertainties are estimated, they are generally not explicated.

The value calculated by Mougeot (2018) of 212 ± 0.15 is currently the best estimate of the ^{40}K $\text{EC}_{\text{ground}}/\beta^+$. It is slightly higher than two other recent estimates, our value of 164 or that of Pradler et al. (2013) of 150. Given a broad consensus in calculated $\text{EC}_{\text{ground}}/\beta^+$ over several decades and via a variety of methods, it appears highly likely that it falls in the range 150–212 (Fig. 2).

6 Comparison with ^{22}Na

To test the validity of our ^{40}K $\text{EC}_{\text{ground}}/\beta^+$ estimate, we use the same calculations to estimate the experimentally constrained $(\text{EC}/\beta^+)^*$ value for ^{22}Na decay. ^{22}Na is radionu-

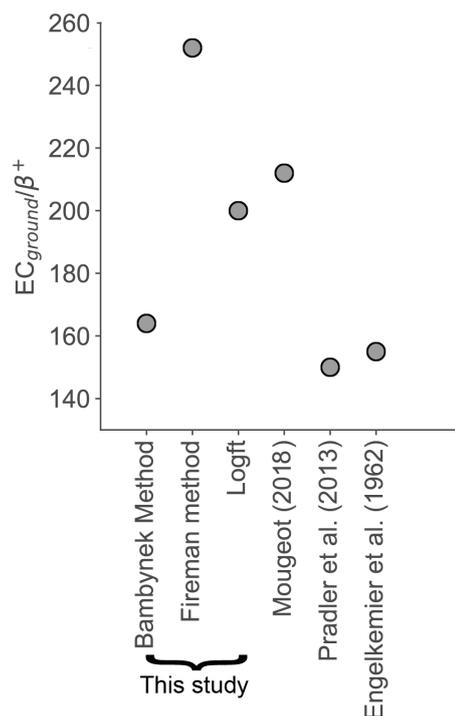


Figure 2. Comparison of theoretically calculated $\text{EC}_{\text{ground}}/\beta^+$ of ^{40}K in this study using three methods: (1) Bambynek method (Bambynek et al., 1977), (2) Fireman method (Fireman, 1949), and (3) LogFT (ENSDF collaboration, 2020, LogFT). The value of Chen (2017) is not included in the figure as it is an extreme outlier, and the authors do not explain the method they use to reach this value. Our calculated ratios are compared to previous evaluations in the literature (Engelkemier et al., 1962; Pradler et al., 2013; Mougeot, 2018). Uncertainties in these values are either intractable or, in the case of Mougeot (2018), too small to plot. Note the consistency in the estimated ratio from all of the methods. Uncertainties are unknown in all cases except Mougeot (2018), where the uncertainty is too small to plot.

clide with a half-life of ~ 2.6 years, it occurs in nature as a low-abundance cosmogenic nuclide produced by spallation of ^{40}Ar and is also produced synthetically by proton irradiation for use in positron emission tomography. Like ^{40}K , it decays by electron capture and positron emission. The main $\text{EC}-\beta^+$ pair for ^{22}Na decays initially to the excited state of ^{22}Ne , followed by a 1.27 MeV γ emission (Fig. 3; Bé et al., 2010). This pair has a $(\text{EC}/\beta^+)^*$ of approximately 0.1 and accounts for $> 99.9\%$ of the total decay. A second $\text{EC}-\beta^+$ pair decays directly to the ground state of ^{22}Na with an $(\text{EC}/\beta^+)_{\text{ground}}$ of ~ 0.02 , but it is a minor component. Here, we calculate the $(\text{EC}/\beta^+)^*$ for the main branch. Unlike ^{40}K , the dominant decay of ^{22}Na is the β^+ decay mode. This is due to the greater difference in energy between the initial and final states, as positron decay will have a greater possibility of occurring in decays with a greater mass differences between initial and final states (Emery, 1975). ^{22}Na is not a perfect analogue; however, it is probably the best choice that

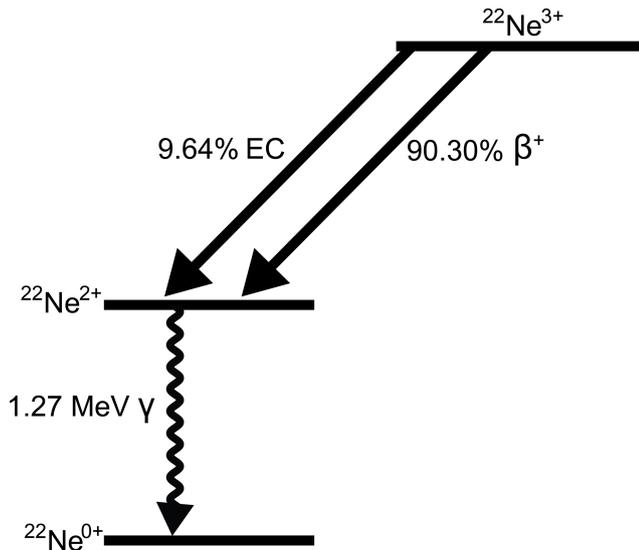


Figure 3. Decay scheme of ^{22}Na following Bé et al. (2010) and Leutz et al. (1965). An additional EC and β^+ decay pair that corresponds to approximately 0.056% of the total decay of ^{22}Na has been omitted for clarity.

has both a tractable theoretical calculation and a wealth of experimental data that can be used readily for verification.

Unlike ^{40}K , there are numerous measurements of the electron capture to positron ratio for decay to the excited state of ^{22}Ne (Fig. 4; Kreger, 1954; Vatai et al., 1968; Williams, 1964; McCann and Smith, 1969; MacMahon and Baerg, 1976; Bosch et al., 1977; Baerg, 1983; Schmidt-Ott et al., 1984; Sýkora and Povinec, 1986; Kunze et al., 1990; Nähle et al., 2008). Measurement of $(\text{EC}/\beta^+)^*$ for ^{22}Na is accomplished by measurement of both of the gammas (which come from both the EC* and the β^{+*}) and the x-rays (which only come from the EC branch). Relative to the ^{40}K $\text{EC}_{\text{ground}}/\beta^+$, the ^{22}Na $(\text{EC}/\beta^+)^*$ is a more straightforward measurement because of the higher activity (meaning higher count rate) and the higher energy of the x-ray emitted from the Auger electron, which an electron from the same atom that is emitted as a vacancy of an inner shell is filled. In a decay to the excited state of ^{22}Ne , the de-excitation 1.28 MeV γ will be associated with both electron capture and positron decay. However, those measured 1.28 MeV γ that are not accompanied by two 0.511 MeV annihilation photons can be used to distinguish between both processes. We use the experimental measurements to verify our calculations described above for ^{40}K .

Following a similar calculation using the Fermi method, our preferred method, to that used for our proposed estimate of the ^{40}K $\text{EC}_{\text{ground}}/\beta^+$, we estimate an $(\text{EC}/\beta^+)^*$ of approximately 0.11. This is within the range of measured values of 0.105–0.115 (Fig. 4), suggesting that our calculation strategy of the ^{40}K $\text{EC}_{\text{ground}}/\beta^+$ is accurate, and lends fur-

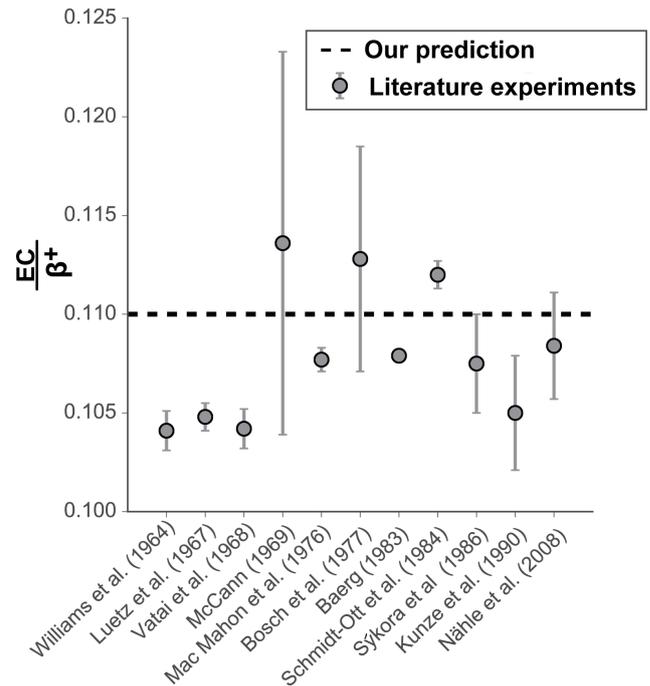


Figure 4. Comparison of experimentally measured $(\text{EC}/\beta^+)^*$ ratios of ^{22}Na (grey circles) adapted from Kunze et al. (1990) with our calculated value (dashed black line). Note the concordance of the theoretical and experimental determinations. The uncertainty in the Baerg (1983) determination is smaller than the symbol.

ther confidence to the existence of the current unmeasured ^{40}K electron capture to ground-state decay.

7 Experimental verification of $\text{EC}_{\text{ground}}$ decay mode

In both β^- and β^+ decay, an electron or positron is emitted that allows for direct detection and verification of the decay process. In contrast, electron capture cannot be detected directly. Methods to experimentally verify electron capture rely on indirect processes associated with the rearrangement of the atom following the capture of the orbital electron. Once the electron is captured the atom will rearrange itself to fill the vacancy, resulting in the emission of a characteristic x-ray or Auger electron with an energy defined by the binding energy of the shell vacancy of the daughter nucleus.

In the case of ^{40}K , verification of the $\text{EC}_{\text{ground}}$ decay can be achieved by measuring the characteristic x-rays (Di Stefano et al., 2017). The orbital electron with the highest probability of capture is from the K-shell; if this electron is captured, the resulting vacancy in the K-shell may be filled by an electron from any of the other shells (e.g., L, M), and a characteristic x-ray is emitted with an energy dependent on the particular shell that fills the vacancy. It is not necessary, however, that all K-capture processes result in the emission of an x-ray. By the Auger effect, a radiationless transfer may occur wherein the K-shell vacancy is replaced by two

vacancies in the next outer shell, L, or one in the next two shells; L and M. The energies of the Auger electrons emitted in these transitions depend upon the ^{40}Ar product resulting from K-capture. Both electron capture decays to the ground and excited state of ^{40}Ar ($^{40}\text{Ar}^{2+}$) result in the same electron configuration and x-ray emissions. Di Stefano et al. (2017) suggested tagging x-rays with the de-excitation γ associated with electron capture to $^{40}\text{Ar}^{2+}$, which has a lifetime on the order of $\sim 10^{-12}$ s (Di Stefano et al., 2017). Measuring these tagged x-rays experimentally will be challenging since it requires identifying a low-probability decay mode with x-ray signals present against a high background from the $^{40}\text{Ar}^{2+}$ state. Further, as illustrated in Di Stefano et al. (2020), it is expected that 50 EC* decays occur for every 1 EC_{ground} decay; therefore, a detector efficiency of $\geq 98\%$ is required to make sure that there is fewer than one mis-tagged EC* decay for each true EC_{ground} decay. The experiment therefore requires an x-ray spectrometer able to resolve the Ar-K x-ray from other x-rays in the background and accurately account for the x-ray- γ -ray coincidence efficiency ($\geq 98\%$) to quantify x-ray emission rates in excess of those from the $^{40}\text{Ar}^{2+}$ state. Given the complexity involved in this experiment, a pilot study was conducted at SUERC to measure characteristic x-rays from a KCl source. The experiment was not successful because the detector was not able to resolve the Ar-K x-ray sufficiently but demonstrates the potential of this method to detect the x-rays, given a sufficiently high-resolution detector. Full details are provided in the Supplement.

Ongoing attempts are being made to verify this decay mode by careful detection of the characteristic x-rays by the Potassium decay (KDK) experiment (Di Stefano et al., 2017; Stukel, 2018). Experimental verification has implications for (1) rare event physics, as it is a vital component in constraining the irreducible background and verifying results in the DArk MATter (DAMA) experiment (Pradler et al., 2013); (2) the theory of β decay (Fermi, 1934), as it is the only third-order unique forbidden electron capture decay known (Audi et al., 2003); and (3) K-Ar and $^{40}\text{Ar}/^{39}\text{Ar}$ geochronology, for which it is currently overlooked due to lack of experimental evidence. We further expand on the implications for geochronology below.

8 Relevance for geochronology

Geochronology with the K-Ar system requires either both the branching ratio and the total decay constant, or in the case of an $^{40}\text{Ar}/^{39}\text{Ar}$ age wherein the fluence monitor age is constrained independently of its K-Ar systematics (Merrihue and Turner, 1966), only the total decay constant. Using lower and upper bound values of $\text{EC}_{\text{ground}}/\beta^+$ corresponding to 150 and 212 as described above, the decay constants calculated by Min et al. (2000) ($\lambda_{\text{EC}^*} = 0.580 \pm 0.014 \times 10^{-10} \text{ a}^{-1}$ and $\lambda_T = 5.463 \pm 0.107 \times 10^{-10} \text{ a}^{-1}$), and the β^+/β^- from Engelkemeir et al. (1962) ($1.12 \pm 0.14 \times 10^{-5}$), we calculate a

β^+ decay constant of $5.47 \pm 0.69 \times 10^{-15} \text{ a}^{-1}$ and a range of $\text{EC}_{\text{ground}}$ decay constants of $8.2\text{--}11.6 \times 10^{-13} \text{ a}^{-1}$. Combining these values with the Min et al. (2000) quantities yields a new partial decay constant for ^{40}K to ^{40}Ar ($\lambda_{40\text{Ar}}$) that ranges from $0.588\text{--}0.592 \times 10^{-10} \text{ a}^{-1}$ and total decay constant (λ_T) that ranges from $5.471\text{--}5.475 \times 10^{-10} \text{ a}^{-1}$. These ranges are within the uncertainties calculated by Min et al. (2000) for decay constants that do not include the $\text{EC}_{\text{ground}}$ decay mode. Existing and modified constraints on the decay modes are given in Table 1. Our preferred decay constants are those calculated with the $\text{EC}_{\text{ground}}/\beta^+ = 212$ from Mougeot (2018).

Consequently, K-Ar (and $^{40}\text{Ar}/^{39}\text{Ar}$) ages calculated with these new decay constants will be younger than those calculated using the Min et al. (2000) decay constants. K-Ar dates are most sensitive to shifts in the decay constant because they incorporate the branching ratio, which is more strongly affected than the total ^{40}K decay constant. K-Ar ages will decrease by $\sim 1.4\%\text{--}2.0\%$ at 1 Ma, $\sim 1.1\%\text{--}1.5\%$ at 1 Ga, and $\sim 0.6\%\text{--}0.8\%$ at 4.5 Ga (Fig. 5). Ages determined using the $^{40}\text{Ar}/^{39}\text{Ar}$ method, for which the fluence monitor age is independently constrained (e.g., Kuiper et al., 2008; Rivera et al., 2011), are much less sensitive to the change in decay constant. Using Eq. (5) from Renne et al. (1998), and assuming calibration to a monitor with an age of 28.2 Ma, ages < 28 Ma increase only slightly, by $< 0.002\%$. There is no age difference at 28.2 Ma, the fluence monitor age. Ages then decrease for ages > 28.2 Ma, with ages decreased by $0.07\%\text{--}0.10\%$ at 2.5 Ga, and by $0.09\%\text{--}0.13\%$ at 4.5 Ga (Fig. 5).

The age of fluence monitors such as the Fish Canyon tuff sanidine (e.g., Morgan et al., 2014) determined by intercomparison with astronomically tuned ages of ash beds (Kuiper et al., 2008; Rivera et al., 2011) is also sensitive to revision of decay constants. Using the data published by Kuiper et al. (2008), and incorporating an $\text{EC}_{\text{ground}}$ decay mode, we calculate a new age for Fish Canyon sanidine of 28.200 ± 0.044 Ma, nominally lower but indistinguishable from the published value of 28.201 ± 0.044 Ma. Overall, the effects of an $\text{EC}_{\text{ground}}$ decay mode are unlikely to be significant for most current applications of $^{40}\text{Ar}/^{39}\text{Ar}$ geochronology. However, given the levels of analytical precision attainable by the K-Ar dating approach when dating geologically recent materials by K-Ar (e.g., Altherr et al., 2019), the $\text{EC}_{\text{ground}}$ decay mode will impact the accuracy of this chronometer.

9 Conclusion

The Fermi theory of β decay has decades of experimental support and is well established. We demonstrate this here by using these theories to accurately calculate the decay rate of a ^{22}Na , a nuclide with an experimentally verified decay rate. We have used this information to demonstrate the high likelihood that the suspected second electron capture decay mode of ^{40}K exists. Based on the cal-

Table 1. Evaluations of decay mode branches and total decay constant used in age determination. $\lambda_{40\text{Ar}}$ is the partial decay constant for the ^{40}Ar branch, including both the EC^* and $\text{EC}_{\text{ground}}$ components. Uncertainties from the β^+/β^- and $\text{EC}_{\text{ground}}/\beta^+$ do not substantially change the uncertainties in $\lambda_{40\text{Ar}}$ or λ_T .

Parameter	Value $\pm 2\sigma$	Relative unc. (%)	References
Previous values			
λ_{EC^*}	$0.580 \pm 0.014 \times 10^{-10} \text{ a}^{-1}$	2.4	Min et al. (2000)
λ_T	$5.463 \pm 0.107 \times 10^{-10} \text{ a}^{-1}$	2.0	Min et al. (2000)
λ_{β^+}	$5.47 \pm 0.69 \times 10^{-15} \text{ a}^{-1}$	13	Engelkemeir et al. (1962)
Modified values with lower-bound $\text{EC}_{\text{ground}}/\beta^+ = 150$			
$\lambda_{\text{EC}_{\text{ground}}}$	$8.2 \pm 1.0 \times 10^{-13} \text{ a}^{-1}$	13	This work
$\lambda_{40\text{Ar}}$	$0.588 \pm 0.014 \times 10^{-10} \text{ a}^{-1}$	2.4	This work
λ_T	$5.471 \pm 0.107 \times 10^{-10} \text{ a}^{-1}$	2.0	This work
Modified values with upper-bound $\text{EC}_{\text{ground}}/\beta^+ = 212$			
$\lambda_{\text{EC}_{\text{ground}}}$	$11.6 \pm 1.5 \times 10^{-13} \text{ a}^{-1}$	13	This work
$\lambda_{40\text{Ar}}$	$0.592 \pm 0.014 \times 10^{-10} \text{ a}^{-1}$	2.4	This work
λ_T	$5.475 \pm 0.107 \times 10^{-10} \text{ a}^{-1}$	2.0	This work

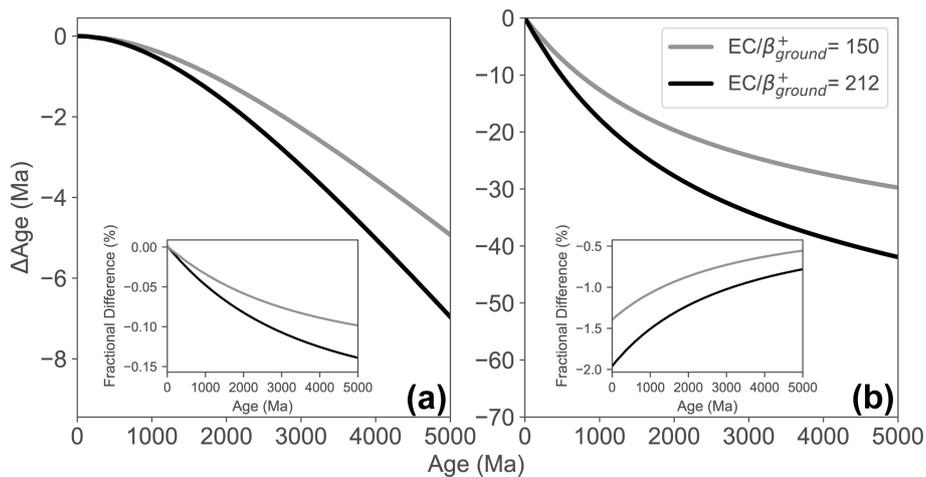


Figure 5. Change in age, ΔAge , is the age of a given sample with the decay mode to ground state included, subtracted from the age with the decay mode to ground state omitted. Panel (a) shows the change in age using the $^{40}\text{Ar}/^{39}\text{Ar}$ equation with independently calibrated standards using both the lower bound ($\text{EC}/\beta^+ = 150$; grey) and upper bound ($\text{EC}/\beta^+ = 212$; black). Panel (b) shows the change in age using the K-Ar equation using both the lower bound ($\text{EC}/\beta^+ = 150$; grey) and upper bound ($\text{EC}/\beta^+ = 212$; black). Inset figures in each panel show the fractional differences in age by the inclusion of both the upper- and lower-bound EC/β^+ value. The larger difference in ages for the K-Ar system is due to the dependence on both the total decay constant and branching ratio.

culations of Mougeot (2018), the best estimate of the partial decay constant for ^{40}K direct to ground-state ^{40}Ar is $11.6 \pm 1.5 \times 10^{-13} \text{ a}^{-1}$ (2σ), and other calculations are no lower than about $8.2 \times 10^{-13} \text{ a}^{-1}$. Combining this with the decay constants published by Min et al. (2000) results in revised values of $\lambda_{40\text{Ar}} = 0.592 \pm 0.014 \times 10^{-10} \text{ a}^{-1}$ and $\lambda_T = 5.475 \pm 0.107 \times 10^{-10} \text{ a}^{-1}$. This addresses a long-standing question in K-Ar and $^{40}\text{Ar}/^{39}\text{Ar}$ geochronology and provides future workers with confidence that the ^{40}K $\text{EC}_{\text{ground}}$ decay mode exists. Just as important as providing support for its ex-

istence, we also demonstrate that the magnitude of this decay mode is small enough that neglecting it has not yet resulted in significantly biased geochronological $^{40}\text{Ar}/^{39}\text{Ar}$ data. The same cannot be stated for the K-Ar dating approach, especially for geologically young materials.

Despite the strong grounding in theory, the $\text{EC}_{\text{ground}}$ decay mode has yet to be detected. The next step is experimental verification to determine the branching ratio. This will allow for a more complete evaluation of uncertainties associated

with the decay mode and the branching ratio. This experiment is difficult but not intractable.

Data availability. All data used are available in the Supplement and can be accessed publicly via <https://doi.org/10.6084/m9.figshare.13280795> (Carter et al., 2020).

Supplement. The supplement related to this article is available online at: <https://doi.org/10.5194/gchron-2-355-2020-supplement>.

Author contributions. The study was conceived by JC and RBI. JC, AJC, and DCWS calculated the ratio of electron capture to beta activities and measured x-rays. JC, RBI, DFM and MMT calculated the effects for geochronology. JC wrote the manuscript with contributions from all authors.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. Jack Carter studentship funded by the UK Space Agency (grant no. ST/P001289/1). NERC are thanked for continued funding of the National Environmental Isotope Facility (Ar/Ar laboratory) at SUERC NE/S011587/1. Marissa M. Tremblay acknowledges The Royal Society (NF171365) for funding. Ryan B. Ickert thanks the members of the Geochron Club for discussion. We thank two anonymous reviewers who helped clarify and focus the manuscript and Clare Warren for editorial handling.

Financial support. This research has been supported by the UK Space Agency (grant no. ST/P001289/1).

Review statement. This paper was edited by Clare Warren and reviewed by two anonymous referees.

References

Aldrich, L. T. and Nier, A. O.: Argon 40 in potassium minerals, *Phys. Rev.*, 74, 876–877, <https://doi.org/10.1103/PhysRev.74.876>, 1948.

Altherr, R., Mertz-Kraus, R., Volker, F., Kreuzer, H., Henjes-Kunst, F., and Lange, U.: Geodynamic setting of Upper Miocene to Quaternary alkaline basalts from Harrat al ‘Uwayrid (NW Saudi Arabia): Constraints from KAr dating, chemical and Sr-Nd-Pb isotope compositions, and petrological modelling, *Lithos*, 330, 120–138, <https://doi.org/10.1016/j.lithos.2019.02.007>, 2019.

Audi, G., Bersillon, O., Blachot, J., and Wapstra, A. H.: The NUBASE evaluation of nuclear and decay properties, *Nucl. Phys. A*, 729, 3–128, <https://doi.org/10.1016/j.nuclphysa.2003.11.001>, 2003.

Baerg, A. P.: Electron capture to positron branching ratios in the decay of ^{22}Na and ^{44}Sc , *Can. J. Phys.*, 61, 1222–1226, <https://doi.org/10.1139/p83-155>, 1983.

Bahcall, J. N.: Electron Capture and Nuclear Matrix Elements of Be^7 , *Phys. Rev.*, 128, 1297–1301, <https://doi.org/10.1201/9780429502811-65>, 1962.

Bambynek, W., Behrens, H., Chen, M. H., Crasemann, B., Fitzpatrick, M. L., Ledingham, K. W. D., Genz, H., Mutterer, M., and Intemann, R. L.: Orbital electron capture by the nucleus, *Rev. Mod. Phys.*, 49, 77–221, <https://doi.org/10.1103/RevModPhys.49.77>, 1977.

Bé, M.-M., Chisté, V., Dulieu, C., Mougeot, X., Browne, E., Chechev, V., Kuzmenko, N., Kondev, F., Luca, A., and Galan, M.: Table of radionuclides (Vol. 5-A = 22 to 244), *Mongraphie BIPM-5*, available at: <https://hal-cea.archives-ouvertes.fr/cea-02476352/document> (last access: 12 November 2020), 2010.

Beckinsale, R. D. and Gale, N. H.: A reappraisal of the decay constants and branching ratio of ^{40}K , *Earth Planet. Sc. Lett.*, 6, 289–294, [https://doi.org/10.1016/0012-821X\(69\)90170-8](https://doi.org/10.1016/0012-821X(69)90170-8), 1969.

Bethe, H. A. and Bacher, R. F.: Nuclear physics A. Stationary states of nuclei, *Rev. Mod. Phys.*, 8, 82–229, <https://doi.org/10.1103/RevModPhys.8.82>, 1936.

Bosch, H. E., Davidson, J., Davidson, M., and Szybisz, L.: The electron capture to positron emission ratios in the decay of ^{22}Na and ^{65}Zn , *Z. Phys. A*, 280, 321–327, <https://doi.org/10.1007/BF01435440>, 1977.

Carter, J., Ickert, R., Mark, D., Tremblay, M. M., Cresswell, A., and Sanderson, D.: Potassium counting experiment.csv, figshare, Dataset, <https://doi.org/10.6084/m9.figshare.13280795.v1>, 2020.

Chen, J.: Nuclear Data Sheets for $A = 40$, *Nuclear Data Sheets*, 140, 1–376, <https://doi.org/10.1016/j.nds.2017.02.001>, 2017.

Cresswell, A. J., Carter, J., and Sanderson, D. C. W.: Dose rate conversion parameters: Assessment of nuclear data, *Radiat. Meas.*, 120, 195–201, <https://doi.org/10.1016/j.radmeas.2018.02.007>, 2018.

Cresswell, A. J., Sanderson, D. C. W., and Carter, J.: Review of nuclear data for naturally occurring radionuclides applied to environmental applications, *Eur. Phys. J. Plus*, 134, 69, <https://doi.org/10.1140/epjp/i2019-12437-1>, 2019.

Di Stefano, P. C. F., Brewer, N., Fijałkowska, A., Gai, Z., Goetz, K. C., Grzywacz, R., Hamm, D., Lechner, P., Liu, Y., Lukosi, E., and Mancuso, M.: The KDK (potassium decay) experiment, *arXiv [preprint]*, arXiv:1711.04004, 10 November 2017.

Di Stefano, P. C. F., Brewer, N., Fijałkowska, A., Gai, Z., Goetz, K. C., Grzywacz, R., Hamm, D., Lechner, P., Liu, Y., Lukosi, E., and Mancuso, M.: The KDK (potassium decay) experiment, *J. Phys. Conf. Ser.*, 1342, 012062, IOP Publishing, <https://doi.org/10.1088/1742-6596/1342/1/012062>, 2020.

Emery, G. T.: Ionization through Nuclear Electron Capture and Internal Conversion, in: *Atomic Inner-Shell Processes*, Academic Press, New York, 201–231, <https://doi.org/10.1016/B978-0-12-196901-1.50010-8>, 1975.

Engelkemeir, D. W., Flynn, K. F., and Glendenin, L. E.: Positron Emission in the Decay of K^{40} , *Phys. Rev.*, 126, 1818–1822, <https://doi.org/10.1103/PhysRev.126.1818>, 1962.

Endt, P. M.: Energy levels of $A = 21$ –44 nuclei (VII), *Nucl. Phys. A*, 521, 1–400, [https://doi.org/10.1016/0375-9474\(90\)90598-G](https://doi.org/10.1016/0375-9474(90)90598-G), 1990.

- Endt, P. M. and Van der Leun, C.: Energy levels of $A = 21\text{--}44$ nuclei (V), Nucl. Phys. A, 214, 1–625, [https://doi.org/10.1016/0375-9474\(73\)91131-7](https://doi.org/10.1016/0375-9474(73)91131-7), 1973.
- Endt, P. M. and Van der Leun, C.: Energy levels of $A = 21\text{--}44$ nuclei (VI), Nucl. Phys. A, 310, 1–751, [https://doi.org/10.1016/0375-9474\(78\)90611-5](https://doi.org/10.1016/0375-9474(78)90611-5), 1978.
- ENSDF Collaboration: LOGFT, available at: https://www-nds.iaea.org/public/ensdf_pgm/, last access: 12 November 2020.
- Fermi, E.: Versuch einer Theorie der β -Strahlen. I, Z. Phys., 88, 161–177, <https://doi.org/10.1007/BF01351864>, 1934.
- Fireman, E. L.: On the Decay of K^{40} , Phys. Rev., 75, 1447, <https://doi.org/10.1103/PhysRev.75.1447.2>, 1949.
- Garner, E. L., Murphy, T. J., Gramlich, J. W., Paulsen, P. J., and Barnes, I. L.: Absolute isotopic abundance ratios and the atomic weight of a reference sample of potassium, J. Res. Natl. Bur. Stand. A Phys. Chem., 79A, 713–725, <https://doi.org/10.6028/jres.079A.028>, 1975.
- Huber, P.: Determination of antineutrino spectra from nuclear reactors, Phys. Rev. C, 84, 024617, <https://doi.org/10.1103/PhysRevC.84.024617>, 2011.
- Krane, K. S. and Halliday, D.: Introductory nuclear physics, 3rd edn., Wiley, New York, 845 pp., 1987.
- Kreger, W. E.: K Capture to positron ratio for Na^{22} , Phys. Rev., 96, 1554–1555, <https://doi.org/10.1103/PhysRev.96.1554>, 1954.
- Kuiper, K. F., Deino, A., Hilgen, F. J., Krijgsman, W., Renne, P. R., and Wijbrans, J. R.: Synchronizing rock clocks of Earth history, Science, 320, 500–504, <https://doi.org/10.1126/science.1154339>, 2008.
- Kunze, V., Schmidt-Ott, W. D., and Behrens, H.: Remeasurement of capture to positron decay ratios in ^{22}Na and ^{65}Zn and comparison with theory, Z. Phys. A, 337, 169–173, <https://doi.org/10.1007/BF01294288>, 1990.
- Leutz, H., Schulz, G., and Wenninger, H.: The decay of potassium-40, Z. Phys., 187, 151–164, <https://doi.org/10.1007/BF01387190>, 1965.
- MacMahon, T. D. and Baerg, A. P.: The electron capture to positron branching ratio in the decay of ^{22}Na , Can. J. Phys., 54, 1433–1437, <https://doi.org/10.1139/p76-168>, 1976.
- Marshall, B. D. and DePaolo, D. J.: Precise age determinations and petrogenetic studies using the K-Ca method, Geochim. Cosmochim. Ac., 46, 2537–2545, [https://doi.org/10.1016/0016-7037\(82\)90376-3](https://doi.org/10.1016/0016-7037(82)90376-3), 1982.
- McCann, M. F. and Smith, K. M.: Direct measurement of the K electron capture to positron emission ratio in the decay of ^{22}Na , J. Phys. A-Gen. Phys., 2, 392–397, <https://doi.org/10.1088/0305-4470/2/3/018>, 1969.
- McDougall, I. and Harrison, T. M.: Geochronology and Thermochronology by the $^{40}\text{Ar}/^{39}\text{Ar}$ Method, 2nd edn., Oxford University Press, Oxford, 269, 1999.
- Merrihue, C. and Turner, G.: Potassium-argon dating by activation with fast neutrons, J. Geophys. Res.-Sol. Ea., 71, 2852–2857, <https://doi.org/10.1029/JZ071i011p02852>, 1966.
- Min, K., Mundil, R., Renne, P. R., and Ludwig, K. R.: A test for systematic errors in $^{40}\text{Ar}/^{39}\text{Ar}$ geochronology through comparison with U/Pb analysis of a 1.1-Ga rhyolite, Geochim. Cosmochim. Ac., 64, 73–98, [https://doi.org/10.1016/S0016-7037\(99\)00204-5](https://doi.org/10.1016/S0016-7037(99)00204-5), 2000.
- Morgan, L. E., Mark, D. F., Imlach, J., Barfod, D., and Dymock, R.: FCs-EK: A new sampling of the Fish Canyon Tuff $^{40}\text{Ar}/^{39}\text{Ar}$ neutron flux monitor, Geol. Soc. Spec. Publ., 378, 63–67, <https://doi.org/10.1144/SP378.21>, 2014.
- Mougeot, X.: Improved calculations of electron capture transitions for decay data and radionuclide metrology, Appl. Radiat. Isot., 134, 225–232, <https://doi.org/10.1016/j.apradiso.2017.07.027>, 2018.
- Mougeot, X. and Helmer, R. G.: ^{40}K – Comments on evaluation of decay data, LNHB/INEEL, available at: http://www.nucleide.org/DDEP_WG/Nuclides/K-40_com.pdf (last access: 12 November 2020), 2009.
- Nähle, O., Kossert, K., and Klein, R.: Activity standardization of ^{22}Na , Appl. Radiat. Isot., 66, 865–871, <https://doi.org/10.1016/j.apradiso.2008.02.028>, 2008.
- Pradler, J., Singh, B., and Yavin, I.: On an unverified nuclear decay and its role in the DAMA experiment, Phys. Lett. B, 720, 399–404, <https://doi.org/10.1016/j.physletb.2013.02.033>, 2013.
- Preece, K., Mark, D. F., Barclay, J., Cohen, B. E., Chamberlain, K. J., Jowitt, C., Vye-Brown, C., Brown, R. J., and Hamilton, S.: Bridging the gap: $^{40}\text{Ar}/^{39}\text{Ar}$ dating of volcanic eruptions from the “Age of Discovery”, Geology, 46, 1035–1038, <https://doi.org/10.1130/G45415.1>, 2018.
- Renne, P. R.: $^{40}\text{Ar}/^{39}\text{Ar}$ age of plagioclase from Aca-pulco meteorite and the problem of systematic errors in cosmochronology, Earth Planet. Sc. Lett., 175, 13–26, [https://doi.org/10.1016/S0012-821X\(99\)00287-3](https://doi.org/10.1016/S0012-821X(99)00287-3), 2000.
- Renne, P. R., Sharp, W. D., Deino, A. L., Orsi, G., and Civetta, L.: $^{40}\text{Ar}/^{39}\text{Ar}$ dating into the historical realm: Calibration against Pliny the Younger, Science, 277, 1279–1280, <https://doi.org/10.1126/science.277.5330.1279>, 1997.
- Renne, P. R., Swisher, C. C., Deino, A. L., Karner, D. B., Owens, T. L., and DePaolo, D. J.: Intercalibration of standards, absolute ages and uncertainties in $^{40}\text{Ar}/^{39}\text{Ar}$ dating, Chem. Geol., 145, 117–152, [https://doi.org/10.1016/S0009-2541\(97\)00159-9](https://doi.org/10.1016/S0009-2541(97)00159-9), 1998.
- Renne, P. R., Mundil, R., Balco, G., Min, K., and Ludwig, K. R.: Joint determination of ^{40}K decay constants and $^{40}\text{Ar}^*/^{40}\text{K}$ for the Fish Canyon sanidine standard, and improved accuracy for $^{40}\text{Ar}/^{39}\text{Ar}$ geochronology, Geochim. Cosmochim. Ac., 74, 5349–5367, <https://doi.org/10.1016/j.gca.2010.06.017>, 2010.
- Renne, P. R., Balco, G., Ludwig, K. R., Mundil, R., and Min, K.: Response to the comment by W.H. Schwarz et al. on “Joint determination of ^{40}K decay constants and $^{40}\text{Ar}^*/^{40}\text{K}$ for the Fish Canyon sanidine standard, and improved accuracy for $^{40}\text{Ar}/^{39}\text{Ar}$ geochronology” by P.R. Renne et al. (2010), Geochim. Cosmochim. Ac., 75, 5097–5100, <https://doi.org/10.1016/j.gca.2011.06.021>, 2011.
- Rivera, T. A., Storey, M., Zeeden, C., Hilgen, F. J., and Kuiper, K.: A refined astronomically calibrated $^{40}\text{Ar}/^{39}\text{Ar}$ age for Fish Canyon sanidine, Earth Planet. Sc. Lett., 311, 420–426, <https://doi.org/10.1016/j.epsl.2011.09.017>, 2011.
- Schmidt-Ott, W.-D., Lauerwald, J., Bosch, U., Dornhöfer, H., Schrewe, U. J., and Behrens, H.: Electron-capture to positron ratio in the decays of ^{22}Na and ^{65}Zn : Proceedings of the 7th International Conference on Atomic Masses and Fundamental Constants AMCO-7, Technische Hochschule Darmstadt, Lehrdruckerei, Darmstadt, 3–7 September 1984.
- Steiger, R. and Jäger, E.: Subcommittee on geochronology: convention on the use of decay constants in geo-

- and cosmochronology, *Earth Planet. Sc. Lett.*, 36, 359–362, [https://doi.org/10.1016/0012-821X\(77\)90060-7](https://doi.org/10.1016/0012-821X(77)90060-7), 1977.
- Stukel, M.: Characterization of Large Area Avalanche Photodiodes For The Measurement of The Electron Capture Decay Of ^{40}K To The Ground State Of ^{40}Ar , MSc Thesis, Queen's University, Australia, 159 pp., 2018.
- Sýkora, I. and Povinec, P.: Measurement of electron capture to positron emission ratios in light and medium nuclides, *Nucl. Instrum. Methods Phys. Res. B*, 17, 467–471, [https://doi.org/10.1016/0168-583X\(86\)90189-8](https://doi.org/10.1016/0168-583X(86)90189-8), 1986.
- Vatai, E., Varga, D., and Uchrin, J.: Measurement of the ϵ/β^+ ratio in the decay of ^{22}Na and ^{74}As , *Nucl. Phys. A*, 116, 637–642, [https://doi.org/10.1016/0375-9474\(68\)90396-5](https://doi.org/10.1016/0375-9474(68)90396-5), 1968.
- Wang, M., Audi, G., Kondev, F. G., Huang, W. J., Naimi, S., and Xu, X.: The AME2016 atomic mass evaluation (II). Tables, graphs and references, *Chinese Phys. C*, 41, 030003, <https://doi.org/10.1088/1674-1137/41/3/030003>, 2017.
- Wasserburg, G. J. and Hayden, R. J.: $\text{A}^{40}\text{-K}^{40}$ dating, *Geochim. Cosmochim. Ac.*, 7, 51–60, [https://doi.org/10.1016/0016-7037\(55\)90045-4](https://doi.org/10.1016/0016-7037(55)90045-4), 1955.
- Williams, A.: Measurement of the ratio of electron capture to positron emission in the decay of Na-22 , *Nucl. Phys.*, 52, 324–332, [https://doi.org/10.1016/0029-5582\(64\)90696-0](https://doi.org/10.1016/0029-5582(64)90696-0), 1964.
- Yukawa, H. and Sakata, S.: On the Theory of the β -Disintegration 30 and the Allied Phenomenon, *Proc. Phys. Math. Soc. Jpn.*, 3rd Series, 17, 467–479, available at: https://www.jstage.jst.go.jp/article/ppmsj1919/17/0/17_0_467/_pdf (last access: 12 November 2020), 1935.