Technical note: on LA–ICP-MS U–Pb dating of unetched and etched apatites

Fanis Abdullin et al.: LA–ICP-MS U–Pb dating of apatites

Fanis Abdullin1, Luigi A. Solari2, Jesús Solé3, Carlos Ortega-Obregón2

1CONACyT–Centro de Geociencias, Campus Juriquilla, UNAM, Querétaro, 76230, Mexico
2Centro de Geociencias, Campus Juriquilla, UNAM, Querétaro, 76230, Mexico
3LANGEM, Instituto de Geología, UNAM, Ciudad Universitaria, CDMX, 04510, Mexico

Correspondence: Fanis Abdullin (fanis@geociencias.unam.mx)

Abstract

The same unetched and chemically etched apatite grains from five rock samples were dated U–Pb via laser ablation inductively-coupled plasma mass spectrometry (LA–ICP-MS). The objective of this study is to assert whether chemical etching required for apatite fission track analysis impacts the precision and accuracy of same-grain U–Pb ages. The results of our experiment suggest that etching has no significant effect on the accuracy of apatite U–Pb ages obtained by LA–ICP-MS. Thus, LA–ICP-MS can be used safely for apatite fission track and U–Pb double dating.

Short summary

Unetched and etched apatite grains from five samples were dated with U–Pb using laser ablation inductively-coupled plasma mass spectrometry. Our experiment indicates that etching needed for apatite fission track dating has no effect on the obtaining accurate U–Pb ages; therefore, the laser ablation-based technique may be used for apatite fission track and U–Pb double dating.
1 Introduction

Apatite, $\text{Ca}_5(\text{PO}_4)_3[\text{F,Cl,OH}]$, is the most common phosphate mineral in the Earth’s crust and can be found in practically all igneous and metamorphic rocks, in many ancient and recent sediments as well as in certain mineral deposits (Piccoli and Candela, 2002; Morton and Yaxley, 2007; Webster and Piccoli, 2015). This accessory mineral is often used as a natural thermochronometer for fission track, He, U–Th and U–Pb dating (e.g., Zeitler et al., 1987; Wolf et al., 1996; Ehlers and Farley, 2003; Hasebe et al., 2004; Donelick et al., 2005; Chew and Donelick, 2012; Chew et al., 2014; Cochrane et al., 2014; Liu et al., 2014; Spikings et al., 2015; Glorie et al., 2017).

Presently, apatite fission track (AFT) ages can be obtained rapidly by using laser ablation inductively-coupled plasma mass spectrometry (LA–ICP-MS) for direct measurement of “parent nuclides”, i.e., $^{238}\text{U}$ contents (Cox et al., 2000; Svojtka and Košler, 2002; Hasebe et al., 2004; Donelick et al., 2005; Vermeesch, 2017). LA–ICP-MS technique may be used to obtain $^{238}\text{U}$ for AFT dating, together with isotope ratios needed for U–Pb dating (e.g., Chew and Donelick, 2012; Liu et al., 2014; Glorie et al., 2017; Bonilla et al., 2020; Nieto-Samaniego et al., 2020).

Hasebe et al. (2009) previously performed an important experimental study, during which they demonstrated that chemical etching required for fission track dating has no significant effect on the accuracy of $\text{U}$ measurement by LA–ICP-MS method. After chemical etching of apatites, a smaller volume of ablated material is analyzed by LA–ICP-MS. The influence of etching needed for AFT dating on the precision and accuracy of same-grain U–Pb dating analyzed via LA–ICP-MS remains to be quantified. To investigate this issue, the same unetched and etched apatite grains extracted from five rock samples were analyzed using LA–ICP-MS for U–Pb dating. The
chosen rock samples have either emplacement or metamorphic ages ranging from the Cretaceous to the Neoproterozoic (see Table 1 for further details).

--- Table 1 ---

2 Sample descriptions

2.1 OV-0421 (Tres Sabanas Pluton, Guatemala)

This sample is a two mica-bearing deformed granite belonging to the Tres Sabanas Pluton, which is located of Guatemala City, Guatemala. For sample OV-0421, an emplacement age of 115 ± 4 (2σ) Ma was proposed based on zircon U–Pb data (Torres de León, 2016). A cooling age of 102 ± 1 (2σ) Ma, obtained with K–Ar (on biotite), has also been reported by the same author.

2.2 MCH-38 (Chiapas Massif Complex, Mexico)

MCH-38 is an orthogneiss from the Permian Chiapas Massif Complex. This rock was sampled to the west of Unión Agrarista, the State of Chiapas, southeastern Mexico. There is no reported age for this sample. Some zircon U–Pb dates obtained for the Chiapas Massif Complex (Weber et al., 2007, 2008; Ortega-Obregón et al., 2019) suggest that a Lopingian (260–252 Ma) crystallization or metamorphic age may be assumed for sample MCH-38.
2.3 TO-AM (Totoltepec Pluton, Mexico)

TO-AM is a granitic rock, sampled ca. 5 km west of Totoltepec de Guerrero, the State of Puebla, southern Mexico. There is no reported radiometric data for sample TO-AM. Previous geological studies indicate that the Pennsylvanian–Cisuralian Totoltepec Pluton was emplaced over a ca. 23 million year period (from ca. 308 to ca. 285 Ma; e.g., Kirsch et al., 2013).

2.4 CH-0403 (Altos Cuchumatanes, Guatemala)

CH-0403 was collected 5 km ESE of Barillas, in the Altos Cuchumatanes, Guatemala. It consists of a gray to green granodiorite. Five zircon aliquots of sample CH-0403 were dated using isotope dilution thermal-ionization mass spectrometry, yielding a lower intercept date of 391 ± 8 (2σ) Ma that is interpreted as its approximate crystallization age (Solari et al., 2009).

2.5 OC-1008 (Oaxacan Complex, Mexico)

This sample is a paragneiss from the Grenvillian Oaxacan Complex, southern Mexico. OC-1008 was collected in the federal road which connects Nochixtlán to Oaxaca. It was demonstrated that this sample underwent granulite facies metamorphism at 1000–980 Ma (Solari et al., 2014).
3 Analytical procedures

Accessory minerals were concentrated using conventional mineral separation techniques such as rock crushing, sieving, Wilfley table, Frantz magnetic separator, and bromoform. Approximately 300 apatite grains were extracted from each rock sample and mounted with their surfaces parallel to the crystallographic c-axis in a 2.5 cm diameter epoxy mount. Mounted crystals were polished to expose their internal surfaces (i.e., up to 4π geometry). For this experiment, complete crystals lacking visible inclusions and other defects, such as cracks, were carefully selected for analysis.

Sample preparation was performed at Taller de Molienda and Taller de Laminación, Centro de Geociencias (CGEO), Campus Juriquilla, Universidad Nacional Autónoma de Mexico (UNAM). Single spot analyses were performed with a Resonetics RESOlution™ LPX Pro (193 nm, ArF excimer) laser ablation system, coupled to a Thermo Scientific iCAP™ Qc quadrupole ICP-MS at Laboratorio de Estudios Isotópicos (LEI), CGEO, UNAM. During this experimental work, LA–ICP-MS-based sampling was performed in central parts of the selected apatite grains before and after chemical etching (in 5.5M HNO₃ at 21 °C for 20 s to reveal spontaneous fission tracks), as shown schematically in Fig. 1. The LA–ICP-MS protocol used for apatite analyses, as given in Table 2, was established on the basis of numerous experiments carried out at LEI during the past five years, and can be used for U–Pb and fission track double dating plus multielemental analysis (Abdullin et al., 2018; Ortega-Obregón et al., 2019). Corrected isotopic ratios and errors were calculated using Iolite 3.5 (Paton et al., 2011) and the VizualAge data reduction scheme (Petrus and Kamber, 2012). UcomPbine (Chew et al., 2014) was used to model ²⁰⁷Pb/²⁰⁶Pb initial values and thus force a ²⁰⁷Pb correction that considers the common Pb (non-radiogenic Pb) incorporated by apatite standards at the moment of their crystallization (see also Ortega-Obregón et al., 2019).
The “First Mine Discovery” apatite from Madagascar, with a mean U–Pb age of ca. 480 Ma (Thomson et al., 2012; Chew et al., 2014), was used as a primary reference material. The results for measured isotopes using NIST-612 (Pearce et al., 1997) were normalized using $^{43}\text{Ca}$ as an internal standard and taking an average CaO content of 55% (i.e., for F-apatites).

Tera and Wasserburg Concordia diagrams (T–W; Tera and Wasserburg, 1972) are used in apatite U–Pb dating, because the LA–ICP-MS-derived U–Pb results are generally discordant. The lower intercept in the T–W plot is considered as a mean apatite U–Pb age that should have geological significance (crystallization or cooling age, the age of mineralization or metamorphic event). Apatite U–Pb ages were calculated with IsoplotR (Vermeesch, 2017, 2018) and described below. Detailed information on our U–Pb experiments is given in Table S1 in the Supplement.

--- Figure 1 ---

--- Table 2 ---

4 Results

4.1 OV-0421

For rock sample OV-0421, 41 unetched apatites yielded a lower intercept age of 106 ± 4 (2σ) Ma with a mean square weighted deviation (MSWD) of 1.07, passing the chi-squared test with the $P(\chi^2)$ value of 0.35 (see in Fig. 2). Virtually the same U–Pb date, 107 ± 5 (2σ) Ma, was obtained after chemical etching of the same apatite grains, yielding a MSWD of 1.13 and a $P(\chi^2)$ of 0.27. Both these apatite U–Pb ages lie between the zircon U–Pb date of 115 ± 4 (2σ) Ma (i.e.,
crystallization age) and the biotite K–Ar age of 102 ± 1 (2σ) Ma (i.e., cooling age), which were previously obtained for the same granite sample by Torres de León (2016).

4.2. MCH-38

For orthogneiss sample MCH-38, the lower intercept in T–W yield a U–Pb age of 245 ± 6 (2σ) Ma (obtained from 41 unetched apatites) with a MSWD of 0.28 and a $P(\chi^2)$ of 1. Etched apatite grains from MCH-38 yielded an age of 240 ± 4 (2σ) Ma with a MSWD of 0.36 and a $P(\chi^2)$ of 1 (Fig. 2). Our U–Pb results are in close agreement with geochronological data reported from the Chiapas Massif Complex in previous studies (Damon et al., 1981; Torres et al., 1999; Schaaf et al., 2002; Ortega-Obregón et al., 2019). For instance, Torres et al. (1999) compiled biotite K–Ar ages, most of which lie within Early–Middle Triassic period. Triassic cooling ages in the Chiapas Massif Complex were also detected by Rb–Sr in mica–whole rock pairs that range from 244 ± 12 (2σ) Ma to 214 ± 11 (2σ) Ma (Schaaf et al., 2002).

4.3 TO-AM

Unetched apatites (32 crystals; Fig. 2) from granite TO-AM yielded a lower intercept date of 303 ± 5 (2σ) Ma with a MSWD of 0.6 and a $P(\chi^2)$ of 0.96. After etching, a slightly younger age of 299 ± 3 (2σ) Ma was obtained, with a MSWD of 0.89 and a $P(\chi^2)$ of 0.65. These apatite U–Pb ages are in line with the zircon U–Pb ages of 306 ± 2 (2σ) Ma to 287 ± 2 (2σ) Ma reported for the Pennsylvanian–Cisuralian Totoltepec Pluton (e.g., see details in Kirsch et al., 2013).
36 unetched apatite grains from sample CH-0403 yielded a lower intercept U–Pb age of 345 ± 10 (2σ) Ma with a MSWD of 0.7 and a $P(\chi^2)$ of 0.9, whereas etched grains yielded an age of 334 ± 8 (2σ) Ma with a MSWD of 1.37 and a $P(\chi^2)$ of 0.08 (Fig. 2). These cooling dates are considerably younger if compared to the CH-0403 emplacement age of 391 ± 8 (2σ) Ma (Solari et al., 2009).

41 unetched apatites belonging to the sample OC-1008 yielded a U–Pb age of 839 ± 12 (2σ) Ma with a MSWD of 0.98 and a $P(\chi^2)$ of 0.50. After etching, the same apatite crystals yielded an age of 830 ± 10 (2σ) Ma with a MSWD of 1.24 and a $P(\chi^2)$ of 0.14 (Fig. 2). Both these apatite U–Pb ages are significantly younger than the age of granulate facies metamorphism in the Grenville-aged Oaxacan Complex (1 Ga to 980 Ma, Solari et al., 2014), and thus, should be considered as cooling ages.

--- Figure 2 ---

5 Discussion and concluding remarks

Most rock samples, except OV-0421, yielded slightly younger apatite U–Pb ages after chemical etching (up to 3.3% in sample CH-0403). However, the lower intercept U–Pb ages obtained from unetched apatite grains are identical within errors to the U–Pb ages obtained on the same etched
grains (see diagram in Fig. 3). The results of our experimental study demonstrate that the chemical etching, required for the AFT analysis, has no important effect on the accuracy of apatite U–Pb ages determined via LA–ICP-MS. Thus, as a main conclusion of this experimental study, LA–ICP-MS can be used safely to obtain simultaneously AFT and U–Pb ages (i.e., double dating), as it was already done in some studies without previous proof (e.g., Chew and Donelick, 2012; Liu et al., 2014; Glorie et al., 2017; Bonilla et al., 2020; Nieto-Samaniego et al., 2020).

--- Figure 3 ---

**Supplement**

The supplement related to this article is available online at: https://...

**Author contributions**

Conceptualisation, investigation, and writing of the original draft were done by FA. LS and COO provided technical support. LS and JS acquired funding and resources, supervised the study, and reviewed the manuscript.

**Competing interests**

The authors declare that they have no conflict of interest.
Acknowledgements

The authors are grateful to Juan Tomás Vázquez Ramírez and Ofelia Pérez Arvizu for their help with sample preparation for this study. Professor Stuart Thomson is acknowledged for sharing Madagascar apatite. Dr. Michelangelo Martini kindly provided sample TO-AM that was useful for our experimental study. Dr. Ziva Shulaker, Dr. Jakub Sliwinski, and Professor Axel Schmitt are acknowledged for their constructive comments that improved our manuscript.

Financial support

This research has been supported by PAPIIT DGAPA UNAM (grant no. IN101520 to LS).

Figure caption

Figure 1

Illustration displaying the LA–ICP-MS-based U–Pb dating of the same apatite crystal before and after chemical etching (i.e., etched in 5.5M nitric acid at 21 °C for 20 s). Spot diameter of 60 µm.

Figure 2

Tera–Wasserburg Concordia diagrams for the U–Pb results of unetched and etched apatites from samples OV-0421, MCH-38, TO-AM, CH-0403, and OC-1008. MSWD – mean square weighted deviation, Ngr – number of grains dated. Errors are given in 2σ.

Figure 3

Plot showing the lower intercept U–Pb ages obtained on unetched and etched apatite grains.


Figure 1

LA-ICP-MS apatite U-Pb dating before etching

Chemical etching (5.5M nitric acid, 21 °C for 20 s)

LA-ICP-MS apatite U-Pb dating after etching
Figure 3

The graph illustrates the lower intercept U-Pb ages (Ma) obtained on unetched and etched apatite samples. The data points represent different samples, with annotations showing the sample names (e.g., sample OC-1008, sample CH-0403, sample TO-AM, sample CH-38, sample OV-0421). The graph shows a trend line indicating the relationship between the ages and the unetched versus etched conditions.
Table 1

Lithology, locality, and zircon U–Pb data for the selected experimental rock samples.

<table>
<thead>
<tr>
<th>Sample</th>
<th>Unit and locality</th>
<th>Rock type</th>
<th>Zircon U–Pb age</th>
<th>References</th>
</tr>
</thead>
<tbody>
<tr>
<td>OV-0421</td>
<td>Tres Sabanas Pluton, Guatemala</td>
<td>deformed granite</td>
<td>115 ± 4 Ma</td>
<td>Torres de León (2016)</td>
</tr>
<tr>
<td>MCH-38</td>
<td>Chiapas Massif Complex, Mexico</td>
<td>orthogneiss</td>
<td>ca. 260 to ca. 252 Ma (?)</td>
<td>Weber et al. (2007, 2008)</td>
</tr>
<tr>
<td>TO-AM</td>
<td>Totoltepec Pluton, Mexico</td>
<td>granite</td>
<td>ca. 308 to ca. 285 Ma (?)</td>
<td>Kirsch et al. (2013)</td>
</tr>
<tr>
<td>CH-0403</td>
<td>Altos Cuchumatanes, Guatemala</td>
<td>granodiorite</td>
<td>391 ± 8 Ma</td>
<td>Solari et al. (2009)</td>
</tr>
<tr>
<td>OC-1008</td>
<td>Oaxacan Complex, Mexico</td>
<td>paragneiss</td>
<td>990 ± 10 Ma</td>
<td>Solari et al. (2014)</td>
</tr>
</tbody>
</table>
Table 2

LA–ICP-MS protocol established at LEI to be applied for simultaneous apatite U–Pb and fission-track in-situ double dating plus multielemental analysis (REEs, Y, Sr, Mn, Mg, Th, U, and Cl).

**ICP-MS operating conditions**

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Value</th>
</tr>
</thead>
<tbody>
<tr>
<td>Instrument</td>
<td>Thermo Scientific™ iCAP™ Qc</td>
</tr>
<tr>
<td>Forward power</td>
<td>1450 W</td>
</tr>
<tr>
<td>Carrier gas flow rate</td>
<td>~1 L/min (Ar) and ~0.35 L/min (He)</td>
</tr>
<tr>
<td>Auxiliary gas flow rate</td>
<td>~1 L/min</td>
</tr>
<tr>
<td>Plasma gas flow rate</td>
<td>~14 L/min</td>
</tr>
<tr>
<td>Nitrogen</td>
<td>~3.5 mL/min</td>
</tr>
</tbody>
</table>

**Data acquisition parameters**

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Value</th>
</tr>
</thead>
<tbody>
<tr>
<td>Mode of operating</td>
<td>STD (standard mode)</td>
</tr>
<tr>
<td>Sampling scheme</td>
<td>2NIST-612–2MAD–1DUR–10apt–</td>
</tr>
<tr>
<td>Background scanning</td>
<td>15 s</td>
</tr>
<tr>
<td>Data acquisition time</td>
<td>35 s</td>
</tr>
<tr>
<td>Wash-out time</td>
<td>15 s</td>
</tr>
</tbody>
</table>

<table>
<thead>
<tr>
<th>Measured isotopes</th>
</tr>
</thead>
<tbody>
<tr>
<td>26Mg 31P 35Cl 41Ca 44Ca 55Mn 88Sr</td>
</tr>
<tr>
<td>89Y 139La 146Ce 144Pr 146Nd 147Sm</td>
</tr>
<tr>
<td>153Eu 157Gd 159Tb 165Dy 165Ho 166Er</td>
</tr>
<tr>
<td>160Tm 173Yb 173Lu 203Hg 204Pb 206Pb</td>
</tr>
<tr>
<td>207Pb 208Pb 232Th 238U [total = 29]</td>
</tr>
</tbody>
</table>

**Laser ablation system**

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Value</th>
</tr>
</thead>
<tbody>
<tr>
<td>Ablation cell</td>
<td>RESOlution™ Laurin Technic S-155</td>
</tr>
<tr>
<td>Model of laser</td>
<td>Resonetics RESOlution™ LPX Pro</td>
</tr>
<tr>
<td>Wavelength</td>
<td>193 nm (Excimer ArF)</td>
</tr>
<tr>
<td>Repetition rate</td>
<td>4 Hz</td>
</tr>
<tr>
<td>Energy density</td>
<td>*4 J/cm²</td>
</tr>
<tr>
<td>Mode of sampling</td>
<td>spot diameter of 60 µm</td>
</tr>
</tbody>
</table>

Note: MAD – “First mine Discovery” U–Pb apatite standard from Madagascar; DUR – Durango apatite from Cerro de Mercado mine (Mexico); apt – unknown apatite crystals. (*) Laser pulse energy of 4 J/cm², which was measured directly on target with a Coherent™ laser energy meter.