### 1 Chemical Abrasion: The Mechanics of Zircon Dissolution

- 2
- 3 Alyssa J. McKanna<sup>1,2</sup>, Isabel Koran<sup>2</sup>, Blair Schoene<sup>2</sup>, and Richard A. Ketcham<sup>3</sup>
- 4
- 5 <sup>1</sup>Los Alamos National Laboratory, EES-16, Los Alamos, NM 87545, USA
- 6 <sup>2</sup>Department of Geosciences, Guyot Hall, Princeton University, Princeton, NJ 08544, USA
- 7 <sup>3</sup>Jackson School of Geosciences, The University of Texas Austin, Austin, TX 78712, USA
- 8
- 9 Correspondence: Alyssa J. McKanna (ajmckanna@lanl.gov)

### 10 Abstract

11 Chemical abrasion is a technique that combines thermal annealing and partial

- 12 dissolution in hydrofluoric acid (HF) to selectively remove radiation-damaged portions
- 13 of zircon crystals prior to U-Pb isotopic analysis, and it is applied ubiquitously to zircon
- 14 prior to U-Pb isotope dilution thermal ionization mass spectrometry (ID-TIMS). The
- 15 mechanics of zircon dissolution in HF and the impact of different leaching conditions
- 16 on the zircon structure, however, are poorly resolved. We present a microstructural
- 17 investigation that integrates microscale X-ray computed tomography ( $\mu$ CT), scanning
- 18 electron microscopy, and Raman spectroscopy to evaluate zircon dissolution in HF. We
- show that μCT is an effective tool for imaging metamictization and complex dissolutionnetworks in three dimensions. Acid frequently reaches crystal interiors via fractures
- spatially associated with radiation damage zoning and inclusions to dissolve soluble
- high-U zones, some inclusions, and material around fractures leaving behind a more
- crystalline zircon residue. Other acid paths to crystal cores include the dissolution of
- 24 surface-reaching inclusions and the percolation of acid across zones with high defect
- 25 densities. In highly crystalline samples dissolution is crystallographically-controlled
- 26 with dissolution proceeding almost exclusively along the *c*-axis. Increasing the leaching
- 27 temperature from 180 °C to 210 °C results in deeper etching textures, wider acid paths,
- 28 more complex internal dissolution networks, and greater volume losses. How a grain
- 29 dissolves strongly depends on its initial radiation damage content and defect
- 30 distribution as well as the size and position of inclusions. As such, the effectiveness of
- 31 any chemical abrasion protocol for ID-TIMS U-Pb geochronology is likely sample-
- 32 dependent. We also briefly discuss the implications of our findings for deep-time (U-
- 33 Th)/He thermochronology.

34

### 35 1 Introduction

- 36
- 37 Zircon U-Pb dating by isotope dilution thermal ionization mass spectrometry (ID-TIMS)
- 38 produces high-precision dates that the Earth science community depends on to calibrate
- 39 geologic time (Bowring and Schmitz, 2003; Schoene, 2014). Zircon crystals affected by
- 40 radiation damage caused by alpha recoil events in the <sup>238</sup>U, <sup>235</sup>U, and <sup>232</sup>Th decay series
- 41 and the spontaneous fission of <sup>238</sup>U (Holland and Gottfried 1955; Weber et al., 1990;
- 42 Murakami et al., 1991; Meldrum et al. 1998; Trachenko et al., 2002; Ewing et al., 2003) –
- 43 can lose radiogenic Pb or more rarely U by diffusion, leaching, or recrystallization
- 44 compromising the accuracy of U-Pb ages (Mezger, 1997; Nasdala et al., 1998; Geisler et
- 45 al., 2002). Open system behavior can sometimes be identified graphically on a concordia
- 46 diagram when there is a mismatch between the  ${}^{238}U/{}^{206}Pb$  and the  ${}^{235}U/{}^{207}Pb$  isotopic
- 47 clocks, but sometimes discordia lines closely track concordia making Pb-loss difficult to
- 48 detect, thereby complicating age interpretations from zircon datasets (Mezger, 1997;
- 49 Schoene, 2014).

50 Chemical abrasion, a technique that combines thermal annealing to induce partial

- 51 structural recovery and leaching in hydrofluoric acid (HF) to selectively remove soluble,
- 52 radiation-damaged portions of crystals prior to U-Pb isotopic analysis, revolutionized
- 53 the field's ability to date zircon crystals affected by open system behavior (Mundil et al.,
- 54 2004; Mattinson, 2005; 2011). Still, many chemically abraded U-Pb zircon datasets
- 55 exhibit anomalously young, concordant dates that are often attributed to residual Pb-
- loss or, in rare instances older, reversely discordant dates (Mattinson et al., 1996,
- 57 Davydov et al., 2010; Schoene et al., 2010a; Schmitz and Davydov, 2012; Meyers et al.,
- 58 2012). Undetected open-system behavior can potentially bias or lead to the assignment
- 59 of inappropriate age uncertainties in critical geologic interpretations where ~100 ka
- 60 precision and accuracy matter such as correlations between terrestrial flood volcanism
- 61 and biotic crises or between biostratigraphic and radioisotopic calibrations constructed
- 62 to study key climate transitions in Earth history (Schoene et al., 2010a; Schmitz and
- Davydov, 2012). This ongoing challenge has recently prompted the ID-TIMS U-Pb
  community to more closely evaluate how different chemical abrasion protocols which
- contracting to more closely evaluate now different chemical abrasion protocols which
   can vary considerably both within and between individual laboratories affect
- 66 geochronological results (Huyskens et al., 2016; Widmann et al., 2019) and to explore
- 67 different frameworks for interpreting crystallization ages and uncertainties in complex
- 68 U-Pb datasets (e.g., Schoene, 2014).
- 69 Despite the near-universal acceptance of chemical abrasion, the mechanics of zircon
- 70 dissolution during acid digestion are poorly documented in the literature. Previous
- 71 work has demonstrated that acid dissolves U-rich rims and can reach the interior of
- some grains to preferentially dissolve U-rich zones in zircon cores (Mundil et al., 2004;

- 73 Mattinson, 2005; 2011). However, no study to date has systematically documented
- 74 zircon dissolution textures given a range of zircon types and leaching conditions, nor
- 75 leveraged such findings to gain a mechanistic understanding of the microstructural
- 76 processes that occur during partial dissolution in HF. Such an understanding would
- improve Pb-loss mitigation efforts and help ensure the accuracy of high-precision ID-
- 78 TIMS zircon U-Pb dates. In this study, we present the first three-dimensional (3D) view
- 79 of zircon dissolution based on microscale X-ray computed tomography data (μCT)
- acquired before and after leaching in HF. We evaluate zircon crystals from different
  geological settings with different degrees of radiation damage treated at different
- leaching conditions (180 °C vs. 210 °C, 4 h vs. 12 h). These data are paired with
- secondary electron images of etched grain surfaces and Raman spectral data used to
- 84 track changes in zircon crystallinity. In addition to achieving valuable new insights into
- 85 the mechanics of zircon dissolution, our  $\mu$ CT data reveal exciting opportunities for
- 86 quickly and non-destructively imaging radiation damage zoning in zircon in 3D which
- 87 has broader implications for zircon chronology.
- 88

## 89 2 Methods

- 90
- 91 **2.1** Samples
- 92

Table 1. Zircon samples.

Sample	Age & Rock Type	Radiation Damage	α-dose (α/g) Min Max	
AS3	Mesoproterozoic anorthosite	Intermediate-to-high	2×10 <sup>17</sup>	>1×10 <sup>19</sup>
SAM-47	Archean granitoid	Intermediate-to-high	6×10 <sup>17</sup>	2×10 <sup>18</sup>
KR18-04	Neoproterozoic rhyolite	Low-to-intermediate	5×10 <sup>16</sup>	7×10 <sup>17</sup>
BOM2A	Paleogene trachyte	Low-to-intermediate	6×10 <sup>15</sup>	2×10 <sup>17</sup>

93

94 Our study focuses on four zircon samples (AS3, SAM-47, KR18-04, and BOM2A) that

95 together span nearly the full radiation damage spectrum (Table 1). AS3 is an

96 intermediate-to-high damage sample from the Mesoproterozoic Duluth Complex

97 anorthositic series, emplaced during the North American Midcontinent Rift (Paces and

98 Miller, 1993; Schmitz et al., 2003; Takehara et al., 2018; Swanson-Hysell et al., 2020). The

sample of AS3 used in this study is the same as that studied by Takehara et al. (2018)

100 which was collected from the same locality as that of Paces and Miller (1993)

101 (92°09′32.4″, 46°45′43.4″). AS3 crystals are coarse-grained, orange to orangish-brown,

and fractured. Most grains are tabular prisms or anhedral shards and many show

103 evidence of hydrothermal alteration (Takehara et al., 2018). SAM-47 is an intermediate-



Figure 1. Summary of our experimental workflow.

- to-high damage Archean (3.32 3.29 Ga) sample from the Corunna Downs Granitic
- 105 Complex of the Emu Pools Supersuite in the eastern Pilbara Craton (-21°24′29.01″,
- 106 119°46′21.03″) (Barley and Pickard, 1999; Smithies et al., 2003; van Kranendonk et al.,
- 107 2007). Grains are euhedral, brown, and translucent. KR18-04 is an intermediate-to-low
- 108 damage sample from a Neoproterozoic rhyolite body associated with the
- 109 glaciolacustrine Konnarock Formation of Virginia, USA (MacLennan et al., 2020)
- 110 (36°41′47.95″, 81°24′22.08″). Grains are small, transparent, pink-orange and prismatic.
- 111 BOM2A is our lowest-damage sample from a Paleocene trachyte dike in Mumbai, India
- associated with rifting following the main phase of Deccan Traps volcanism (Basu et al.,
- 113 2020). Crystals are small, transparent, colorless, and prismatic.
- 114
- Aliquots of unannealed and annealed (900 °C for 48 h) grains from each of the four
- 116 zircon samples were set aside at the start of the study, mounted, polished, and
- 117 characterized using Raman spectroscopy to quantify the degree of radiation damage
- 118 present in each sample, as key bands in the zircon Raman spectrum broaden
- 119 predictably with increasing damage (Nasdala et al., 2001; Palenik et al., 2003; Váczi and
- 120 Nasdala, 2017). Annealed grain mounts were also imaged using optical microscopy,
- 121 cathodoluminescence (CL) imaging, and/or backscattered electron (BSE) imaging to
- 122 characterize growth textures for each sample (Fig. 1).
- 123

#### 124 **2.2** Workflow for partial dissolution experiments

125



Figure 2: Color, reflected light photomicrographs of zircon crystals mounted on tape for  $\mu$ CT imaging. (a) Photomicrograph of annealed grains prior to chemical abrasion (b) Photomicrograph of zircon residues following chemical abrasion. The after images illustrate how chemical abrasion anneals color centers in grains and renders grains colorless.

- 126 A diagram depicting our experimental workflow is presented in Fig. 1. Separate
- 127 aliquots of the four zircon samples were annealed in quartz crucibles in air at 900 °C for
- 12848 hours in a box furnace. Annealing conditions follow the recommendations of
- 129 Huysken et al., (2016) who demonstrated that hotter annealing temperatures likely
- 130 restore crystallinity to domains affected by Pb loss. Annealing durations within the ID-
- 131 TIMS U-Pb community typically range between 48 h and 60 h (Huysken et al. 2016;
- 132Widmann et al., 2019). Annealing studies of radiation damage in zircon demonstrate
- 133that annealing only weakly depends on heating duration after the first few hours of
- heating (Ginster et al., 2019, their Fig. 1). Thus, the difference between 48 and 60 h is not
- expected to significantly change zircon crystallinity or affect chemical abrasion
- 136 outcomes.
- 137
- 138 Annealed grains were mounted on sticky tape (~6 mm diameter circles fashioned using
- a hole punch) and imaged using optical microscopy. The four sticky tape mounts were
- then stacked on top of a pushpin and loosely secured with tape for μCT imaging
- 141 (Cooperdock et al., 2016). After imaging, grains were removed from the sticky tape and
- 142 transferred to individual Teflon microcapsules for leaching in concentrated HF in a Parr
- 143 Instrument Company pressure digestion vessel at 180 °C or 210 °C for a duration of 4 or
- 144 12 h. The chosen temperatures bracket the range commonly used for chemical abrasion
- 145 by the ID-TIMS U-Pb community (Huyskens et al., 2016; Widmann et al., 2019).
- 146 Leaching durations were selected based on the sample's initial radiation damage
- 147 content. Most intermediate-to-high damage zircon crystals (AS3 and SAM-47) were
- 148 chemically abraded at shorter durations to ensure that intact zircon residues remained

- 149 (as opposed to dust), although one subset of AS3 grains were leached at 180 °C for the
- 150 full 12 h. The intermediate-to-low damage samples (KR18-04 and BOM2A) maintained
- structural integrity over longer leaching durations, so grains were leached for the full 12
- 152 h period commonly used for chemical abrasion.
- 153

After partial dissolution, residues – the portions of zircon crystals that survive chemical
abrasion – were rinsed in Milli-Q water, dried down, and carefully transferred to fresh
sticky tape. Mounted residues were then re-imaged using optical microscopy and μCT
to generate a "before" and "after" imagery dataset. Microphotographs of annealed
grains and chemically abraded zircon residues are presented in Fig. 2. Following μCT,

- residue mounts were carbon coated, and secondary electron (SE) images of residue
- 160 surfaces were acquired using a scanning electron microscope (SEM). Raman spectra
- 161 were measured for a subset of zircon residues to characterize samples' crystallinities.
- 162

# 163 **2.3 Instrumentation and analyses**

164

165 Chemical abrasion was carried out using equipment and clean lab space at Princeton166 University. CL and BSE electron images of polished mounts were acquired using the

167 XL30 FEG SEM at the PRISM Imaging and Analysis Center at Princeton University

- 168 equipped with a mini-Gatan CL detector and a semiconductor BSE detector. Most
- images were acquired using a 10 kV accelerating voltage, 10 mm working distance, and
- 170 spot size 5. SE images of chemically abraded zircon residues were captured using the
- 171 Quanta FEG 200 Environmental-SEM also at the PRISM Imaging and Analysis Center.
- 172 This system is equipped with a Schottky field emission gun and Everhart-Thornley
- 173 secondary electron detector. SE images were acquired using low vacuum mode (~0.4 to
- 0.8 Torr) to minimize charging due to sample topography. Scans used a 10 kV
  accelerating voltage, 10 to 10.5 mm working distance, and spot size 4 or 5.
- 176

177 All X-ray computed tomography data were collected at the High-Resolution X-ray

177 All X-ray computed tomography data were conected at the High-Resolution X-ray

- 178 Computed Tomography Facility at the University of Texas at Austin using a Zeiss
  179 Xradia 620 Versa. Measurements were made with X-rays set to 120 kV and 15 W and
- 179 Aradia 620 versa. Measurements were made with X-rays set to 120 kV and 15 W and 180 prefiltered with the LE3 filter. For each scan 2401 views were obtained over a 360°
- prefiltered with the LE3 filter. For each scan 2401 views were obtained over a 360°
  rotation at 4 s per view on the 4x detector. 16-bit TIFF images were reconstructed at 1.62
- $\mu$ m/voxel, using a beam hardening correction setting of 1.8 in the Xradia Reconstructor
- 183 software. All 2D and 3D visualizations and quantitative measurements were made
- using Object Research Systems (ORS) Dragonfly software. Crystallographic dimensions
- 185 for BOM2A and KR18-04 were measured using the ruler function. Volume estimates for
- 186 these two samples were made using the software's "Upper OTSU" segmentation
- 187 function. This function differentiates zircon from inclusions, dissolution features, and
- 188 background (tape/air) based on grayscale intensity. Total volume is calculated by



Figure 3: Representative images of annealed AS3 and SAM-47 zircon that have not been treated by chemical abrasion. (a) SEM images of annealed AS3 zircon. I. A zircon with simple growth zoning. Arrows highlight dark hydrothermal alteration zones associated with fine-scale fractures. Some fractures cross-cut compositional zones. II. Zircon with an unfractured high-damage, CL-black rim and a fractured core. III. Zircon with row of fractures that cross-cuts a zone. IV. Zircon with a large melt inclusion oriented parallel to the c-axis. V. Zircon with convolute growth zoning. (b) Representative images of annealed SAM-47 zircon. I. Reflected light images showing fine-scale concentric growth zoning. II. BSE images showing that some grains are finely fractured. Some of these fractures pass though mineral inclusions (arrows).

- adding the number of selected high-intensity zircon voxels together. The volume of one
- 190 voxel is ~4.25  $\mu$ m<sup>3</sup>.
- 191 Raman spectra were acquired using the Horiba LabRAM Evolution Raman
- 192 spectrometer in the High-Pressure Mineral Physics Laboratory at Princeton University.
- 193 Measurements were made using either a 632.81 nm or 532 nm diode laser. The laser
- 194 power to the sample surface was ~8.5 to 17 mW and ~7.5 to 30 mW for the red and
- 195 green lasers, respectively. The instrument was calibrated daily using the silicon 520.7
- 196 cm<sup>-1</sup> Raman band and the automated protocol implemented within the Horiba Scientific
- 197 LabSpec6 software (Itoh and Shirono, 2020). Additionally, a quartz reference spectrum
- 198 was acquired daily to verify the accuracy of measured peak positions (Krishnam, 1945).
- All measurements were made using an 1800 g/mm grating, a 100  $\mu$ m slit, a 400 to 100
- $\mu$ m confocal pin hole, and either an Olympus 100x/0.9na lens or Mitutoyo 50x or 20x
- 201 long working distance objective lens.
- 202 This setup has a spectral resolution better than 2 cm<sup>-1</sup> and a spatial resolution of <1 to
- 203 ~5 μm. Polynomial background subtractions and Gaussian-Lorentzian peak fits were

- 204 made using LabSpec6 software. Peak widths have estimated uncertainties on the order
- of 10% (2 $\sigma$ ) based on tests of measurement and peak fit reproducibility. All reported
- 206 peak widths (full width at half maximum, FWHM) have been corrected for
- 207 instrumental broadening following the approach of Váczi (2014). A Raman spectrum for
- 208 a synthetic zircon grown using a Li-Mo flux method (Hanchar et al., 2001) was acquired
- as a loose analog for undamaged zircon.

210



Figure 4: Representative images of annealed KR18-04 and BOM2A zircon that have not been treated by chemical abrasion. (a) CL images of annealed KR18-04 zircon with fine concentric or broad, faint growth patterns. All scall bars are 50  $\mu$ m. (b) Representative BSE (top) and CL (bottom) images of annealed BOM2A zircon showing broad concentric growth zoning. Arrows highlight the frequent occurrence of apatite inclusions.

- 211 3 Results
- 212

# 213 **3.1 Images of polished grain mounts**

- 214
- 215 SEM and reflected light images of annealed AS3 and SAM-47 grains are presented in
- 216 Fig. 3. CL images of AS3 grains display broad concentric or convoluted zoning patterns
- 217 with evidence of hydrothermal alteration. Many crystals are finely fractured, and some
- 218 have large melt inclusions oriented elongate to the *c*-axis. Some fractures and alteration
- 219 zones cross-cut compositional zones. SAM-47 crystals are not CL luminescent. Reflected
- 220 light images acquired under the Raman microscope, however, reveal fine-scale

- 221 concentric zoning. BSE images indicate that some crystals are finely fractured and
- included. Many inclusions are cross-cut by fractures. SEM images of annealed KR18-04
- and BOM2A grains are presented in Fig. 4. Both samples exhibit concentric zoning with
- some faint, broad zones. Fractures are rare. Many BOM2A crystals have needle-like
- apatite inclusions.
- 226
- 227

Table 2. Minimum and maximum  $v_3$ (SiO<sub>4</sub>) FWHM values for unannealed, annealed, and chemically abraded zircon samples.

Sample	v <sub>3</sub> (SiO <sub>4</sub> ) FWHM (cm <sup>-1</sup> ) <sup>a</sup>		Sample	v <sub>3</sub> (SiO <sub>4</sub> ) FWHM (cm <sup>-1</sup> ) <sup>a</sup>	
Sample	Min	Max	Sample	Min	Max
AS3			KR18-04		
Unannealed	5.7	35.9	Unannealed	2.6	11.9
Annealed	3.6	20.0	Annealed	1.6	5.0
Chemically Abraded			Chemically Abraded		
180 °C, 4 h	5.0	11.0	180 °C, 12 h	1.9	3.4
210 °C, 4 h	5.2	14.0	210 °C, 12 h	2.2	4.5
180 °C, 12 h	4.0	11.0			
SAM-47			BOM2A		
Unannealed	10.4	24.1	Unannealed	1.8	5.4
Annealed	6.9	14.6	Annealed	2.3	5.2
Chemically Abraded			Chemically Abraded		
180 °C, 4 h	4.8	11.9	180 °C, 12 h	1.6	3.0
210 °C, 4 h	6.8	9.7	210 °C, 12 h	1.8	3.0

228

## 229 **3.2 Raman spectroscopy**

230

## **3.2.1** Polished grain mounts of unannealed and annealed samples

232

233 Key bands in the zircon Raman spectrum – most notably the  $v_3$ (SiO<sub>4</sub>) asymmetric SiO<sub>4</sub>

stretching band near ~1008 cm<sup>-1</sup> and the external  $E_g$  mode near ~357 cm<sup>-1</sup> – broaden and

shift to lower frequencies with increasing radiation damage (Nasdala et al. 1995, Zhang

et al. 2000, Nasdala et al. 2001, Anderson et al. 2020a, Härtel et al. 2021). Multiple

237 Raman analyses were made on several grains from each sample set to assess

intracrystalline variations in radiation damage. Measured  $v_3(SiO_4)$  and  $E_g$  peak widths

and positions are reported in Table S1. Peak width ranges for the  $v_3$ (SiO<sub>4</sub>) band for all

samples are summarized in Table 2. Equivalent alpha doses ( $\alpha$ /g) for unannealed

samples were derived using the relationship between the  $v_3(SiO)_4$  peak width and

equivalent alpha dose for Sri Lankan zircon (Palenik et al., 2003; Váczi and Nasdala,

243 2017) (Table 1). This relationship – calculated assuming an equivalent damage

accumulation interval of 375 Ma to account for the partial annealing of radiation

245 damage in Sri Lankan zircon – nicely fits the dataset of unannealed zircon presented by

Nasdala et al. (2001), suggesting that the relationship is broadly appropriate for zircon

247 from a wide range of geological environments.



Figure 5: Raman  $v_3$ (SiO<sub>4</sub>) and  $E_g$  peak width data for intermediate-to-high damage samples AS3 and SAM-47. (a) Results for unannealed and annealed (900 °C for 48 hour) zircon samples. Alpha dose estimates for unannealed zircon samples derived from measured  $v_3$ (SiO<sub>4</sub>) peak widths are shown on the right y-axis (Váczi and Nasdala, 2017). Slopes (m) for unannealed and annealed samples were calculated assuming a simple linear regression. Gray boxes mark the plot area presented in (b). (b) Results for chemically abraded residues compared to annealed samples. Reported slopes are inclusive of all leaching conditions. (c) Representative spectrum of synthetic zircon with peak assignments.

248 Unannealed AS3 and SAM-47 grains have intermediate-to-high degrees of radiation

- 249 damage with strong inter- and intra-crystalline variations (Fig. 5a, Table 2). CL black
- regions in AS3 samples that yielded anomalous zircon spectra with fluorescent artifacts
- 251 indicative of altered material were excluded from radiation damage estimates. Rims in
- 252 SAM-47 samples have accumulated more radiation damage than cores, indicating that
- rims are enriched in actinides relative to cores. Alpha dose estimates for both AS3 and
- SAM-47 span above and below the estimated alpha dose threshold assigned to fission
- track percolation  $1.9 \times 10^{18} \alpha/g$  (Ketcham et al., 2013). Importantly, this threshold also
- corresponds to key transitions in zircon material properties including density (Holland
- and Gottfried, 1955; Murakami et al., 1991; Ewing et al., 2003). Unannealed KR18-04 and
- 258 BOM2A zircon samples have low-to-intermediate levels of radiation damage and a
- 259 lesser degree of radiation damage zoning (Fig. 6a, Table 2).
- 260



Figure 6: Raman  $v_3$ (SiO<sub>4</sub>) and  $E_g$  peak width data for lower damage samples KR18-04 and BOM2A. (a) Results for unannealed and annealed (900 °C for 48 hour) zircon samples. Alpha dose estimates for unannealed zircon samples derived from  $v_3$ (SiO<sub>4</sub>) peak width measurements are shown on the right y-axis (Váczi and Nasdala, 2017). Slopes (m) for unannealed and annealed samples were calculated assuming a simple linear regression. Gray boxes mark the plot area presented in (b). (b) Results for chemically abraded residues compared to annealed samples. Reported slopes are inclusive of all leaching conditions.

- 261 Raman peak widths in annealed AS3, SAM-47, and KR18-04 samples are narrower than
- their unannealed counterparts indicating partial annealing of radiation damage (Fig. 5a,
- 263 Fig. 6a, and Table 2) (Zhang et al., 2000; Geisler et al., 2001a; 2001b; Ginster et al., 2019;
- Härtel et al., 2021). Peak width ranges for each sample are also more restricted implying
- that annealing has decreased the magnitude of inter- and intra-crystalline variations in
- 266 radiation damage. Annealing has had minimal effect on the crystallinity of BOM2A. The
- 267 most crystalline annealed BOM2A and KR18-04 samples have peak widths that closely
- 268 approach that of synthetic zircon. The slight differences between the natural and
- 269 synthetic samples could reflect minor residual radiation damage or slight differences in
- 270 lattice strain related to zircon composition and other intrinsic defects.
- 271
- 272 The relationship between the  $v_3$ (SiO<sub>4</sub>) and  $E_g$  peak widths steepens upon annealing in
- 273 each of the four samples, since the two Raman peaks have different temperature
- sensitivities (Härtel et al., 2021). This observation once again confirms that thermal



Figure 7: Raman  $v_3$ (SiO<sub>4</sub>) and  $E_8$  peak width results for all chemically abraded zircon residues.

annealing is not the inverse of radiation damage accumulation as demonstrated by
previous annealing studies (e.g., Zhang et al., 2000; Geisler et al., 2002; Ginster et al.,
2019). As such, we caution against using the Váczi and Nasdala (2017) v<sub>3</sub>(SiO<sub>4</sub>)-alpha

277 2017). As such, we callfor against using the vacal and russiand (2017) oscillet applied 278 dose relationship to derive alpha dose estimates for either the annealed or chemically

- 279 abraded samples.
- 280

## 281 3.2.2 Chemically abraded zircon residues

282

283 Raman results for chemically abraded residues broken down by zircon sample are 284 shown in Fig. 5b and Fig. 6b. The broadest peaks for AS3, SAM-47, KR18-04, and 285 BOM2A residues are narrower than their unleached counterparts indicating that HF 286 leaching has dissolved the most damaged material in each sample leaving behind a more crystalline zircon residue. Notably, residue datapoints for SAM-47 and BOM2A 287 288 samples largely plot below (at lower  $v_3$  for a given  $E_g$ ) the annealed datapoints. The slope of the  $v_3(SiO_4)$  and  $E_g$  relationship is also shallower for all four chemically abraded 289 290 sample sets when compared to their annealed sample sets. Taken together, these 291 observations could suggest that additional structural changes occur during HF leaching.

292

In Fig. 7 we compile Raman results for all chemically abraded residues to evaluate how

different leaching conditions affect zircon crystallinity. The spread in datapoints for AS3

residues leached at 180 °C for 12 h is shifted toward narrower values compared to AS3

and SAM-47 residues leached at either 180 °C or 210 °C for 4 h implying that increasing

- the leaching duration results in a more crystalline zircon residue due to the progressive
- 298 dissolution of higher damage domains. Somewhat surprisingly, leaching temperature
- 299 does not appear to have a significant effect on residue crystallinity; AS3 and SAM-47



Figure 8:  $\mu$ CT images of a zircon with density zoning. (a) A single 2D  $\mu$ CT image slice of an annealed – but not leached – AS3 grain with a dark, low-density damaged rim and a light, high-density crystalline core. (b) Semi-transparent 3D rendering of the  $\mu$ CT image stack for the same grain. High-density zircon is teal, and lower-density material is orange-brown. The arrow marks an interior inclusion. The faint stripes are surface indents of surficial inclusions not shown.

- samples leached for 4 h at 180 °C or 210 °C have residues with broadly similar peak
- 301 width distributions, as do KR18-04 and BOM2A samples leached for 12 h at 180 °C or
- 302 210 °C. This could reflect a small n-problem. AS3 residues leached at 180 °C for 12 h
- 303 have universally broader peak widths compared to KR18-04 and BOM2A residues
- 304 treated under the same leaching conditions highlighting that a sample's initial radiation
- 305 damage profoundly affects its residue's crystallinity.
- 306
- 307 3.3  $\mu$ CT imaging of radiation damage zoning
- 308
- 309 The accumulation of radiation damage decreases the density of zircon by 17% from ~4.7 to 3.9 g/cm<sup>3</sup> with the most rapid density change occurring over an alpha dose interval of 310 311 ~1×10<sup>18</sup> to ~4×10<sup>18</sup>  $\alpha/g$  (Holland and Gottfried, 1955; Murakami et al., 1991; Ewing et al., 312 2003; Nasdala et al., 2004). Raman data for unannealed AS3 and SAM-47 grains indicates that these samples have alpha doses spanning above and below this interval. 313 Lower density materials attenuate X-rays less, so metamict zircon should appear darker 314 315 in grayscale µCT image slices than crystalline zircon. Indeed, some AS3 and SAM-47 316 grains exhibit density zoning (Fig. 8), indicating that annealing at 900 °C for 48 h does not significantly increase the density of metamict material. Importantly, µCT does not 317 318 capture variations in radiation damage below the ~1×10<sup>18</sup>  $\alpha$ /g density-change threshold; density zoning is not observed in any of the low-to-intermediate damaged KR18-04 and 319 BOM2A samples. 320
- 321
- 322 3.4 Imaging textures before and after partial dissolution
- 323
- 324 **3.4.1 AS3**





Before(•) and after(•) partial dissolution

Figure 9: SE and  $\mu$ CT images of AS3 grains pre- and post-chemical abrasion (yellow dots and white dots, respectively) at 180 °C for 4 h. (a) I. Semi-transparent 3D renderings of  $\mu$ CT data for Zr17 showing melt inclusions removed by partial dissolution (yellow and white arrows) and newly visible fractures (black arrows). II. 2D  $\mu$ CT image slices showing the removal of a metamict rim and interior zone. III. 2D  $\mu$ CT cross section of the melt inclusion marked by white arrows in I. Newly visible radial fractures have developed along the length of the melt inclusion (black arrow). (b) I. SE images of Zr14 showing the widening of fractures on the grain surface. II. 2D  $\mu$ CT image slices showing a fracture network after partial dissolution. III. 3D rendering of  $\mu$ CT data showing radial fractures (black arrows) around large melt inclusions removed by partial dissolution. (c) SE images of zircon residues illustrating the contrast between a smooth, low damage surface and a higher damage pitted surface (Zr12), curved acid paths and small etch pits (Zr13), blocky fractures (Zr11 top), and dumbbell-like dissolution features (Zr11 bottom). (d) I. SE images of Zr16 showing the removal of fine-scale zones. II. 3D rendering of  $\mu$ CT data with showing the removal of large melt inclusions (yellow arrows), the formation of a parallel fracture sequence (black arrow), and significant volume loss likely due to breakage along the grain center where there are two giant melt inclusions.

- AS3 residues are white and brittle (Fig. 2b). Most residues treated at 180 °C for 12 h and
- a large fraction of grains treated at 180 °C or 210 °C for 4 h broke apart during rinsing or
- 328 transfer from the microcap to the tape. SE and  $\mu$ CT images of grains before and after
- 329 chemical abrasion are presented in Figures 9, 10, 11, S1, and S2. Each figure shows
- results for one of the three leaching conditions 180 °C for 4 h, 210 °C for 4 h, and 180



Figure 10: SE and  $\mu$ CT images of AS3 grains pre- and post-chemical abrasion at 210 °C for 4 h. (a) I. SE images of Zr04, a large crystal broken into four pieces. The rotated piece marked with a yellow arrow shows a nice cross-section of the grain interior. The arrow highlights an example of a branching channel. The higher magnification images show that these channels correlate with dumbbell features that cross-cut zones of relatively low (i) or high (ii) radiation damage. iii shows etch pit arrays likely indicative of dislocations loops or low-angle grain boundaries. II. 3D rendering of the  $\mu$ CT data shows the development of a complex dissolution network in the crystal's interior. III. 2D  $\mu$ CT image slice showing that the intensive fracturing observed in 3D is restricted to narrow plane within the crystal. (b) I. Semi-transparent 3D rendering of  $\mu$ CT data for Zr03 showing a large melt inclusion. II. SE images show elongated, channel-like dumbbells (low magnification) and the apparent removal of fine-scale zones (high magnification). III. 2D  $\mu$ CT image slice showing wide acid paths in the grain interior.

- 331 °C for 12 h. Here we briefly summarize key observations. We refer the reader to the
- 332 figure captions for additional context.
- 333
- 334 Damaged zircon is more soluble in HF than crystalline zircon. µCT images show that
- 335 low-density, high damage rims and interior zones dissolved early and at low
- temperatures (180 °C for 4 h). SE images also document the removal of fine zones early
- in the leaching process. Etching in SE images reflects the removal of soluble defects
- such as partially-annealed radiation damage, dislocations, low-angle grain boundaries,



Figure 2. SE and  $\mu$ CT images of AS3 pre- and post-chemical abrasion at 180 °C for 12 h. **(a)** I. SE images of sample Zr27 showing a row of dumbbells along the length of the zircon crystal. The higher magnification SE image shows a sponge-like surface texture. II. A series of 2D  $\mu$ CT image slices progressively stepping down to view structures beneath the crystal's surface. The yellow arrows highlight the same dumbbell features marked on the SE image in I. The teal arrows highlight fractures, many of which radiate from dumbbell features. The white arrows mark another series of dumbbells on the bottom side of the crystal. III. Cross-sectional 2D  $\mu$ CT image slices of a-a' and b-b' as labeled in II. White arrows mark second set of dumbbells with a different crystallographic orientation. IV. Semi-transparent 3D rendering of  $\mu$ CT data with arrows highlighting a large melt inclusion. The dissolution of this inclusion likely caused the grain to break into two pieces. The white arrows mark the same row of dumbbells as indicated by the white arrows in II. **(b)** I. Semi-transparent 3D rendering of  $\mu$ CT data for Zr28 prior. II. SE image of the husk-like zircon shell with large dumbbell features. **(c)** SE images of zircon residues Zr26, Zr32, and Zr23 showing cobble stone, straw, and lace-like textures.

- and intrinsic point defects. Low damage zones have smooth surfaces, whereas higher
- 340 damage zones have pitted or sponge-like surfaces due to etching of closely-spaced,
- radiation-related defects in SE images (Fig. 9c and Fig. 10aI). For spatial reference,
- 342 fission tracks are ~16.7 μm and alpha recoil tracks (clusters of alpha recoil tracks

stemming from a single decay chain) average ~125 nm in length prior to annealing or
etching (Ewing et al., 2003; Jonckheere, 2003). Etch pits are not observed in μCT images
due to the dataset's lower spatial resolution.

346

347 The shape of etch pits is independent of the nature of the defect (Jonckheere and Van den haute, 1996; Jonckheere et al., 2005; 2022). A pit's surface symmetry instead reflects 348 349 crystallographically-controlled dissolution, and etch pit geometries vary with crystallographic orientation (Gleadow et al., 1976; Yamada et al., 1995). As such, while 350 351 individual diamond-shaped etch pits resemble SE images of etched fission tracks 352 presented by others for zircon and other minerals (e.g., Jones et al., 2022), these likely reflect other defect types such as lattice dislocations. Fission tracks are expected to 353 354 anneal during the pre-leach 900 °C heating step (e.g., Yamada et al. 1995; 2007), 355 although some pits could reflect fission tracks that were pre-etched geologically. Given the limited abundance, spacing, and larger size of many diamond and pyramid-like 356 357 etch pits, we find them unlikely to represent alpha recoil tracks. Etch pit arrays that do 358 not correlate with expected zoning patterns (Fig. 10a-iii) are interpreted as dislocation 359 loops or low-angle grain boundaries.

360

Etch textures are subtle at low temperatures and short leaching durations. At hotter
temperatures and longer leaching durations, etched zones have deeper, sponge-like
textures indicative of a greater degree of dissolution. When leached at 180 °C for 12 h,
only a heavily dissected crystalline husk, a collection of perforated straw-like zones, or

- a cobble stone-like residue is sometimes all that remains.
- 366

Other interesting textures in AS3 residues include geometrical dissolution features that cross-cut radiation damage zones as highlighted in Fig. 10a and Fig. 11a which we refer to as dumbbells. Some dumbbells cross-cut zones of relatively high damage, while others cross-cut zones of relatively low damage. Dumbbells are oriented normal to the of the crystal (the *c*-axis). 3D rendering of  $\mu$ CT data reveal that dumbbells are surface expressions of complex, fracture networks that are spatially restricted to specific zones.

373 The geometrical shape of dumbbells and the wide, branching, and channel-like

- appearance of some fractures in SE imaging, indicate that these fracture networks are
- focal points for crystallographically-controlled dissolution.
- 376

Our μCT dataset also generates new insights into the fate of inclusions. In μCT image
slices of unleached grains, inclusions appear dark with grayscale intensities marginally

above that of background (air and tape) due to their low density and mean atomic

- number relative to that of zircon. We interpret an inclusion to have dissolved if its gray-
- 381 scale intensity decreases to that of background, if its size or morphology changes after
- 382 leaching, or if an acid path leads to the inclusion. We find that inclusions dissolved at

each leaching condition investigated. Radial fractures are commonly present arounddissolved inclusions in residues (Fig. 9b-I).

#### 385

#### 386 3.4.2 SAM-47

- 387
- Like AS3 residues, SAM-47 residues are white and brittle (Fig. 2b). Many residues broke
- during sample transfer, especially those leached at 210 °C. SE and  $\mu$ CT images of SAM-47 grains before and after chemical abrasion at 180 °C or 210 °C for 4 h are presented in
- 391 Figures 12, 13, S3, and S4. Some SAM-47 grains have density zoning with dark, high-



Before(o) and after(o) partial dissolution

Figure 12: SE and  $\mu$ CT images of SAM-47 grains pre- and post-chemical abrasion at 180 °C for 4 h. (a) I. SE images of Zr05 showing deep grooves on the grain's surface and a sponge-like etch texture. II. Opaque 3D rendering of µCT data showing that these surface fractures are only apparent after partial dissolution. III. Semi-transparent 3D rendering of µCT data with yellow arrows marking inclusions removed by partial dissolution. IV. 2D µCT image slices highlighting an example of an acid path into the grain interior (black arrow) and the removal of concentric zones (teal arrow). (b) I. 2D µCT image slices showing the removal of fine-scale concentric zones (teal arrow) and a mineral inclusion (yellow arrows) in Zr03. II. Semi-transparent 3D rendering of µCT data with yellow arrows depicting the removal of more mineral inclusions. (c) I. 2D  $\mu$ CT image slices of Zr09 showing the removal of a lowdensity rim (teal arrow) and an acid path into the grain interior (black arrow). II. Semi-transparent 3D rendering of µCT data highlighting the removal of inclusions (yellow arrows) and the formation of a large fracture (black arrow). (d) SE images of Zr53 showing crystal-shaped voids interpreted as dissolved surface-reaching inclusions (yellow arrow) and the fractures that crosscut these voids (black arrow). (e) SE images of Zr33 again showing fractures cross-cutting inclusions removed by partial dissolution (yellow arrows) and a smooth grain surface. II. 2D µCT image slices showing a convolute pattern of material dissolved from the crystal core. (f) I. 2D µCT image slices highlighting a lowdensity rim on Zr10. II. Semi-transparent 3D rendering of µCT data showing the removal of this rim.

damage rims and light, crystalline cores. One crystal exhibits concentric density zoning 392 393 in the grain interior. Like for AS3, these low-density zones dissolve at low leaching

- 394 temperatures and durations (180 °C, 4 h).
- 395

SE images of SAM-47 residues treated at 180 °C for 4 h show a range of surface textures 396 397 (Fig. 12). Some grains have smooth, unetched surfaces while others are more strongly 398 etched indicating inter- and intra-crystalline variations in radiation damage. Lowintensity chemical abrasion removes surface-reaching inclusions as evidenced by large 399 400 prismatic voids on grain surfaces. Most of these voids are crosscut by fractures. Other

- grains have finer sinuous fracture patterns not associated with inclusions. µCT images 401
- 402 show that acid has reached the interior of most zircon residues treated at 180 °C for 4 h and dissolved inclusions and fine-scale concentric and convolute zones from crystal
- 403

interiors.

404 405

SE images of SAM-47 residues treated at 210 °C for 4 h are more strongly etched with 406

- deep sponge-like textures (Fig. 13). Etch pits are larger with diamond-like shapes 407
- similar to those observed in AS3 crystals treated at either 210 °C for 4 h or 180 °C for 12 408
- 409 h, and fractures are wider. SE images indicate the dissolution of surface-reaching
- inclusions, and the shell-like appearance of some residues hints at the removal of 410
- 411 interior zones. µCT images of residues treated at 210 °C for 4 h reveal that concentric
- 412 zones and inclusions have been dissolved from crystal cores. Acid paths are wider and
- 413 more interconnected, and fractures crosscut dissolved mineral inclusions. We observe
- fracture patterns similar to the dumbbell features in AS3. Drawing a line normal to 414
- 415 dumbbell features in a µCT image slice of Zr30 forms a continuous concentric zone (Fig.



Before(•) and after(•) partial dissolution

Figure 13: SE and  $\mu$ CT images of SAM-47 grains pre- and post-chemical abrasion at 210 °C for 4 h. (a) I. SE images of Zr30 showing wide fractures, the removal of mineral inclusions (yellow arrows), and a moderately etched surface. II. 2D  $\mu$ CT image slices highlighting dumbbell-like features (yellow arrows) interpreted to cross-cut what could be a concentric zone (yellow dashed line). The black arrow exhibits how fractures radiate from the dumbbell features. III. Semi-transparent 3D rendering of  $\mu$ CT data. Yellow arrows correlate to those in I. The black arrow highlights how the fractures observed on the surface propagate through the crystal interior. (b) I. SE images of Zr36 showing fractures, diamond-shaped etch pits, and the targeted removal of an interior zone (yellow arrow). II. Semi-transparent 3D rendering of  $\mu$ CT data. The yellow arrow highlights the grain's shell-like appearance because of significant dissolution in the grain's interior. III. 2D  $\mu$ CT image slices showing the removal of mineral inclusions (yellow arrows), oscillatory zones (teal arrow), and dumbbell-like fractures that appear to cross-cut compositional zones (white arrows). (c) SE images of dog-chewed zircon residues Zr25, Zr27, and Zr25.

- 416 13a-II). Other fractures radiate from the dumbbell features. In sample Zr36 (Fig. 13b-III)
- 417 dumbbell features connect dissolved concentric zones both to one another and to the
- 418 grain surface in a scaffold-like pattern.
- 419





Figure 14: SE and  $\mu$ CT images of KR18-04 grains pre- and post-chemical abrasion at 180 °C for 12 h. (a) I. A low magnification SE image of zircon samples Zr38, Zr27, and Zr28 and higher magnification images of Zr27 showing close up images of rectangular and triangular etch pits and the removal of a surface-reaching inclusion (yellow arrow). II. Semi-transparent 3D rendering of  $\mu$ CT data for Zr27. Arrows highlight an inclusion inferred to have survived partial dissolution. (b) I. SE image of Zr40 with linear etch pit arrays likely indicative of dislocations. II. Semi-transparent 3D rendering of  $\mu$ CT data for Zr45. Teal arrows highlight a large inclusion inferred to have survived partial dissolution, while yellow arrows mark inclusions that dissolved. (c) I. Semi-transparent 3D rendering of  $\mu$ CT data for Zr45. Teal arrows highlight a large inclusion inferred to have survived partial dissolution, while yellow arrows mark inclusions that dissolved. Black arrows mark acid paths. II. 2D  $\mu$ CT image slices. Teal arrows mark the same multi-phase inclusion in I. Black arrows mark acid paths not apparent in the before imagery dataset. (d) Semi-transparent 3D rendering of  $\mu$ CT data for Zr36. Yellow arrows highlight surface-reaching inclusions removed by partial dissolution, resulting in a large cavity in the grain's interior.

421

#### 422 3.4.3 KR18-04

- 423
- 424 KR18-04 residues are transparent and colorless (Fig. 2b). Most residues remained intact
- 425 during rinsing and transfer. Only grains with large, pre-existing fractures broke apart.
- 426  $\mu$ CT and SE images of KR18-04 grains before after chemical abrasion at 180 °C or 210 °C
- 427 for 12 h are presented in Fig. 14 and Fig. 15, respectively. SE images of residues treated



Before(•) and after(•) partial dissolution

Figure 15: SE and  $\mu$ CT images of KR18-04 grains pre- and post-chemical abrasion at 210 °C for 12 h. (a) I. SE image of Zr13 showing dissolved inclusions (yellow arrows) and the removal of oscillatory zones (teal arrows). II. Opaque 3D rendering of  $\mu$ CT data. III. Semi-transparent 3D rendering of  $\mu$ CT data. IV. Representative 2D  $\mu$ CT image slices indicate that a significant amount of zircon material was dissolved from the grain's interior. Yellow arrows correlate to those in I. (b) I. SE image of Zr11 showing deep etch pits on (100) with the long axes oriented parallel to the crystal's *c*-axis. Etch pits are absent from other crystal faces. II. Semi-transparent 3D rendering of  $\mu$ CT data acquired before partial dissolution. III. Opaque and Semi-transparent 3D renderings and a representative 2D  $\mu$ CT image slice of the sample after partial dissolution. Black arrows highlight acid paths into the grain interior. (d) Semi-transparent 3D rendering of  $\mu$ CT data for Zr21. Yellow arrows mark an inclusion that dissolved. The black arrow highlights the acid path that inexplicably cut into the grain interior. (e) SE image of Zr10 with deep prismatic etch pits present on some grain surfaces but not others.

- 428 at 180 °C show intact grains with mildly etched surfaces (Fig. 14). Etch pits on (100) are
- 429 small, prismatic, and generally rectangular, while etch pits on other crystal faces are
- 430 more triangular, again highlighting that the shape of etch pits is crystallographically
- 431 controlled. Linear etch pit arrays are indicative of dislocation loops.

- 432
- 433 Large crystal-shaped voids on grain surfaces once again indicate that leaching dissolves
- 434 surface-reaching inclusions. μCT images of residues treated at 180 °C suggest that
- 435 leaching dissolves some but not all mineral inclusions from crystal interiors. For
- 436 example, the large multi-phase inclusion in Fig. 14c-I is interpreted to have survived
- 437 partial dissolution since 1) there is apparent change to the grayscale intensities of either
- 438 phase relative to that of background, 2) there is no apparent change to the inclusion's
- size or morphology, and 3) there is no evidence that an acid path has reached theinclusion. Beam hardening effects (the halo-like effect around high-density zircon)
- 441 make it challenging to identify whether or not smaller inclusions survive chemical
- abrasion. In such cases, grayscale intensity values cannot be used to identify whether or
- 443 not an inclusion dissolved. Some residues treated at 180°C have fractures or acid paths
  444 that lack obvious precursors in the before imagery dataset. Qualitatively, before-and-
- after  $\mu$ CT imagery suggest minimal volume loss and a slight shortening of prismatic
- 446 grain's c-to-a aspect ratio.
- 447

448 SE images of residues treated at 210 °C show the removal of fine concentric zones and 449 surface-reaching inclusions. Etch pits are well-preserved on some crystal faces including (100) and entirely absent on others. Etch pits are generally larger than those 450 observed in 180 °C residues. Many are deep, rectangular, and well-faceted. The long 451 452 axes of deep rectangular pits align parallel to the crystallographic *c*-axis, while the long 453 axes of shallower rectangular pits align parallel to the *a*-axis. Etch pit clusters have a 454 sponge-like texture. µCT images of residues treated at 210 °C show that acid has dissolved inclusions and zircon material from grain interiors. Some grains have deep 455 456 carveouts from crystal interiors with no obvious structural precursor in the before 457 imagery dataset. Before-and-after imagery suggest higher volume loss and a more 458 pronounced shortening of some grains' aspect ratios.

459

# 460 **3.4.4 BOM2A**

461

462 BOM2A residues are transparent and colorless (Fig. 2b). All residues remained intact during rinsing and transfer. SE and µCT images of BOM2A gains before and after 463 464 chemical abrasion at 180 °C or 210 °C for 12 h are presented in Fig. 16 and Fig. 17, 465 respectively. Etch pits are small and rectangular in SE images of residues treated at 180 466 °C (Fig. 16). Some etch pits are isolated while others are interconnected. Some surfaces have deep voids that penetrate the grain interior but do not correlate with inclusions. 467 468 µCT images qualitatively suggest minor volume loss with a slight shortening of the crystal's *c*-axis. Chemical abrasion dissolves surface-reaching inclusions and some – but 469 470 not all – inclusions from crystal interiors. Some residues have fractures that are spatially

471 associated with inclusions.

472



Figure 16. SE and  $\mu$ CT images of BOM2A grains pre- and post-chemical abrasion at 180 °C for 12 h. (a) I. Opaque 3D rendering of  $\mu$ CT data for Zr12. II. SE image of grain surface with close up image of clustered and isolated rectangular etch pits. The black arrow points to a void in the crystal perhaps related to a surficial inclusion not apparent in the pre-chemical abrasion dataset, and the yellow arrow highlights another interesting dissolution feature. III. Semi-transparent 3D rendering of  $\mu$ CT data showing inclusions removed by partial dissolution (yellow arrows) and inclusions inferred to have survived (teal arrows). IV. 2D  $\mu$ CT image slices with yellow arrows depicting inclusions dissolved during chemical abrasion and black arrows highlighting acid paths. (b) Semi-transparent 3D rendering of  $\mu$ CT data for Zr03 showing inclusions removed by partial dissolution (yellow arrows) and inclusions inferred to have survived (teal arrows). The black arrow highlights an acid path cutting through the crystal interior. (c) I. Semi-transparent 3D rendering of  $\mu$ CT data for Zr15 suggesting a slight shortening along the *c*-axis. II. 2D  $\mu$ CT image slices showing inclusions inferred to have survived partial dissolution. (d) Opaque 3D rendering of  $\mu$ CT data for Zr10 showing the removal of large, protruding apatite inclusions by partial dissolution.

- 473 SE images of residues treated at 210 °C show that etch pits are preserved on some
- 474 crystal faces but not others suggesting a crystallographic control on either etch pit
- formation or preservation (Fig. 17). Like KR18-04 residues leached under the same
- 476 conditions, etch pits are larger with well-developed facets at hotter leaching conditions.
- 477 Some etch pits are isolated while others interconnect to form acid paths into grain
- 478 interiors. The long axes of deep, prismatic etch pits on (100) align with the crystal's *c*-
- 479 axis, while the long axes of shallower etch pits align with the crystal's *a*-axis. Some SE



Before(•) and after(•) partial dissolution

Figure 17: SE and  $\mu$ CT images of BOM2A grains pre- and post-chemical abrasion at 210 °C for 12 h. (a) I. SE images of Zr31 showing deep fractures penetrating the grain's interior. Close up images show well-faceted etch pits on (100) some of which are isolated whiles others are interconnected. The long axes of deep, octahedral etch pits are oriented parallel to the *c*-axis, whereas the long axes of shallower etch pits are oriented parallel to the *a*-axis. II. Semi-transparent 3D rendering of µCT data again highlighting the development of large fractures. III. 2D µCT image slices. Teal arrows highlight inclusions that were dissolved, the yellow arrow points to a surface-reaching inclusion that acted as an acid path into the grain interior, and the black arrow highlights acid paths not observed in the before imagery dataset. (b) SE images of Zr40 that demonstrates how some crystallographic faces are strongly etched while others are pristine. Etch pits are again strongly prismatic and sometimes interconnected. The yellow arrow points to a void where there once was an inclusion. (c) I. Semi-transparent 3D rendering of  $\mu$ CT data for Zr34 showing a significant shortening of the crystal's *c*-axis. II. 2D µCT image slices. The yellow arrows highlight surface-reaching inclusions removed by partial dissolution. Black arrows mark acid paths not apparent in the before imagery dataset. (d) Semi-transparent 3D rendering of µCT data for Zr36. Teal arrows highlight inclusions inferred to have survived partial dissolution. Yellow arrows highlight inclusions that were dissolved. (e) Semi-transparent 3D rendering of µCT data for Zr28 showing significant volume loss from the grain interior. (f) Opaque 3D rendering of µCT data for Zr18. Yellow arrows highlight how some topographic features are preserved during partial dissolution despite significant volume loss. Note how crystal facets are better developed after partial dissolution. (g) Low magnification SE images of Zr34 and Zr32 showcasing the crystallographic-dependence of surface etching and acid paths that cut deep into grain interiors.

480 images show that acid has penetrated deeply into grain interiors forming what look like

- 481 caverns. Many of these caverns lack obvious precursors in in the before imagery
- 482 dataset. μCT images show that the dissolution of surface-reaching inclusions allows
- acid into crystal cores. Fractures in SE images are sometimes associated with large
- 484 mineral inclusions. Like the 180 °C leach, we find that leaching at 210 °C dissolves some
- 485 but not all interior inclusions. Qualitatively, volume loss appears greater at 210 °C,
- and the *c*-axis is considerably shorter in most crystals after partial dissolution. Before-
- and-after images show that some surface topographic features are preserved during
- chemical abrasion. Some residues are more strongly faceted than they were prior to
- 489 chemical abrasion.
- 490

# 491 **3.5 Quantifying volume loss and changes to crystal morphology**

492

All quantitative measurements made using the ruler and segmentation functions in

- 494 Dragonfly ORS software for samples KR18-04 and BOM2A are presented in
- 495 supplementary Tables S2 and S3 and summarized in Fig. 18. Leaching at 180 °C for 12 h
- 496 causes a ~5 to 10 % decrease in the length of a crystal's *c*-axis (Fig. 18a). Increasing the
- 497 leaching temperature to 210 °C results in a greater degree of shortening on the order of
- 498 ~15 to 30 %. In contrast, the length of a crystal's *a*-axes shows little (maximum <4 %) to
- 499 no change after leaching at 180 °C or 210 °C (Fig. 18b). Consequently, the aspect ratio
  500 (c/a) of a crystal decreases during chemical abrasion (Fig. 18c). A 2 % change in a crystal
- 501 with an initial axis length of 80  $\mu$ m equates to a change of 1.6  $\mu$ m which is
- 502 approximately the spatial resolution of our  $\mu$ CT dataset (1.62  $\mu$ m). As such, we take ~2
- 503 % to be a minimum estimate for our measurement error.
- 504

Estimated volume losses are presented in Fig. 18d. Fine-scale dissolution features andsmall mineral inclusions are sometimes missed by grayscale segmentation method used

- 507 due to a combination of beam hardening effects which manifest as bright halos around
- 508 zircon edges and the relatively low spatial resolution of the  $\mu$ CT dataset. As such,
- 509 volume loss estimations are considered first-order approximations for minimum
- volume loss. We find that chemical abrasion at 180 °C for 12 h dissolves ~5 to 10 % of a
- grain by volume, whereas chemical abrasion at 210  $^\circ$ C for 12 h dissolves ~25 to 50 % of a
- 512 grain by volume. Although there is considerable overlap between the BOM2A and
- 513 KR18-04 datasets at both leaching conditions, KR18-04 values are skewed toward higher
- 514 volume losses because KR18-04 grains have more radiation damage.
- 515
- 516 Despite clear evidence for dominantly *c*-axis dissolution, there is only a weak
- 517 correlation between a grain's aspect ratio and volume loss; crystals with aspect ratio's



Figure 18: Data plots summarizing crystal morphology, volume, and surface area measurements for KR18-04 and BOM2A. (a) Boxplot showing how the length of a grain's *c*-axis changes during chemical abrasion. In all box plots, the central line represents the dataset's median, the box extends to the dataset's 25<sup>th</sup> and 75<sup>th</sup> percentiles, the whiskers extend to include the full data range excluding outliers, and circles markers are outliers that exceed the 99% confidence interval. (b) Boxplot showing how the length of a grain's *a*-axis changes during chemical abrasion. (c) Boxplot showing how a grain's aspect ratio (*c*/*a*) changes during chemical abrasion. (d) Boxplot showing estimated volume loss during chemical abrasion. (e) Scatter plot showing the relationship between a grain's initial aspect ratio and estimated volume loss. (f) Scatter plot showing the relationship between a grain's initial surface-to-volume ratio and estimated volume loss.

- 518 <2.5 dissolve more readily than crystals with aspect ratio >2.5 in the samples leached at
- 519 210 °C (Fig. 18e). There is no correlation between a grain's initial surface area-to-
- 520 volume ratio and volume loss (Fig 18f).

521	4 Discussion
522	
523	4.1 The Mechanics of Zircon Dissolution
524	
525	4.1.1 Higher damage grains (~2×10 <sup>17</sup> $\alpha$ /g to >1×10 <sup>19</sup> $\alpha$ /g)
526	
527	In addition to dissolving high-damage, low-density rims acid easily accesses crystal
528	cores to dissolve inclusions and interior zones at short leaching durations (4 h) leaving
529	behind an inclusion-free residue with a higher degree of crystallinity in higher damage
530	grains like AS3 and SAM-47 (Fig. 5, Fig. 9 – Fig. 13). The most common acid path into
531	crystal cores in higher damage samples are fractures that are spatially associated with
532	radiation damage zoning and inclusions (Fig. 10a, Fig. 11a-b, and Fig. 13b). While
533	fractures are common in CL and BSE images of annealed AS3 and SAM-47 grains,
534	fractures are rare in $\mu$ CT images of annealed grains. This discrepancy reflects the
535	difference in spatial resolution between the two imaging methods. Fractures are visible
536	in $\mu$ CT images of residues because dissolution has widened them sufficiently.
537	
538	Radial or concentric fracturing related to internal stresses caused by volume expansion
539	of radiation-damaged domains is a common feature in zircon (Chakoumakos et al.,
540	1987; Lee and Tromp, 1995). Fracturing has also been attributed to differential stresses
541	caused by volume reduction of damaged domains during annealing (Geisler et al.
542	2001a, Geisler et al. 2002). CL images of annealed AS3 zircon illustrate that fractures
543	related to radiation damage zoning are indeed common (Fig. 3a-I, a-II, and a-III). Some
544	of these fractures exhibit evidence of hydrothermal alteration indicating that they are
545	geological in nature (Fig. 3a-I). We consider dual radiation damage accumulation and
546	annealing fracturing mechanisms to best explain why some residue fractures crosscut
547	zones of relatively high damage while others crosscut zones of relatively low damage
548	(Fig. 10a). Radiation damage zoning fracturing mechanisms also explain why complex
549	fracture networks are spatially restricted to certain zones (Fig. 10a, Fig. 11a).
550	
551	Radial fractures are evident around dissolved melt inclusions in AS3 residues (Fig. 9a-
552	b), and fractures that crosscut mineral inclusions are a common in both annealed SAM-
553	47 samples and chemically abraded residues (Fig 3b, Fig. 12d-e). BSE images of
554	unannealed SAM-47 grains confirm that some fractures formed prior to the thermal
555	annealing, however, we consider it likely that some fractures developed during thermal
556	annealing at 900 °C, since zircon and inclusions have different coefficients of thermal
557	expansion (e.g., Subbarao et al., 1990; Hovis et al., 2015). Stress fractures around
558	inclusions have long been used to identify heat treatment in gemstones
559	(Crowningshield and Nassau, 1981; Nassau 1981).
560	

- 561 While fractures related to radiation damage zoning and inclusions are the major
- 562 highways providing acid access to crystal interiors, SE images of overlapping etch pits
- indicate that acid also percolates across regions with high defect densities including
- zones of higher radiation damage (Fig. 10a, Fig. 11) and regions with dislocation loops
- 565 (Fig. 10a-iii). Increasing the temperature or duration of acid leaching results in more
- 566 pronounced and interconnected etching textures on grain surfaces, wider acid paths,
- and the formation of more complex dissolution networks within crystal cores.
- 568

# 569 **4.1.2 Lower damage grains (~6**×10<sup>15</sup> to 7×10<sup>17</sup> $\alpha$ /g)

570

571 The mechanics of zircon dissolution are considerably different for lower damage

- samples KR18-04 and BOM2A. Fractures spatially associated with large mineral
- 573 inclusions still play an important role as acid conduits to grain interiors (Fig. 14d, Fig.
- 574 16a, Fig. 17a). Fractures may be geological in nature or form during thermal annealing
- 575 (Crowningshield and Nassau, 1981; Nassau, 1981; Subbarao et al., 1990; Hovis et al.,
- 576 2015). Fracturing related to radiation damage zoning, however, does not meaningfully
- 577 contribute to zircon dissolution in samples with lower radiation damage and more
- 578 muted intracrystalline variations.
- 579



Figure 19: The zircon crystal structure (Hazen and Finger, 1979; Finch and Hanchar, 2003) rendered using CrystalMaker® software. ZrO<sub>8</sub> polyhedra are in light gray and SiO<sub>4</sub> tetrahedra are in teal. **(a)** Projection on (100) looking down the *a*-axis. **(b)** Projection on (001) looking down the *c*-axis. The yellow circle highlights the corner-sharing bonds between the SiO<sub>4</sub> tetrahedra and the ZrO<sub>8</sub> polyhedra.

580

- 581 Other mechanisms by which acid reaches a grain interior's is via the dissolution of
- 582 surface reaching inclusions (Fig. 14d, 15a, 16d) and the percolation of acid across
- regions with higher defect densities and overlapping etch pits (Fig. 15, Fig. 16a, Fig. 17a-
- b). In some samples, chemical abrasion dissolves large volumes from crystal cores
- 585 without clear structural reasons (Fig. 15c-d, Fig. 17e). This could reflect the dissolution
- 586 of zones with more radiation damage, but the pattern of the dissolved material does not

- obviously conform with the zonation patterns expected for these samples. Combined,
  these various acid paths lead to the dissolution of some but not all interior inclusions
  and some zones with higher degrees of radiation damage.
- 590
- 591 Importantly, µCT measurements indicate that dissolution in highly crystalline material
- is crystallographically-controlled and strongly anisotropic. Most dissolution occurs
- along the *c*-axis. Etch pits preserved on (100) suggest that dissolution along the *a*-axis is
- 594 mostly limited to the dissolution of defects that intersect the grain surface. In the (100)
- and (010) projections of the zircon structure ZrO<sub>8</sub> polyhedra share edges with adjacent
- ZrO<sub>8</sub> polyhedra and SiO<sub>4</sub> tetrahedra (Fig. 20) (Hazen and Finger, 1979; Finch and
  Hanchar, 2003). Whereas in the (001) projection of the zircon structure, ZrO<sub>8</sub> polyhedra
- share edges with adjacent ZrO<sub>8</sub> polyhedra and *corners* with adjacent SiO<sub>4</sub> tetrahedra. We
- 599 infer that corner sharing bonds in the (001) plane are easier to break during dissolution
- 600 than the solely edge-sharing bonds in the (100) and (010) planes causing faster
- dissolution along the *c*-axis. Increasing the leaching temperature from 180 °C to 210 °C
- 602 leads a more significant shortening of a crystal's aspect ratio and greater volume loss. In
- 603 lower damage grains that lack fractures, surface-reaching inclusions, and
- 604 interconnected defect zones, grains predominantly dissolve from rim-to-core along the
- 605 crystal's *c*-axis (Fig. 14a, Fig. 15b, and Fig. 16c).
- 606

607 In the two samples analyzed, leaching temperature and a crystal's bulk radiation 608 damage has the strongest control over volume loss. Results for BOM2A show that 609 crystals with very high aspect ratios might dissolve more slowly than more equant 610 grains, since dissolution along the *c*-axis is the rate-limiting process. A grain's initial 611 surface-to-volume ratio does not affect volume loss.

612

# 613 4.2 Implications for ID-TIMS U-Pb geochronology

614

615 4.2.1 Zircon U-Pb ages and trace element analyses

616

617 The goal of this study is to construct a mechanistic understanding of zircon dissolution618 and identify possible implications for U-Pb dating and coupled trace element analyses

619 upon which future geochronological and geochemical investigations – such as the

620 single-crystal stepwise partial dissolution experiments that are currently underway by

- 621 authors AJM and BS can build.
- 622

As discussed above, how a zircon dissolves strongly depends on its initial radiation

- 624 damage content and the distribution of radiation damage and other defects within the
- 625 crystals and associated fractures. Dissolution also depends on the size and distribution
- of inclusions within a grain and the extent to which fractures develop around these

- 627 inclusions. To a lesser degree, crystal morphology also affects dissolution. As such, the
  628 effectiveness of any chemical abrasion protocol will inherently be sample dependent.
  629
- 630 Here we briefly consider the idealized case of a concentrically-zoned magmatic zircon.
- 631 This discussion focuses on the dissolution of an intact single crystal; some ID-TIMS U-
- 632 Pb studies analyze polished half-grains or targeted portions of a single crystal.
- 633 Magmatic crystallization of zircon occurs over a period of time within a magma
- 634 chamber. As such, zircon cores are intrinsically older than rims and often differ
- 635 compositionally. A rim-to-core model for zircon dissolution implicitly suggests that
- dissolving more zircon during chemical abrasion by either increasing the temperatureor duration of leaching will remove a greater portion of a crystal's rim and bias its U-Pb
- 637 or duration of leaching will remove a greater portion of a crystal's rim and bias its U-Pb638 date and trace element content toward an older value and core composition,
- 639 respectively. This is especially concerning for geochronological studies of volcanic rocks
- 640 where the youngest U-Pb date or population of dates is often taken to represent the age
- 641 of a volcanic eruption.
- 642

643 Our results suggest that the majority of zircon crystals evaluated in this study do not644 predominantly dissolve from rim-to-core. While increasing the intensity of chemical

- abrasion leads to greater volume loss, much of that added loss comes from the
- 646 dissolution of interior zones. Consequently, a typical U-Pb analysis of a zircon residue
- is more likely to reflect the absence of soluble high U zones irrespective of age variation
- 648 within single grains. As such, analyses of zircon residues are more likely to broadly
- 649 represent mixed core-rim ages and trace element compositions. The proportion of rim-
- to-core material will inherently be sample- and leaching condition-dependent. Only
- 651 grains with low radiation damage, few-to-no inclusions, and no pre-existing fractures
- are likely to conform to a rim-to-core dissolution model with dissolution predominantly
   progressing along a crystal's *c*-axis; however, since rim material on (100) is preserved
- 654 due to limited dissolution along *a*, there likely remains a mixed core-rim age component
- 655 to each analysis.
- 656
- 657 4.2.2 Inclusions and zircon trace element analyses
- 658

659 Integrating chemical abrasion ID-TIMS U-Pb dates with trace element analyses (TEA) of

- the same volume of dissolved zircon can provide important information about
- 661 petrogenetic processes (Schoene et al. 2010b). The integrative TEA approach, however,
- broadly assumes that inclusions are also dissolved during chemical abrasion, such that
- the final volume analyzed is zircon as opposed to a zircon-inclusion mixture. While
- 664 geochronologists generally endeavor to select inclusion-free grains, this is not possible –
- or desired for all zircon samples, and not all inclusions can be identified optically with
- 666 a standard binocular picking scope.

667

- 668 The fate of inclusions during chemical abrasion has never been rigorously investigated.
- 669 Our data suggest that inclusions are readily dissolved in grains with intermediate-to-
- 670 high radiation damage densities due to the development of stress fractures that form
- either geologically or during thermal annealing at 900 °C. These findings strongly
- 672 emphasize that the annealing step of chemical abrasion is important not just for
- 673 minimizing leaching-induced elemental and isotopic fractionation (Mattinson, 2005;
- 674 2011), but also for building acid paths into grain interiors to dissolve inclusions. In
- lower damage grains, our findings suggest that small inclusions armored by highly
- 676 crystalline zircon can survive 12 h of chemical abrasion at 210 °C. As such, some lower
- damage residues may be susceptible to inclusion contamination. Increasing the leaching tamparature from 180 °C to 210 ° improves the likelihood that in ducing will be
- 678 temperature from 180 °C to 210 ° improves the likelihood that inclusions will be
- 679 removed, but it does not guarantee it.
- 680

# 681 4.3 Imaging radiation damage zoning: Implications for (U-Th)/He thermochronology

The accumulation of radiation damage in zircon has a profound impact on not only on

- 683 U-Pb geochronology, but also on He diffusion kinetics and deep-time zircon (U-Th)/He
- thermochronology (Guenthner et al. 2013; Cherniak, 2019; Anderson et al. 2017; 2020b).
- 685 While cathodoluminescence imaging and Raman 2D spectral mapping have previously
- been used to either qualitatively or quantitatively characterize the distribution ofradiation damage in polished zircon grains prior to laser ablation zircon (U-Th)/He
- 688 analyses (Danišík et al., 2017; Anderson et al., 2017; 2020a), finding a method for rapid
- 689 and non-destructive 3D characterization of strong radiation damage zoning in
- 690 unpolished grains for single-crystal zircon (U-Th)/He dating has remained elusive.  $\mu$ CT
- 691 offers an exciting new way to quickly screen zircon grains for strong radiation damage
- 692 zoning prior to (U-Th)/He analysis. Strongly zoned grains could either be excluded
- from datasets or corrections could be applied to account for expected intracrystalline
- 694 variations in He diffusivity.  $\mu$ CT data can also be used to identify mineral phases or
- 695 inclusions and intergrowths that might impact He systematics (Cooperdock et al., 2016;
- 696 Cooperdock and Stockli, 2018; Cooperdock et al. 2022), and improve alpha ejection
- 697 corrections by providing zoning information and generating more robust surface area-
- 698 to-volume estimates (Cooperdock et al., 2019).

# 699 5 Conclusions

- 700 In this study we present a microstructural investigation of four zircon samples covering
- a range of ages and radiation damage densities evaluated before and after chemical
- abrasion in HF acid in pressure digestion vessels at 180  $^{\circ}$ C or 210  $^{\circ}$ C for 4 h or 12 h.
- 703 Results yield new insights into the mechanics of zircon dissolution and advance  $\mu$ CT as

an effective tool for the rapid – and non-destructive – imaging of strong radiationdamage zoning in zircon in 3D.

706 How a zircon dissolves strongly depends on the degree of radiation damage and the 707 nature of intracrystalline variations. Dissolution also depends on the size and placement of inclusions. In lower damage zircon (~6×10<sup>15</sup> to 7×10<sup>17</sup>  $\alpha$ /g), dissolution is strongly 708 709 anisotropic; dissolution dominantly progresses along the *c*-axis with minimal 710 dissolution occurring along *a*. Acid reaches the interior of many lower damage crystals 711 via the dissolution of surface-reaching inclusions, fractures that crosscut inclusions, and 712 by the percolation of acid across closely-spaced, soluble defects to remove interior zones 713 with higher degrees of damage and some – but not all – mineral inclusions. In addition 714 to inclusions and radiation damage, acid also preferentially attacks intrinsic defects 715 such as dislocation loops.

- 716 In higher damage samples  $(2 \times 10^{17} \alpha/g \text{ to } >1 \times 10^{19} \alpha/g)$ , acid readily dissolves low-
- 717 density, high-damage rims and regularly accesses crystal cores to dissolve inclusions
- and interior zones with higher degrees of damage resulting in a more crystalline
- residue. The most common acid path into the interior of the higher damage samples
- 720 analyzed are planar fractures associated with radiation damage zoning, fractures that
- form around inclusions, and acid percolation across regions with high-defect densities
- which forms sponge-like textures. Fractures reflect differential stress caused by volume
- expansion and/or reduction of radiation damaged domains and inclusions. Some
- fractures are geological in nature, but a subset of fractures like formed during the
- thermal annealing step conducted prior to leaching (900  $^{\circ}$ C for 48 h); these results
- highlight the important role that the annealing step of chemical abrasion plays in
- 727 generating pathways for acid to reach crystal interiors.
- 728 Increasing the leaching temperature or duration leads to the development of wider acid
- 729 paths, more extensive dissolution networks, and the development of deeper sponge-like
- 730 surface textures. In the lower damage samples analyzed, increasing the leaching
- temperature by 30 °C resulted in an increase in volume loss of up to  $\sim$ 40 %.
- 732 The effectiveness of any chemical abrasion protocol for ID-TIMS U-Pb geochronology
- 733 will ultimately be sample-dependent. Most residue dates like reflect a mixture of core
- and rim material, although the proportion of rim relative to core is expected to be both
- 735 sample- and leaching-condition dependent. Future microstructural investigations
- should focus on a wider range of zircon ages, morphologies, and geological
- rank environments of formation to help build a broader intuition for how different zircon
- 738 populations dissolve. Other future studies could integrate textural data with
- 739 geochemical and geochronological analyses of leachates and residues to further

- 740 elucidate the mechanics of dissolution. Studies that evaluate how different annealing
- conditions affect zircon micro-fracturing or the rate of dissolution would also be
- 742 beneficial.
- 743 **Supplement.** The supplement to this article is available online at:
- Author Contributions. AJM designed and conducted the experiments. All authors
   participated in the interpretation of the experimental results. AJM prepared the figures
- 746 and manuscript.
- 747 **Competing Interests.** The authors declare no competing interests.
- 748 Acknowledgements. We would like to thank Jessie Maisano of the University of Texas
- 749 High Resolution CT facility for helping us to acquire  $\mu$ CT data during the height of the
- 750 global pandemic, and Tom Duffy of Princeton University for use of his Raman system.
- 751 Thank you also to Mami Takehara of the National Institute of Polar Research in Tokyo,
- Japan for providing the hydrothermally altered AS3 zircon crystals used in this study.
- 753 We are indebted to Tyler McKanna for providing computing resources, and we thank
- 754 Dawid Szymanowksi for many hours of valuable discussion. We would also like to
- acknowledge the constructive commentary from the geochronology community,
- 756 reviewers, and editor Charles, Magee, Magdalena Huyskens, Fernando Corfu, and
- 757 Daniel Condon that helped strengthen this manuscript.
- **Financial support.** This work was supported by research funds provided by the
- 759 Department of Geosciences at Princeton University granted to Alyssa J. McKanna as
- 760 part of her Harry Hess Postdoctoral Fellowship.
- 761 **Review Statement.**

# 762 **References**

- Anderson, A. J., Hodges, K. V., and van Soest, M. C.: Empirical constraints on the effects of
  radiation damage on helium diffusion in zircon, Geochim. Cosmochim. Acta, 218, 308–322,
  https://doi.org/10.1016/j.gca.2017.09.006, 2017.
- 766
- Anderson, A. J., Hanchar, J. M., Hodges, K. V., and van Soest, M. C.: Mapping radiation
  damage zoning in zircon using Raman spectroscopy: Implications for zircon chronology, Chem.
  Geol., 538, 119494, https://doi.org/10.1016/j.chemgeo.2020.119494, 2020a.
- 770
- 771 Anderson, A. J., van Soest, M. C., Hodges, K. V., and Hanchar, J. M.: Helium diffusion in
- zircon: Effects of anisotropy and radiation damage revealed by laser depth profiling, Geochim.
- 773 Cosmochim. Acta, 274, 45–62, https://doi.org/10.1016/j.gca.2020.01.049, 2020b.

- Barley, M. and Pickard, A.: An extensive, crustally-derived, 3325 to 3310 ma silicic
- volcanoplutonic suite in the eastern Pilbara craton: evidence from the Kelly Belt, Mcphee Dome
- and Corunna Downs Batholith, Precambrian Research, 96, 41–62, 1999.
- 777 Basu, A. R., Chakrabarty, P., Szymanowski, D., Ibañez-Mejia, M., Schoene, B., Ghosh, N., and
- 778 Georg, R. B.: Widespread silicic and alkaline magmatism synchronous with the Deccan Traps
- flood basalts, India, Earth Planet. Sci. Lett., 552, 116616,
- 780 https://doi.org/10.1016/j.epsl.2020.116616, 2020.
- 781
- 782 Bowring, S. A. and Schmitz, M. D.: High-precision U-Pb zircon geochronology and the
- stratigraphic record, in: Reviews in Mineralogy and Geochemistry Zircon, vol. 53, edited by:
  Hanchar, J. M. and Hoskin, P. W. O., 305–326, https://doi.org/10.2113/0530305, 2003.
- 785
- 786 Chakoumakos, B. C., Murakami, T., Lumpkin, G. R., and Ewing, R. C.: Alpha-decay induced
- fracturing in zircon: The transition from the crystalline to the metamict state, Science, 236,
  1556–1559, 1987.
- 788 789

792

- Cherniak, D. J.: Diffusion of helium in radiation-damaged zircon, Chem. Geol., 529, 119308,
  https://doi.org/10.1016/j.chemgeo.2019.119308, 2019.
- Cooperdock, E. H. G. and Stockli, D. F.: Unraveling alteration histories in serpentinites and
  associated ultramafic rocks with magnetite (U-Th)/He geochronology, Geology, 44, 967–970,
  https://doi.org/10.1130/g38587.1, 2016.
- 796
- Cooperdock, E. H. G. and Stockli, D. F.: Dating exhumed peridotite with spinel (U–Th)/He
  chronometry, Earth Planet. Sci. Lett., 489, 219–227, https://doi.org/10.1016/j.epsl.2018.02.041,
  2018.
- 800
- Cooperdock, E. H. G., Ketcham, R. A., and Stockli, D. F.: Resolving the effects of 2-D versus 3D grain measurements on apatite (U–Th) / He age data and reproducibility, Gchron., 1, 17–41,
  https://doi.org/10.5194/gchron-1-17-2019, 2019.
- 804
- Cooperdock, E. H. G., Hofmann, F., Collins, R. M., Carrera, A., Takase, A., and Celestian, A. J.:
  Technical note: Rapid phase identification of apatite and zircon grains for geochronology using
  X-ray micro-computed tomography, GChron., Preprint, https://doi.org/10.5194/gchron-2022-7,
  2022.
- 808 809
- Crowningshield, R. and Nassau, K.: The Heat and Diffusion Treatment of Natural and Synthetic
  Sapphires, J. Gemmology, 17, 528–541, https://doi.org/10.15506/jog.1981.17.8.528, 1981.
- 811 812
- 813 Danišík, M., McInnes, B. I. A., Kirkland, C. L., McDonald, B. J., Evans, N. J., and Becker, T.:
- 814 Seeing is believing: Visualization of He distribution in zircon and implications for thermal
- 815 history reconstruction on single crystals, Sci. Adv., 3, e1601121,
- 816 https://doi.org/10.1126/sciadv.1601121, 2017.
- 817

- Davydov, V. I., Crowley, J. L., Schmitz, M. D., and Poletaev, V. I.: High-precision U-Pb zircon 818 819 age calibration of the global Carboniferous time scale and Milankovitch band cyclicity in the 820 Donets Basin, eastern Ukraine, Geochem. Geophys. Geosyst., 11, n/a-n/a, 821 https://doi.org/10.1029/2009gc002736, 2010. 822 823 Ewing, R. C., Meldrum, A., Wang, L., Weber, W. J., and Corrales, L. R.: Radiation effects in 824 zircon, in: Reviews in Mineralogy & Geochemistry, vol. 53, edited by: Hanchar, J. M. and 825 Hoskin, P. W. O., 387-425, https://doi.org/10.2113/0530387, 2003. 826 827 Finch, R. J. and Hanchar, J. M.: Structure and chemistry of zircon and zircon-group minerals, in: 828 Reviews in Mineralogy and Geochemistry Zircon, vol. 53, edited by: Hanchar, J. M. and Hoskin, 829 P. W. O., 1–25, https://doi.org/10.2113/0530001, 2003. 830 831 Geisler, T., Pidgeon, R. T., Bronswijk, W. van, and Pleysier, R.: Kinetics of thermal recovery 832 and recrystallization of partially metamict zircon: a Raman spectroscopic study, Eur. J. Mineral., 13, 1163–1176, https://doi.org/10.1127/0935-1221/2001/0013-1163, 2001a. 833 834 835 Geisler, T., Ulonska, M., Schleicher, H., Pidgeon, R. T., and Bronswijk, W. van: Leaching and differential recrystallization of metamict zircon under experimental hydrothermal conditions, 836 837 Contrib. Mineral. Petr., 141, 53-65, https://doi.org/0.1007/s004100000202, 2001b. 838 839 Geisler, T., Pidgeon, R. T., Bronswijk, W. van, and Kurtz, R.: Transport of uranium, thorium, 840 and lead in metamict zircon under low-temperature hydrothermal conditions, Chem. Geol., 191, 841 141–154, https://doi.org/10.1016/s0009-2541(02)00153-5, 2002. 842 843 Ginster, U., Reiners, P. W., Nasdala, L., and N., C. C.: Annealing kinetics of radiation damage in zircon, Geochim. Cosmochim. Acta, 249, 225-246, https://doi.org/10.1016/j.gca.2019.01.033, 844 845 2019. 846 847 Gleadow, A. J. W., Hurford, A. J., and Quaife, R. D.: Fission track dating of zircon: Improved 848 etching techniques, Earth Planet. Sci. Lett., 33, 273-276, https://doi.org/10.1016/0012-849 821x(76)90235-1, 1976. 850 851 Guenthner, W. R., Reiners, P. W., Ketcham, R. A., Nasdala, L., and Giester, G.: Helium 852 diffusion in natural zircon: Radiation damage, anisotropy, and the interpretation of zircon (U-853 Th)/He thermochronology, Am. J. Sci., 313, 145–198, https://doi.org/10.2475/03.2013.01, 2013. Hanchar, J. M., Finch, R. J., Hoskin, P. W. O., Watson, E. B., Cherniak, D.J., Mariano, A. N.: 854 855 Rare earth elements in synthetic zircon: part 1. Synthesis, and rare earth element and phosphorus 856 doping, Am Mineral, 86, 667-680, 2001. 857 Härtel, B., Jonckheere, R., Wauschkuhn, B., and Ratschbacher, L.: The closure temperature(s) of 858 zircon Raman dating, GChron., 3, 259–272, https://doi.org/10.5194/gchron-3-259-2021, 2021. 859 860 Hazen, R. M. and Finger, L. W.: Crystal structure and compressibility of zircon at high pressure,
- 861 Am. Mineral., 64, 196–201, 1979.

862

863 Holland, H. D. and Gottfried, D.: The effect of nuclear radiation on the structure of zircon, Acta Crystallogr., 8, 291–300, https://doi.org/10.1107/s0365110x55000947, 1955. 864 865 866 Hovis, G., Abraham, T., Hudacek, W., Wildermuth, S., Scott, B., Altomare, C., Medford, A., 867 Conlon, M., Morris, M., Leaman, A., Almer, C., Tomaino, G., and Harlov, D.: Thermal expansion of F-Cl apatite crystalline solutions, Am. Mineral., 100, 1040-1046, 868 https://doi.org/10.2138/am-2015-5176, 2015. 869 870 871 Huyskens, M. H., Zink, S., and Amelin, Y.: Evaluation of temperature-time conditions for the 872 chemical abrasion treatment of single zircons for U-Pb geochronology, Chem. Geol., 438, 25-873 35, https://doi.org/10.1016/j.chemgeo.2016.05.013, 2016. 874 875 Itoh, N., and Shirono, K.: Reliable estimation of Raman shift and its uncertainty for a non-doped 876 Si substrate (NMIJ CRM 5606-a), J. Raman Spectroscopy, 51(12), 2496-2504, 877 http://doi.org/10.1002/jrs.6003, 2020. 878 879 Jonckheere, R.: On the densities of etchable fission tracks in a mineral and co-irradiated external 880 detector with reference to fission-track dating of minerals, Chem. Geol., 200, 41-58, 881 https://doi.org/10.1016/s0009-2541(03)00116-5, 2003. 882 883 Jonckheere, R. and Van den haute, P.: Observations on the geometry of etched fission tracks in 884 apatite: Implications for models of track revelation, Am. Mineral., 81, 1476–1493, 1996. 885 886 Jonckheere, R., Enkelmann, E., and Stübner, K.: Observations on the geometries of etched 887 fission and alpha-recoil tracks with reference to models of track revelation in minerals, Radiat. 888 Meas., 39, 577–583, https://doi.org/10.1016/j.radmeas.2004.08.008, 2005. 889 890 Jonckheere, R., Aslanian, C., Wauschkuhn, B., and Ratschbacher, L.: Fission-track etching in 891 apatite: A model and some implications, Am. Mineral., 107, 1190-1200, https://doi.org/10.2138/am-2022-8055, 2022. 892 893 894 Jones, S., Kohn, B., and Gleadow, A.: Etching of fission tracks in monazite: Further evidence 895 from optical and focused ion beam scanning electron microscopy, Am. Mineral., 107, 1065-896 1073, https://doi.org/10.2138/am-2022-8002, 2022. 897 898 Ketcham, R. A., Guenthner, W. R., and Reiners, P. W.: Geometric analysis of radiation damage connectivity in zircon, and its implications for helium diffusion, Am. Mineral., 98, 350-360, 899 900 https://doi.org/10.2138/am.2013.4249, 2013. 901 902 Krishnam, R. S.: Raman spectrum of quartz, Nature, 155, 142, 1945. 903 904 Lee, J. K. W. and Tromp, J.: Self-induced fracture generation in zircon, J. Geophys. Res. Solid 905 Earth, 100, 17753–17770, https://doi.org/10.1029/95jb01682, 1995. 906

MacLennan, S. A., Eddy, M. P., Merschat, A. J., Mehra, A. K., Crockford, P. W., Maloof, A. C., 907 908 Southworth, C. S., and Schoene, B.: Geologic evidence for an icehouse Earth before the Sturtian 909 global glaciation, Sci. Adv., 6, eaay6647, https://doi.org/10.1126/sciadv.aay6647, 2020. 910 911 Mattinson, J. M., Graubard, C. M., Parkinson, D. L., and McClelland, W. C.: U-Pb reverse 912 discordance in zircon: The role of fine-scale oscillatory zoning and sub-micron transport of Pb, in: Earth Processes Reading the Isotopic Code, edited by: Basu, A., and Hart, S., American 913 914 Geophysical Union, Washington D.C., USA, 355-370, 1996 915 916 Mattinson, J. M.: Zircon U-Pb chemical abrasion ("CA-TIMS") method: Combined annealing 917 and multi-step partial dissolution analysis for improved precision and accuracy of zircon ages, 918 Chem. Geol., 220, 47-66, https://doi.org/10.1016/j.chemgeo.2005.03.011, 2005. 919 920 Mattinson, J. M.: Extending the Krogh legacy: development of the CA-TIMS method for zircon U-Pb geochronology, Can. J. Earth Sci., 48, 95–105, https://doi.org/10.1139/e10-023, 2011. 921 922 923 Meldrum, A., Boatner, L. A., Weber, W. J., and Ewing, R. C.: Radiation damage in zircon and 924 monazite, Geochim. Cosmochim. Acta, 62, 2509–2520, https://doi.org/10.1016/s0016-925 7037(98)00174-4, 1998. 926 927 Meyers, S. R., Siewert, S. E., Singer, B. S., Sageman, B. B., Condon, D. J., Obradovich, J. D., 928 Jicha, B. R., and Sawyer, D. A.: Intercalibration of radioisotopic and astrochronologic time 929 scales for the Cenomanian-Turonian boundary interval, Western Interior Basin, USA, Geology, 930 40, 7–10, https://doi.org/10.1130/g32261.1, 2012. 931 932 Mezger, K. and Krogstad, E. J.: Interpretation of discordant U-Pb zircon ages: An evaluation, J. 933 Metamorph. Geol., 15, 127–140, https://doi.org/10.1111/j.1525-1314.1997.00008.x, 1997. 934 935 Mundil, R., Ludwig, K. R., Metcalfe, I., and Renne, P. R.: Age and timing of the Permian mass 936 extinctions: U/Pb dating of closed-system zircons, Science, 305, 1760-1763, 937 https://doi.org/10.1126/science.1101012, 2004. 938 939 Murakami, T., Chakoumakos, B. C., Ewing, R. C., Lumpkin, G. R., and Weber, W. J.: Alpha-940 decay event damage in zircon, Am. Mineral., 76, 1510-1532, 1991. 941 942 Nasdala, L., Irmer, G., and Wolf, D.: The degree of metamictization in zircon: a Raman 943 spectroscopic study, Eur. J. Mineral., 7, 471–478, https://doi.org/10.1127/ejm/7/3/0471, 1995. 944 945 Nasdala, L., Pidgeon, R. T., Wolf, D., and Irmer, G.: Metamictization and U-Pb isotopic discordance in single zircons: a combined Raman microprobe and SHRIMP ion probe study, 946 947 Mineral. Petrol., 62, 1–27, https://doi.org/10.1007/bf01173760, 1998. 948 949 Nasdala, L., Wenzel, M., Vavra, G., Irmer, G., Wenzel, T., and Kober, B.: Metamictisation of 950 natural zircon: accumulation versus thermal annealing of radioactivity-induced damage, Contrib. 951 Mineral. Petr., 141, 125–144, https://doi.org/10.1007/s004100000235, 2001. 952

Nasdala, L., Reiners, P. W., Garver, J. I., Kennedy, A. K., Stern, R. A., Balan, E., and Wirth, R.: 953 954 Incomplete retention of radiation damage in zircon from Sri Lanka, Am. Mineral., 89, 219–231, 955 2004. 956 957 Nassau, K.: Heat treating ruby and sapphire: Technical aspects, Gems Gemol., 17, 121–131, 958 https://doi.org/10.5741/gems.17.3.121, 1981. 959 960 Paces, J. B. and Miller, J. D.: Precise U-Pb ages of Duluth Complex and related mafic intrusions, 961 northeastern Minnesota: Geochronological insights to physical, petrogenetic, paleomagnetic, and 962 tectonomagmatic processes associated with the 1.1 Ga Midcontinent Rift System, J. Geophys. 963 Res. Solid Earth, 98, 13997–14013, https://doi.org/10.1029/93jb01159, 1993. 964 965 Palenik, C. S., Nasdala, L., and Ewing, R. C.: Radiation damage in zircon, Am. Mineral., 88, 966 770-781, https://doi.org/10.2138/am-2003-5-606, 2003. 967 968 Schmitz, M. D. and Davydov, V. I.: Quantitative radiometric and biostratigraphic calibration of 969 the Pennsylvanian–Early Permian (Cisuralian) time scale and pan-Euramerican 970 chronostratigraphic correlation, GSA Bulletin, 124, 549-577, https://doi.org/10.1130/b30385.1, 971 2012. 972 973 Schmitz, M. D., Bowring, S. A., and Ireland, T. R.: Evaluation of Duluth Complex anorthositic 974 series (AS3) zircon as a U-Pb geochronological standard: new high-precision isotope dilution 975 thermal ionization mass spectrometry results, Geochim. Cosmochim. Acta, 67, 3665-3672, 976 https://doi.org/10.1016/s0016-7037(03)00200-x, 2003. 977 978 Schoene, B.: Treatise on Geochemistry (Second Edition), in: Treatise on Geochemistry, vol. 4, 979 edited by: Holland, H. D. and Turekian, K. K., Treatise on Geochemistry, 341-378, 980 https://doi.org/10.1016/b978-0-08-095975-7.00310-7, 2014. 981 982 Schoene, B., Guex, J., Bartolini, A., Schaltegger, U., and Blackburn, T. J.: Correlating the end-983 Triassic mass extinction and flood basalt volcanism at the 100 ka level, Geology, 38, 387-390, 984 https://doi.org/10.1130/g30683.1, 2010a. 985 986 Schoene, B., Latkoczy, C., Schaltegger, U., and Günther, D.: A new method integrating high-987 precision U-Pb geochronology with zircon trace element analysis (U-Pb TIMS-TEA), Geochim. 988 Cosmochim. Acta, 74, 7144–7159, https://doi.org/10.1016/j.gca.2010.09.016, 2010b. Smithies, R. H., Champion, D. C., and Cassidy, K. F.: Formation of Earth's early Archaean 989 990 continental crust: Precambrian Research, 127, 89-101, 2003. 991 Subbarao, E. C., Agrawal, D. K., McKinstry, H. A., Sallese, C. W., and Roy, R.: Thermal 992 expansion of compounds of zircon structure, J. Am. Ceram. Soc., 73, 1246–1252, 993 https://doi.org/10.1111/j.1151-2916.1990.tb05187.x, 1990. 994

Swanson-Hysell, N. L., Hoaglund, S. A., Crowley, J. L., Schmitz, M. D., Zhang, Y., and Miller, 995 996 J. D.: Rapid emplacement of massive Duluth Complex intrusions within the North American Midcontinent Rift, Geology, 49, 185–189, https://doi.org/10.1130/g47873.1, 2020. 997 998 999 Takehara, M., Horie, K., Hokada, T., and Kiyokawa, S.: New insight into disturbance of U-Pb 1000 and trace-element systems in hydrothermally altered zircon via SHRIMP analyses of zircon from the Duluth Gabbro, Chem. Geol., 484, 168–178, https://doi.org/10.1016/j.chemgeo.2018.01.028, 1001 1002 2018. 1003 1004 Trachenko, K., Dove, M. T., and Salje, E. K. H.: Structural changes in zircon under  $\alpha$ -decay 1005 irradiation, Phys. Rev. B, 65, 180102, https://doi.org/10.1103/physrevb.65.180102, 2002. 1006 1007 Váczi, T.: A new, simple approximation for the deconvolution of instrumental broadening in spectroscopic band profiles, Appl. Spectrosc., 68, 1274–1278, https://doi.org/10.1366/13-07275, 1008 1009 2014. 1010 1011 Váczi, T. and Nasdala, L.: Electron-beam-induced annealing of natural zircon: a Raman 1012 spectroscopic study, Phys. Chem. Mineral., 44, 389-401, https://doi.org/10.1007/s00269-016-1013 0866-x, 2017. 1014 van Kranendonk, M. J., Hugh Smithies, R., Hickman, A. H., and Champion, D.: Review: secular 1015 tectonic evolution of Archean continental crust: interplay between horizontal and vertical processes in the formation of the Pilbara Craton, Australia, Terra Nova, 19, 1–38, 1016 1017 doi:10.1111/j.1365-3121.2006.00723.x, 2007. 1018 Weber, W. J.: Radiation-induced defects and amorphization in zircon, J. Materials Res., 5, 2687-1019 2697, 1990. 1020 Widmann, P., Davies, J. H. F. L., and Schaltegger, U.: Calibrating chemical abrasion: Its effects 1021 1022 on zircon crystal structure, chemical composition and U-Pb age, Chem. Geol., 511, 1–10, 1023 https://doi.org/10.1016/j.chemgeo.2019.02.026, 2019. 1024 1025 Yamada, R., Tagami, T., Nishimura, S., and Ito, H.: Annealing kinetics of fission tracks in 1026 zircon: an experimental study, Chem. Geol., 122, 249-258, https://doi.org/10.1016/0009-1027 2541(95)00006-8, 1995. 1028 1029 Yamada, R., Murakami, M., and Tagami, T.: Statistical modelling of annealing kinetics of 1030 fission tracks in zircon; Reassessment of laboratory experiments, Chem. Geol., 236, 75-91, 1031 https://doi.org/10.1016/j.chemgeo.2006.09.002, 2007. 1032 1033 Zhang, M., Salje, E. K. H., Capitani, G. C., Leroux, H., Clark, A. M., Schlüter, J., and Ewing, R. 1034 C.: Annealing of alpha-decay damage in zircon: a Raman spectroscopic study, J. Phys. Condens. 1035 Matter, 12, 3131, https://doi.org/10.1088/0953-8984/12/13/321, 2000.