

## ***Interactive comment on “Expanding Limits of Laser-Ablation U-Pb Calcite Geochronology” by Andrew R. C. Kylander-Clark***

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One of the nicest aspects of a journal format such as this one is that the reviewers' comments are visible to everyone. Just as questions that follow conference talks can be as, or more illuminating than the talk itself, reviewer comments can be as insightful as a manuscript. Such is the case here, as the reviewer makes several important points that may otherwise be missed upon revision of the manuscript. I will comment on these points during resubmission, but it is important to note that many of the points will not be properly addressed within the manuscript itself. As such, I encourage others to read these comments (and replies) along with the manuscript. The reviewer comments are listed first, and my responses to each of those comments follow. \_\_\_\_\_ Com-

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ment: This is a paper detailing a methodology to do calcite U-Pb dating in which a very specialised instrument – the Nu Instruments P3D- is used which has multiple Daly ion counting detectors and an array of faraday cups. There are likely to be only a few of these instruments globally and so the main thrust of the paper is to show the specific advantages of this setup and to compare data with other, somewhat less sensitive instruments and detector arrays. In a sense this paper is about doing ‘traditional’ calcite dating using a very specialised detector array. Aside from this demonstration of superior sensitivity, there are no particular advances within the paper that improve the way we do calcite dating, but the performance of this instrument is demonstrated to be reliable and is impressive for its ability to analyse samples with minimal U and Pb. It thus may go some distance to opening up the analysis of very low-U calcite samples at higher spatial resolution to address problems that are otherwise challenging. Set up of the instrument. Right away I see a flaw in the set up: there is no ion counter to measure the  $^{232}\text{Th}$  signal, only a faraday cup and with noise level of 8000cps, measuring  $^{232}\text{Th}$  is thus a write-off with no useful data likely to be collected in a majority of samples. Thus, it is unlikely to be quantified in order to measure radiogenic  $^{208}\text{Pb}$  as a contribution to the total  $^{208}\text{Pb}$  signal, which is the largest isotope comprising common Pb. I will come back to this.

Reponse: I agree on this point. However, this is a disadvantage to most multi-collector instruments –  $^{232}\text{Th}$  is measured on a faraday cup. A further disadvantage is the necessity for proper calibration between ion counters. This is especially difficult when calibrating for  $^{232}/^{208}$  measurements, because there is no matrix-matched reference material for such a measurement. The  $^{287}\text{Pb}/^{207}\text{Pb}$  and  $^{208}\text{Pb}/^{206}\text{Pb}$  can be calibrated against NIST and other glasses, so if Th is negligible (though assessing this with a faraday cup would be difficult as pointed out above), then the  $^{208}$ -based correction can be made. Admittedly, I did not try to make a  $^{208}$ -based correction with this data. However, I have struggled to do so with more success than a TW diagram (i.e., a  $^{207}\text{Pb}$ –based correction) for calcite data in the past. I suspect this is due to a combination of: making precise measurements of  $^{232}$ ; making an accurate correction for

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232/208, and possibly an accurate correction of the 208Pb/206Pb ratio (though NIST glass should provide an accurate reference for this latter correction). I have added an explanation of these points to the analytical setup to make this point.

Comment: At the end of this section the author makes a curious statement: Around line 110, The data from the unknowns are all a bit scattered for geological reasons, and were culled to yield single populations for ease of comparison. (Though beyond the scope of this manuscript, the Paleozoic samples are interpreted to have suffered partial Pb loss or new crystal growth in the Cretaceous–Tertiary, and the older Cretaceous sample likely (re)crystallized over an extended period. When I read statements like this that suggest unknowns are all a bit scattered for geological reasons, it makes me wonder if this is just speculation, with some sort of analytical explanation for the scatter, at least in part, at play, and with the author(s) failing to examine the samples in depth to try to find out the explanation. It is so easy to suggest this sort of thing to explain messy data; in fact there are papers that invoke a wide range of unproven processes for scatter (U-loss, U-gain, Pb loss, variable common Pb composition; recrystallisation, etc.) all of which are just ad hoc explanations for scatter. The best approach, however, is to concentrate on such samples and try to really understand them with more measurements, particularly in their geological, hydrological, and textural context. There is little of this in this paper, largely because it is about methods, but this is an important point that all calcite dating people would do well to take more seriously.

Reponse: Agreed. As we, as a community, are able to make more precise measurements on smaller aliquots of material, making sense of scatter becomes more of an issue. This is an issue with all geochronometers to some extent; even zircon, one of the most robust chronometers, can yield data that scatter without a reasonable explanation for a particular large-scale geologic process. Rather the scatter can instead be linked to a process on a smaller-scale, such as those mentioned above (U-loss, U-gain, etc.). Calcite is proving to be one of the more complicated geochronometers, and the



community would greatly benefit from greater in-depth study of the particularly complicated samples. Such an examination in this case would likely detract from the main point of the paper, nevertheless, I can speculate on these particular calcite veins, as we have dated many from this region. The late Cretaceous sample (the most complicated sample) is likely a combination of slightly different vein-filling events. Spots were placed based on pre-analysis that showed regions of high U concentrations, and these groups of spots were scattered across the sample. Many of the samples from this study region yield scatter that correlates to the area of the vein that was analyzed; some of the different regions of the same vein will yield equivalent lower intercepts and different upper intercepts, and others different lower intercepts with similar or distinguishable upper intercepts. This suggests that much of the scatter is caused by multiple calcite-(re)crystallization episodes caused by fluid influx of variable common-Pb compositions. In the case of the Paleozoic sample, there was a similar sample (omitted from this study) that yielded a scatter of data that clearly indicated two (re)crystallization events during the Paleozoic and Cretaceous (clearly distinguishable arrays of data). It is my suspicion that the minor scatter in the Paleozoic samples in this study were caused by minor Pb loss during Cretaceous fluid migration.

Comment: Results Results are on several reference materials and three different 'unknowns'. The WC-1 calcite is taken as the primary calcite reference material and all results are normalised using a secondary normalisation to the WC-1 calcite; this is standard practise and well documented by earlier papers. The three unknowns are not particularly young – 440 Ma, 120 Ma, and 80 Ma and the reference materials also treated as unknowns are Duff Brown (64 Ma) and Ash (3 Ma). Only WC-1 is very radiogenic; the others have a wide spread in U/Pb ratios. The approach used in all samples is to do regression of an array of spots on the assumption that all measured points are syngenetic and formed at the same time in each sample. There is a comparison of the success rate and various other parameters that arise from the measurements of samples on the two instruments with the three set-ups. To no surprise, when U is very low, the faraday cup for its measurement performs relatively badly by comparison, but

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[Discussion paper](#)



it still does work, to be fair. What is a little bit surprising is how the standard Q-ICP-MS performs so well in comparison to the ideal Daly set up of the Nu P3D, which is illustrated well in the plot of figure 1. What I notice about this figure is that, discounting the experiment with faraday 238U, there is almost complete overlap between the P3D setup with Daly 238U and the Q-ICP-MS, and to some extent a bit more scatter in the P3D data. The Daly is of course better at very high and very low count rates, due to its higher saturation count rate and its lower noise at low count rates, but the advantages of the P3D are not anywhere near as significant as I thought they might be.

Reponse: This is an excellent observation; though I'd like to add a few comments. As can be interpreted from figure 1, when count rates are the same (which requires a considerably larger spot if using the Q-ICPMS), the uncertainty in the Q-ICPMS data is similar to that of the P3D. What is striking about this is that, as mentioned in the manuscript, the actual total counts for each isotope on the Q-ICPMS should be lower than that of the P3D, because each isotope is measured in sequence, whereas the P3D can measure all isotopes simultaneously. I would note that the limit of the precision of the Q-ICPMS data is lower than that of the P3D – this is why the P3D data was placed below the Q-ICPMS data in the figure.

Reponse: An important point should be made that the advantage of simultaneous collection is greatly reduced when the precision is limited by one isotope. For example, in Figure 1B, the  $^{206}\text{Pb}/^{238}\text{U}$  precision is commonly limited by  $^{206}\text{Pb}$ , rather than  $^{238}\text{U}$ , so one can count longer on  $^{206}\text{Pb}$  on the quad, thus reducing the advantage of multi-collection. This has long been known, and I was remiss to not point this out; I added a line in the text to correct this. That said, in the case of this experiment,  $^{232}\text{Th}$  and  $^{208}\text{Pb}$  were omitted from the cycle to improve the precision (but measured on the P3D). As noted by this reviewer, these isotopes are of interest if one is to make a  $^{208}\text{Pb}$ -based correction.

Comment: The author composes synthetic sample calculations to illustrate the potential strengths and weaknesses of each instrument and setup and then in section 4.2

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makes suggestions about which instrument and setup is best suited for unknowns. All of this is interesting, but largely a bit academic. The reality is that people who want to date calcite rarely have the luxury of having an initial session to measure their samples, establish a comprehensive picture of a sample's U and Pb concentrations, radiogenic to common proportions, and then have all of these set-ups available to them to then collect optimized data using the instrument/setup of choice. It will only rarely work that way. More often than not, samples have a geological significance and the challenge is to date as many of these samples as possible, as best as one can given the instrumentation available, and not to have to repeat the work unless necessary to answer ambiguities.

Reponse: This is a fair point; I would argue that synthetic dataset gives potential users an idea of the differences between the different types of instrumentation, and it can serve as a guide to non-experts. It is duly noted, however, that one cannot expect to achieve such precision as shown in the theoretical calculations, which is why the comparison is made between that and the actual data. It should be noted that the actual data shows the same trends as the theoretical data; this is discussed in section 4.2.2.

Reponse: It is also true that researchers rarely have the luxury of preablating their samples to determine the U,Pb concentrations as a screening tool. This is not always the case, however, and there are certainly cases in which samples can be triaged prior to analysis. The case of these samples is a perfect example: the samples were provided to me as part of a larger project to understand the faulting history of the eastern seaboard. A large number of veins were collected by a professor and his students, and were prescreened as part of several undergraduate projects on a 213 laser coupled to a quadrupole. Though this instrumentation is insufficient to provide accurate dates, it was instrumental in reducing the costly analysis time with the 193 laser/MC-ICPMS. The samples were sent to me with mapped U/Pb values which were rather accurate; analysis near the regions mapped with high U/Pb gave useable data, and analyzing

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Discussion paper



any other random region almost always gave useless data. This analytical approach yielded far more useful data in a shorter period of time (and much less \$\$) than analyzing 80-100 random spots by high-precision LA-MC-ICPMS on all the collected samples. Not everyone has this luxury, but it is a technique that can be utilized by many.

Comment: The other reality of all in situ dating, whether this be calcite LA-ICP-MS or accessory mineral SIMS or LA-ICP-MS, is that the Poisson 2SE precision on any ratio is never achievable, and usually with in situ dating, there is a +/-2% barrier that one cannot reduce. Calcite standards also have their own issues with absolute age such as WC-1 with its ~2% uncertainty in absolute age, and so to some extent the theoretical plots and analysis of this paper are not as applicable in practice as one would like. The one aspect I cannot find well-described in this contribution is the manner in which the cross calibration of gains and linearity of the multiple Daly detectors have been determined. Has this been done detector by detector using experimental setups that establish ion counting gains and linearity independently? Or, is the use of the WC-1 with its range of count normalisation the method – in other words, if the standard and the other secondary standards give the right data by primary WC-1 normalisation, then all of these linearity and gain issues get accounted for by such a blanket normalisation ‘fudge factor’? I suspect it’s the latter and frankly it probably works ok, but the authors should be a lot more clear on this, since this is an obvious instrumental issue that is normally detailed in papers that collect multiple ion counter data.

Reponse: I would argue that both the mass bias and detector efficiencies can be corrected with a single standard correction. Both are assumed to be linear corrections, so a single standard measurement can correct for the combination of both factors. That said, the two cannot be determined individually using this method. That is, we cannot determine the mass bias of Pb isotopes or the detector efficiencies, but only the combination of both. Determination of a robust detector efficiency is a rather painstaking process, as peak hopping requires recalibration of zoom lens parameters to ensure equal throughput for any given magnet set point. Hence, I included in the submitted

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[Discussion paper](#)



manuscript a small blurb about the advantage of a SC instrument: “Because there is only one SEM on a SC-ICPMS instruments, there is no need to cross calibrate multiple detectors, yielding simpler data reduction and the possibility for making 204- or 208-based common-Pb corrections,” and I further added in section 2, albeit cryptically, “was used first used to correct the 207Pb/206Pb for mass bias, detector efficiency, instrumental drift etc.” Nevertheless, I have also added a small bit of text about the lack of Daly detector calibration prior to analyses.

Comment: Overall, this is a very careful piece of work with good analysis of the data. It offers some insight into the top level performance of the P3D instruments with its multiple Daly detectors, and this helps all of us to understand what the benchmarks are as we evolve the techniques. I think the paper should be published largely as is (aside from the comment on ion counters above), because it is very well written and I can find no real issues with what has been done. However, below I comment on what has NOT been done that should be done on a subsequent comparative paper, ie., a followup study.

Additional comments: Oftentimes in calcite dating, there is little justification in doing it for old samples, and almost all of the action is on younger calcites. This is because there will always be in situ dating uncertainties of +/-2% minimum, and so for a 440 Ma sample this means +/- 10 Ma at least, and many processes that are being studied cannot be resolved when uncertainties get so large. Therefore, the challenge is really to make the method work well for young samples. This study has largely skirted around younger samples but for a methodology to be highly relevant to most geologists, we need the resolution to be a few million years at most; hydrological processes are often fast and subject to disturbance and therefore, revealing tests of methods has to include doing work on very young samples with not so much low U but also very low radiogenic Pb contents. Because textures are complex and often in diagenetic settings, there is no reason to think that all secondary calcite formed at the same time, there is a strong need to be able to calculate single spot ages, like we do in accessory minerals. So far,

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[Discussion paper](#)





few studies have done this.

Response: One could argue that this is applicable to all geochronologic studies, and not just those related to calcite, as this minimum commonly applies to most geochronometers analysed by LA-ICPMS. Calcite, as pointed out in other studies in this volume, can have larger uncertainties than better behaved minerals, and thus this point is well taken. To refer to an earlier comment, in many cases, researchers do not have the luxury to predetermine the age of their sample, and this technique is the closest they can get to understanding the timing of events. That said, if 2% isn't close enough, they shouldn't use this technique, and will have to look elsewhere to solving their problem.

Reponse: Though the natural samples in this dataset are older than many, they span range from which one could reasonably extrapolate their expectations for younger samples. This is the idea of the modelling work presented; a comparison of natural and synthetic data to give the reader an idea of what to expect for samples of different age and U concentration.

Comment: So far the majority of calcite dating studies have ignored  $^{208}\text{Pb}$  and  $^{232}\text{Th}$  and instead used the T-W diagram for age calculations via regression. This has several problems: 1. Regressions assume all spots are the same age when this may not be the case; scatter is often glossed over

Response: Good point: see previous comment on the scatter of the samples presented herein. Note also that one commonly assumes (as the reviewer also pointed out) that the common Pb value is the same for all analyses.

Comment: 2.  $^{207}\text{Pb}$  that is radiogenic is the limiting factor on accuracy and precision. When a sample is young (say 20 Ma) and U is low (0.1 ppm), and when  $^{207}\text{Pb}$  background noise is say 10cps, then the total radiogenic  $^{207}\text{Pb}$  might only be a factor of 1 or 2 above this noise, and no matter what setup is used, that will limit precision. Unless these low counts of  $^{207}\text{Pb}$  are accurate (and there can be issues when close to

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Discussion paper



background), the ages may not be accurate either. The  $^{207}\text{Pb}$  common Pb correction is therefore critical when using the T-W plot, and it has real limitations.

Reponse: This is an important point to highlight. When  $^{207}\text{Pb}$  is near background, the background corrected value may not just be imprecise, but because of possible interferences, the value could be inaccurate.

Comment: 3. This plot fails to use measurements of  $^{232}\text{Th}$  or  $^{208}\text{Pb}$ , which can be useful. Another way to do this in perhaps a more robust way is to use the  $^{232}\text{Th}$  and  $^{208}\text{Pb}$  in the plot with Y-axis being the  $^{208}\text{Pb}_{\text{common}}/^{206}\text{Pb}_{\text{total}}$  and the X axis being  $^{238}\text{U}/^{206}\text{Pb}_{\text{total}}$ . Calcite rarely has high Th and so the subtraction of  $^{208}\text{Pb}$  radiogenic is often trivial, allowing the common Pb correction to be done independently of a radiogenic isotope like  $^{207}\text{Pb}$ .

Reponse: These are all excellent points. The downside, not mentioned here, is twofold: 1) measurement of the  $^{232}/^{208}$  ratio is required for samples that contain Th (that is, one needs to correct for the mass bias, etc., of Th/Pb without a matrix-matched standard), and 2) one needs to calibrate the detectors for  $^{208}\text{Pb}/^{206}\text{Pb}$ . This latter point could be done using NIST, as is done for the  $^{207}\text{Pb}/^{206}\text{Pb}$ .

Comment: When the author stated early on in the manuscript that data are a bit scattered, I think it is possible if not probable that there may be issues in the measurement of the small  $^{207}\text{Pb}$  radiogenic component that could be mitigated if one does not rely upon measurement of  $^{207}\text{Pb}$  and instead takes advantage of the  $^{208}\text{Pb}$  measurement (virtually all common) and  $^{232}\text{Th}$  (usually very low). I discuss this in Parrish et al. (2017) and showed that often this approach works better.

Reponse: When mentioning the possibility of performing a 208-based in the introduction, I failed to reference this paper. This has been rectified in the revised version.

Comment: Secondly, readers of this sort of paper should always remember that the field of ICPMS and particular LA-ICP-MS has seen many orders of magnitude improve-

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Discussion paper



ment in sensitivity by the use of multiple quads, collision cells, and the like. It is never a good idea to take a standard ICP-MS such as that used in the measurements of this paper, as truly representative of the sensitivity of Q-ICP-MS instruments which can achieve sensitivities nearly as good as MC-ICP-MS in like-for-like experiments. This is just a caution. I would also love to see the author undertake some testing of individual spot ages using various methods, spot sizes, and so forth to evaluate the power of the P3D instrument to really outcompete Q-ICP-MS in challenging texturally complex samples that require smaller spot sizes to resolve texturally complex calcite/dolomite growth components. This is where I think the P3D might really have some clear blue water ahead of the other instruments. The application of instrument comparisons on older samples using just the isochron technique is, I think, NOT where the most interesting comparisons of methods of calcite dating are likely to be done. I hope something along these lines might be next project for the author.

Response: Another fair point – natural samples rarely reveal themselves as ideal candidates to serve as examples for the geologic community. This suggestion is certainly one of our aims going forward, as we continue to strive to create the ideal instrument configurations for geologic endeavors.

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Interactive comment on Geochronology Discuss., <https://doi.org/10.5194/gchron-2020-17>, 2020.

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