

Uncertainty in Helium Diffusion in Zircon Limits Thermochronologic Resolution: Application to The Great Unconformity

Matthew Fox¹, Adam G.G Smith¹, Pieter Vermeesch¹, Kerry Gallagher², Andrew Carter^{1,3}

¹ London Geochronology Centre, Department of Earth Sciences, University College London, Gower St., London, WC16BT, UK

² Géosciences Rennes (OSEREN), University of Rennes, Rennes, France

³ School of Natural Sciences, Birkbeck College, Malet St., London, WC1E 7HX, UK

Correspondence to: Matthew Fox (m.fox@ucl.ac.uk)

Abstract. Thermochronology provides a unique perspective on the timing and magnitude of erosion during the generation of unconformities. Recently, thermochronology has been used to reinvigorate a long-standing debate about the origin of the Great Unconformity, a geological feature that has eliminated a time period of almost a billion years from the stratigraphic record at the end of the Precambrian. Our ability to resolve thermal histories is fundamentally limited by how well we know parameters describing the temperature sensitivity of a thermochronometric system. The (U-Th)/He in zircon system is particularly well suited to examine the origin of the Great Unconformity because it is very sensitive to long durations of time at relatively low temperatures (< 200-250°C). Here we determine uncertainties in the Zircon Radiation Damage and Annealing Model (ZRDAAM, Guenther et al. 2013) that describes changes in ⁴He diffusion kinetics as a function of radiation damage accumulation and annealing. We show that the dispersion in predicted zircon (U-Th)/He ages for a given thermal history can be 100's of Ma for a specific amount of radiation damage highlighting that thermal histories are less well resolved than previously appreciated. Additional diffusion experiments and calibration with natural laboratories would provide better constraints on diffusion kinetic parameters.

1. Introduction

Thermochronometry is widely used to constrain the evolution of Earth's surface and upper crust, transforming our understanding of the magnitudes of sedimentation and erosion linked to climate and tectonic change. In many cases, the timings of specific events are of particular importance. The principle behind thermochronometry is that rocks experience temperature changes over geological time-scales: rocks closer to Earth's surface are cooler than deeper ones and so exhumation leads to

Deleted: Origin of

Deleted: Obscured by

Deleted: Thermochronometric Uncertainty

Deleted: Carter³

Deleted: OSUR

Deleted: London Geochronology Centre, Department

Deleted: Earth

Deleted: Gower

Deleted: WC16BT

Deleted: rock that is eroded

Deleted: , and

Deleted: timing

Deleted: ,

Deleted: in the rock record

Deleted: global erosional event

Deleted: represents

Deleted: provide this perspective

Deleted: However, the diffusion kinetics of ⁴He in zircon change dramatically as a result of radiation damage to the crystal lattice. Therefore, our ability to resolve thermal histories is fundamentally limited by how well we know parameters controlling helium diffusion and their uncertainties. ... [1]

Deleted: .

Deleted: 100s

Deleted: and

Moved (insertion) [1]

Deleted: Notably, thermochronometry has been used re- ... [2]

75 **cooling.** The time scales associated with these changes in temperature are determined using the concepts of geochronology and diffusion. **The daughter products of radioactive decay used for thermochronometry have a temperature dependent rate of loss from the target mineral. At high temperatures the daughter products are effectively lost instantaneously.** By contrast, as the rock cools to lower temperatures, the rate of diffusive loss decreases and daughter products are progressively retained until there is effectively no diffusive loss. Therefore a measured age reflects the duration of residence at low temperature (Dodson, 1973; Zeitler et al., 1987; Reiners, 2005; Fox and Carter, 2020), however many thermal histories, that incorporate reheating events will produce the same age.

Moved (insertion) [2]

85 **One of the most widely used methods for deep-time thermochronometry is (U-Th)/He in zircon (ZHe) (Reiners, 2005). The basis of this method is that the radioactive elements uranium and thorium are incorporated into crystals of zircon and the decay product, helium, is trapped in the crystal lattice. Helium diffuses from zircon at high temperatures, but is retained at lower temperatures.** The exact temperature range at which the transition from open-system to closed-system behaviour occurs is dependent on damage to the crystal lattice produced primarily from recoil during decay processes (Guenther et al., 2013; Ketcham et al., 2013). **Radiation damage accumulates at a rate that depends on the amount of uranium, thorium, and samarium in the crystal. This means that two zircon crystals from the same rock, experiencing the same thermal history, could have very different thermal sensitivity. In turn, models accounting for this variable temperature sensitivity as a function of radiation damage can be used to leverage more complex thermal histories than constant kinetic parameter models.**

Moved (insertion) [3]

Moved (insertion) [4]

Thermochronometry has been used recently to understand the development of the Great Unconformity (e.g., DeLucia et al., 2018; Flowers et al., 2020; McDannell et al., 2022), a global feature that marks the boundary between the Precambrian and Phanerozoic. The time-period that the Great Unconformity spans varies depending on location, but on the North American craton, in the Grand Canyon, the erosional event removed stratigraphy between 1300 to 250 million years (Thurston et al., 2022). The origin of the Great Unconformity has been debated for over 125 years, with recent contributions using thermochronometry to gain new insight (McDannell et al., 2022). This approach **inferred** that erosion rates increased across the North American craton during Neoproterozoic glaciation, supporting the hypothesis that the Great Unconformity is the result of glacial erosion (Keller et al., 2019), as opposed to diachronous tectonic events (Flowers et al., 2020).

Deleted: highlighted

Field Code Changed

Both those in favour of a globally, reasonably, synchronous glacial origin (McDannell et al., 2022) and regionally diachronous tectonic origins (Flowers et al., 2020) of the Great Unconformity have interpreted ZHe data using the Zircon Radiation Damage And Annealing Model (ZRDAAM) of Guenther et al. (2013). Importantly, it is clear that it is always better to combine multiple thermochronometric systems as this helps resolve thermal histories and reduces the negative effects of systematic uncertainties for specific systems (McDannell et al., 2022). In ZRDAAM model, a crystal is composed of undamaged and damaged parts that combine to give bulk diffusion kinetics as a function of the amount of radiation damage. Accumulated radiation damage is calculated based on the concentrations of the parent elements and the thermal history of a sample. The damage accumulates at low temperatures, but can be annealed at higher temperatures, calculated using fission track annealing kinetics (Yamada et al., 1995; Rahn et al., 2004; Tagami, 2005; Yamada et al., 2007). Therefore, the diffusivity of helium at a specific temperature is a function of the past thermal history. This makes the overall problem very non-linear so that changing the temperature at some time in the thermal history can have unexpected effects on the resulting age, as also shown for the (U-Th)/He in apatite system by Fox and Shuster (2014).

Using the zircon radiation damage and annealing model (ZRDAAM) with inverse models, researchers have resolved tight temperature constraints on thermal histories over billion-year timescales. For example, Thurston et al. (2022) inferred a 1700 Ma thermal history from ZHe ages in Eastern Grand Canyon. Parts of this history were reported to within less than 10 degrees between 700 and 250 Ma and then again from 15 -7 Ma. It is unclear whether the data really provide such tight constraints on temperatures in the past or whether these are at least partly the consequence of not sufficiently sampling parameter space, model assumptions and/or, potentially, overconfidence in the adopted diffusion kinetic parameters. These aspects are explored in this study. Specifically, we want to know how uncertainty in the estimates of diffusion kinetic parameters may impact on the inference of thermal histories and the associated uncertainties.

Calibration of ZRDAAM has been carried out using measured diffusion kinetics of crystals with well constrained amounts of radiation damage (Guenther et al., 2013). However, the accuracy and precision of this model has not been assessed. In particular, it is unclear how the propagation of uncertainties to model parameters affects the dispersion or sensitivity of predicted thermochronometric ages. Here we show that the uncertainties in the radiation damage model make it challenging to accurately infer the timing and magnitude of unconformities in the deep past. We begin by highlighting why

Moved up [1]: The principle behind thermochronometry is that rocks experience temperature changes over geological timescales: rocks closer to Earth's surface are cooler than deeper ones and so exhumation leads to cooling.

Moved up [2]: The daughter products of radioactive decay used for thermochronometry have a temperature dependent rate of loss from the target mineral. At high temperatures the daughter products are effectively lost instantaneously.

Moved up [3]: One of the most widely used methods for deep-time thermochronometry is (U-Th)/He in zircon (ZHe) (Reiners, 2005). The basis of this method is that the radioactive elements uranium and thorium are incorporated into crystals of zircon and the decay product, helium, is trapped in the crystal lattice.

Moved up [4]: . Radiation damage accumulates at a rate that depends on the amount of uranium, thorium, and

Deleted: The time scales associated with these changes in temperature are determined using the concepts of geochronology. ... [3]

Deleted: By contrast, as the rock cools to lower temperatures, the rate of diffusive decreases and daughter products ... [4]

Deleted: The exact temperature range at which the transition from closed-system to open-system behaviour occurs is ... [5]

Deleted: This approach has successfully been used to infer deep-time erosion rate histories, and in the case of the Great ... [6]

Deleted: origin

Deleted: In this

Deleted: RDAAM

Deleted:

Deleted: known

ZRDAAM needs to be calibrated by accounting for uncertainties and present a new calibration. We then propagate uncertainties from this model calibration into time temperature path uncertainty.

Deleted: our

Deleted: through

Deleted: paths. We show that natural variability in radiation damage annealing parameters causes ZHe ages to be very dispersed even for crystals of the same size and radiation damage levels. Using QTQt (Gallagher, 2012), we show that different diffusion kinetics can lead to the onset of cooling for resolved thermal histories from inverse methods varying by hundreds of millions of years.

Deleted: and annealing

205 2. The existing calibration of the radiation damage model

The rate of diffusion is controlled by the diffusivity and the curvature of the concentration of the diffusant (Fick's Law). Although the production distribution of the diffusant (helium in our case) can be important in some scenarios, diffusion tends to smooth the distribution. More significant is the fact that the diffusivity can vary by orders of magnitude with variations in temperature. The diffusivity at any

210 temperature is given by the Arrhenius equation of the diffusivity D (cm^2/s):

$$D(T) = D_0 e^{\frac{-E_a}{RT}} \quad (1)$$

where D_0 is the frequency factor (cm^2/s), E_a is the activation energy (kJ/mol), R is the gas constant

215 (J/K/mol) and T is the temperature in Kelvin. Taking the logarithm of equation 1, gives:

$$\ln(D(T)) = \frac{-E_a}{R} \frac{1}{T} + \ln(D_0) \quad (2)$$

so that the slope of the line between $\ln(D(T))$ and $1/T$ gives E_a/R and the intercept of the line provides $\ln(D_0)$. Diffusion experiments in which a crystal is step-wise degassed *in vacuo* are used to calculate

220 $D(T)$ for different combinations of specific temperatures and time (Fechtig and Kalbitzer, 1966). The resulting plot can be used to determine the Arrhenius parameters (D_0 , E_a). Importantly, estimates of the two model parameters extracted from this linear inversion covary with one another, i.e., $\ln(D_0)$ is strongly correlated with E_a .

225 Analyzing different zircon crystals with known radiation damage values allows us to assess how diffusion kinetics vary with damage. As it is challenging to visualize both model parameters (D_0 and E_a) for each crystal as a function of radiation damage, it is common to combine the parameters and calculate a diffusivity at a specific temperature or a closure temperature at a specific cooling rate and grain size.

By combining the parameters, however, information on how the two parameters (D_0 and E_a) are correlated is lost. In turn, our ability to accurately infer thermal histories is reduced.

230

Deleted: It

Deleted: and so

Deleted: .

The results of Guenther et al., (2013)'s diffusion experiments highlight two general trends. At low damage values the closure temperature increases with increasing damage. At higher damage values, the

closure temperature decreases with increasing damage. This general behaviour has been reproduced in numerical models conducted at a range of scales (Ketcham et al., 2013; Gautheron et al., 2020). To interpret these trends in ZRDAAM, a model is used in which the diffusion kinetics for a specific radiation damage value are a combination of a theoretical minimally damaged crystal and an extremely damaged crystal. The diffusion kinetics of these end-member crystals need to be estimated. The frequency factor of the minimally damaged, *theoretical*, crystal, (D_0) was estimated by extrapolating the frequency factors of measured crystals down two orders of magnitudes using a power-law relationship. The activation energy for *the* theoretical crystal (E_a) is set as the average of the activation energies of minimally damaged crystals (see Guenthner et al., 2013 for details). Extrapolating values to a minimally damaged crystal, however, will add uncertainties and there is no obvious way to account for these in the power-law relationship. Crucially, this approach does not account for the correlations between the model parameters. This is important because the correlations provide additional information that can yield more precise estimates of model parameters and allow propagation of uncertainties into model predictions. The diffusion kinetics for the extremely damaged crystal are estimated using sample N17 (Guenthner et al., 2013), and also involve correlated model parameters ($^{N17}D_0$ and $^{N17}E_a$). However, the accuracy of this model has only been assessed by looking at general trends in model predictions (Guenthner et al., 2013). Here we attempt to formally quantify the uncertainty in *damage* model parameters and how these uncertainties translate to uncertainties in temperature sensitivity of the ZHe system, and in particular predicted ZHe ages. *We note that additional work, informed by our exploration, could be carried on quantifying uncertainties in other aspects of the ZRDAAM.*

3. A new calibration of the zircon radiation damage and annealing model

In order to account for the correlation between the frequency factor and the activation energy, we model the measured helium diffusivities directly. We use the same diffusion *dataset* and parameterisation as Guenthner et al. (2013). However, in contrast with Guenthner et al. (2013), we determine the diffusion of the end member crystals using the radiation damage model directly, rather than by non-linear extrapolation from high to low radiation damage levels. Our goal is not to simply increase the accuracy of the model parameters but to determine their precision. By tracking correlations in model parameters, *that are the result of correlations within individual diffusion experiments*, we can simulate (U-Th)/He in zircon ages accounting for uncertainties in the original diffusion experiments.

Deleted: .

Deleted: data

The model we fit through each diffusion data set is given by equation 8 of Guenther et al. (2013) and describes the diffusivity D as a function of the amount of damage:

$$\frac{1}{D} = \frac{f'_c}{\frac{1}{\left(\frac{l_{int0}}{l_{int}}\right)^2} * \left(\frac{D_z}{(a * f'_c)}\right)} + \frac{f'_a}{\frac{D_{N17}}{(a * f'_a)^2}}$$

(4)

where f'_c and f'_a are the crystalline and amorphous fractions, respectively, l_{int} and $l_{int,0}$ are parameters describing how far a helium atom can travel within a crystal lattice without encountering damage in a damaged crystal and an undamaged crystal respectively, and a is the grain size, or more accurately, the equivalent spherical radius. D_z and D_{N17} can be calculated using the diffusion kinetics of the undamaged and damaged theoretical crystals, using equation 4. Therefore, for every sample with an estimated amount of damage, we can calculate different diffusivity values for degassing steps using four model parameters $^{N17}D_0$ and $^{N17}E_a$ for D_{N17} and 2D_0 and 2E_a for D_z . In this way, the model predicts the diffusivity at a specific temperature of the data using the model parameters and we do not rely on linear regressions through individual Arrhenius relationships. The log likelihood function is defined as the negative least-squares misfit when fitting all the diffusion data for a given set of model parameters. However, the degassing experiments of Guenther et al. (2013) each have different numbers of steps. The likelihood function involves a summation, related to the number of steps, and so number of data points in each experiment, experiments with more steps would tend to dominate our results. Similarly, key experiments which might have fewer steps would have far less influence on the model parameter estimation. To account for this problem, the least squares misfit for each experiment is weighted accordingly, so that the log-likelihood (LL) function is:

$$LL = -0.5 * \sum_{i=1}^N \frac{1}{M_i} \sum_{j=1}^{M_i} \left(\frac{D_{i,j} - P_{i,j}}{\sigma} \right)^2$$

(5)

where N is the number of crystals analyzed, M_i is the number of degassing steps used in the inversion for a specific crystal, $D_{i,j}$ is the observed diffusivity for a specific crystal at a specific degassing step, $P_{i,j}$ is corresponding predicted diffusivity calculated with Equation 4, and σ is the estimated uncertainty set to $1/\ln(1/s)$ here, based on reported uncertainties.

Deleted: .

Deleted: a known amount of damage, we can calculate different diffusivity values for degassing steps using model parameters $^{N17}D_0$ and $^{N17}E_a$ for D_{N17} and 2D_0 and 2E_a for D_z . We use the Bayesian Markov Chain Monte Carlo (MCMC) method incorporating the Metropolis-Hastings algorithm to sample the full posterior distribution of the model parameters. We tune the proposal distributions

Moved down [5]: to ensure that approximately 20% of the proposed models are accepted as this represents an efficient balance between exploring parameter space and sampling the parameter values. The Markov Chain is initialised with the model parameters of Guenther et al. (2013) and the algorithm runs until 1 million sets of model parameters have been accepted.

Deleted: The likelihood function is defined as a least squares fit to the data.

Formatted: Font: 10 pt, Font colour: Text 1

Formatted: Font: 10 pt, Font colour: Text 1

Formatted: Font: 10 pt, Font colour: Text 1

Formatted: Font: 10 pt, Font colour: Text 1

Formatted: Font: 10 pt

Formatted: Left

Formatted: Font: 10 pt, Font colour: Text 1

We use the Bayesian Markov Chain Monte Carlo (MCMC) method incorporating the Metropolis-Hastings algorithm to sample the full posterior distribution of the model parameters ($^{N17}D_0$ and $^{N17}E_a$ for D_{N17} and 2D_0 and 2E_a). We tune the proposal distributions (which determine how far to perturb model parameters between iterations) **to ensure that approximately 20% of the proposed models are accepted as this represents an efficient balance between exploring parameter space and sampling the parameter values. The Markov Chain is initialised with the model parameters of Guenther et al. (2013) and the algorithm runs until 1 million sets of model parameters have been accepted.** In this way, at each iteration we: 1) propose four model parameters by perturbing the current model parameters; 2) calculate an Arrhenius relationship for each diffusion experiment using the four proposed model parameters and calculate the likelihood; 3) accept the proposed model parameters if they improve the likelihood or accept with a probability determined by the ratio of the likelihoods; 4) update the current model parameters to the proposed model parameters if they are accepted. The advantage of this approach over simply extracting diffusion parameters from individual experiments is that that uncertainty on diffusion data and model correlations are propagated through to uncertainty on model parameters. Preliminary experiments highlighted that the sampling was relatively insensitive to the diffusion kinetics of the damaged N17 crystal. To ensure that the diffusion kinetics of this unique and crucial end member crystal was accurately captured, we reduced the uncertainty of the diffusion data for N17 to 0.1 ln(1/s). **If this uncertainty is not decreased, the diffusion measurements of this crystal are not accurately predicted.**

Moved (insertion) [5]

The comparison of the predicted and observed diffusivities for the best fitting model is shown in Figure 1B. We also show the model fit with the original Guenther et al., (2013) parameters in Figure 1A. From these results, it is clear that neither set of model parameters fit the data perfectly and this is expected with noisy data. However, our model parameters do provide a better fit to the data based on our misfit criterion in equation 5. Results of our analysis are plotted as 4 histograms showing the original model parameters and our inferred model parameters (Figure 2). **These are histograms of model parameters sampled by our MCMC algorithm. These are known as the marginal posterior distributions for each model parameter. The probability of the model parameters is proportional to the height of the model histograms. For the amorphous crystal, our maximum a posteriori model parameter values are close to the original values, shown in red. However, for the low-damaged crystal, the model estimates are quite different, although we note that the original values fall along the strong correlation trend seen in the 2D probability distribution. The original values are a little displaced from the peak (Figure 3)**

Deleted: The

Deleted: predictions are quite different, although we note that the original values are close to the 2D probability peak (Figure 3) which accounts for model correlations. Additionally, we will focus on the overall spread in the values of these model parameters and we explore the importance of this variability in the next section.

showing that these parameters do not fit the data quite as well, but the combination of values are consistent with the linear relationship between 2E_a and 2D_0 . We note also that while there is a strong correlation between $^{N17}D_0$ and $^{N17}E_a$ and between 2D_0 and 2E_a the 2 sets of parameters are not strongly correlated (i.e. $^{N17}D_0$ and $^{N17}E_a$ are effectively independent of 2D_0 and 2E_a). Approximating the posterior distribution as a four-dimensional Gaussian distribution provides the mean values of the parameters, the covariance matrix and the correlation matrix. Below, the parameters are ordered as $^{N17}E_a$, $\log_{10}(^{N17}D_0)$, 2E_a and $\log_{10}(^2D_0)$. The mean values for the four parameters are:

5.61e+04, -3.75e+00, 1.42e+05, 3.54e+00,
with a model covariance matrix of:
3.87e+09, 4.08e+05, 1.62e+06, -2.60e+02,
4.08e+05, 4.31e+01, 1.92e+02, -2.62e-02,
1.62e+06, 1.92e+02, 5.14e+07, 3.86e+03,
-2.60e+02, -2.62e-02, 3.86e+03, 2.95e-01.

In addition to the covariance matrix, it is useful to inspect to the correlation matrix as it easier to see how parameter values correlate without the variable units. The correlation matrix shows strong correlations across some parameters but very low correlations across others:

1., 0.999, 0.004, -0.008
0.999, 1., 0.004, -0.007
0.004, 0.004, 1., 0.991
-0.008, -0.007, 0.991, 1.

Here we can clearly see the strong correlation between the Arrhenius parameters of the two endmember crystals. However, there is relatively little correlation between the different endmember grains.

Formatted: Don't hyphenate

Formatted: Font colour: Auto

4. Propagating model uncertainties

To assess the importance of the uncertainties in the radiation damage and annealing model parameters, we predict ages using a simple thermal model. The time-temperature path is chosen to resemble that of the Minnesota samples from Miltich (2005) and McDannell et al. (2022), and represents a typical inferred time temperature path of a Deep Time target locality. Here the rocks have been below 600°C since 1.5 Ga. From 700 Ma to 650 Ma the rocks cooled from 200 to 150°C, and then gradually to 0°C by 200 Ma before experiencing reheating at 50 Ma to 100°C. Between 50 Ma and the present the rocks

Deleted: McDannell et al. (2022),

cooled linearly to 0°C. Radiation damage accumulates throughout this history such that some zircon crystals that transitioned from open to closed system behaviour during the cooling event at 700 Ma, transitioned back to open behaviour simply due to the accumulation of radiation damage. Some crystals

405 however, with intermediate temperature sensitivity only record the final cooling event. Other crystals, with low-temperature sensitivity, also record cooling associated with the 100 Ma burial event. In terms of constraining thermal history models, this potential to have a wide range of temperature sensitivities within a single sample makes the ZHe method very powerful.

Deleted:

410 To calculate thermochronometric ages we use the radiation damage and annealing model of Guenther et al. (2013) with our updated model parameters. Note, the numerical methods used to implement the diffusion equations and damage models have been used previously (Tripathy-Lang et al., 2015). 20 different crystals are simulated spanning an effective U concentration ($[eU] = [U] + 0.24[Th]$; Cooperdock et al., 2019; Gastil et al., (1967)) interval from 31 to 2828 ppm.

Deleted: Note, this implementation of the model has

415 Initially, the Guenther et al. (2013) model parameters were used to simulate ages. In general, the ages are similar, however, if a reasonable analytical uncertainty of 10% of the age is used, there are parts of the age-[eU] relationship that are distinct. Grain sizes are all set to 70 µm.

Deleted: Ages for these

420 Next, uncertainties from our radiation damage model calibration were propagated through to predict age distributions. To do this 20 different crystal ages with different [eU] values were calculated 1000 times with different model parameters for $^{17}D_0$, $^{17}E_a$, 2D_0 and 2E_a . To do this, we extracted every 10th model from the posterior ensemble of models generated during the MCMC algorithm. This ensures that we are sampling model parameter space in proportion to probability but also that the model correlations are reliably captured. As before the grain sizes were all set to 70 µm.

Deleted: crystals are

Deleted: 200

Deleted: 100th

Deleted: Grain

Deleted: are

Deleted: single

The results highlight the large spread in predicted ages for a representative thermal history, typical for exploiting the ZRDAAM (Figure 4). The overall spread in age between the youngest age and the oldest age is expected given the different temperature sensitivity of the crystals due to radiation damage.

430 However, even for a specific amount of radiation damage there is still a large dispersion in the predicted ages. For example, at [eU] values of about 1600 ppm, ages are expected to vary between 50 and 550 Ma. If a range of grain sizes were also modelled for a specific [eU], the spread would be even larger (Whipp et al., 2021). Furthermore, a myriad of other factors will also contribute to this dispersion (see Fox et al., 2019 for a discussion).

Deleted: .

445

5. Incorporating model parameter uncertainty in inverse models

To assess the effect of imperfectly defined kinetic parameters in the ZHe radiation damage model of Guenther et al., (2014), we modified QTQt to allow sampling of the 4 parameters ($^{N17}D_0$, $^{N17}E_a$, $^{20}D_0$ and $^{20}E_a$) from the joint posterior distribution. The resampling draws values from the 4-dimensional Gaussian with the empirical mean vector and covariance matrix calculated from the joint posterior obtained, as described above. The assumption of Gaussian is not strictly true given the form of the sampled distributions, particularly $\log(^{N17}D_0)$ and $^{N17}E_a$, but the goal is to sample over the range of values appropriately, and then this approximation is good enough. We sample from a zero mean, unit variance normal distribution, use a Cholesky decomposition of the covariance matrix to rescale those values and then add the means, using the triangular factorization method described in Barr and Slezak (1972). Note the covariance matrix incorporates the correlation between the parameters and thus so do the sampled values. We used age uncertainties of 25% for the ZHe age values reported in McDannell et al., (2022) as this uncertainty allowed us to reproduce the characteristics of the thermal history reported in McDannell et al. (2022). We ran QTQt for 2 million iterations with the first million being discarded as burn-in. We ran the model with the original kinetic parameters (no resampling) and with resampling of the 4 kinetic parameters from the joint posterior every 500 iterations (and over the 2 million iterations this resampling effectively reproduces the original posterior distribution of these model parameters). The resampling reproduces the distributions and correlations for the 4 kinetic parameters. However, we do not attempt to infer the optimal value of these parameters, instead we simply incorporate this model parameter uncertainty into the inverse thermal history models and their predictions. Importantly for a given thermal history we use the same radiation damage parameter values for all the grains but due to the different degrees of damage, the resulting diffusion parameters are different for each crystal. The results are given in figure 5a and 5b, which has a summary of the inferred thermal histories. The modelled ages predict the observed ages for the expected models as shown in McDannell et al. (2022). We see that the expected and maximum likelihood thermal histories for the two different kinetic parameter models have a similar form in terms of the timing of cooling events, some variation in the magnitude of cooling and a different range for the credible intervals around the first period of cooling (800-700 Ma). The most significant differences are in the form of the maximum posterior models and the distribution of accepted thermal histories. In particular, the resampling model suggests a second mode or group of solutions that do not have a second period of cooling (around 400-350 Ma). This does not indicate that the solution has not converged, instead the solution has converged to reveal these two groups. We note

Deleted: In order to highlight how uncertainty in diffusion parameters propagate to uncertainty in thermal histories, we use QTQt (Gallagher 2012) and the data from McDannell et al. (2022). It is important to note that our goal is not to determine a new thermal history from the data, but rather to assess uncertainties related to the kinetics. For this reason, we only use only the ZHe data and do not incorporate additional constraints. We use the 22 ZHe single grain age data for the sample “Minnesota” of McDannell et al. (2022), and 2 sets of values for the 4 diffusion parameters based on the MCMC sampling. These correspond to values of 129224.284308, 4.000427, 149756.087043, 4.130064 and 42388.297312, - 5.000653 137904.535009, 3.253373 corresponding to $^{N17}E_a$, $\log10(^{N17}D_0)$, $^{20}E_a$ and $\log10(^{20}D_0)$ for a high value for the frequency factor of N17 and a low value, respectively. We also use the values originally proposed by Guenther et al. (2013) as a reference. The priors for time and temperature were specified to be 1500 ± 1500 M.y. and $150 \pm 150^\circ\text{C}$ respectively. We ran the sampler for 300k iterations burn-in and 300k post-burn-in, accepting models in the conventional MCMC way, such that a proposed model of equal likelihood to the current model will be accepted irrespective of the complexity. Results show that while the general trend of the cooling is very similar, the posterior probabilities are all quite different (Figure 5). This has important implications for our ability to reliably identify cooling signatures. In particular, the part of the thermal history that appears well resolved by the data changes from 1000 Ma to 1500 Ma depending on the choice of radiation damage parameters. This suggests that the resolution in estimating the timing of a given cooling event could be as large as 500 Ma, with obvious implications for resolving the timing of the Great Unconformity. ¶

¶

... [7]

545 that, for the “no resampling” model, the maximum likelihood and maximum posterior models are similar, while these two models tend to fall on two separate modes for the resampling model. In this latter case, the maximum posterior model is a little simpler (fewer time-temperature points) and does not reproduce the observations (lower likelihood) quite as well as the no resampling model. The effect of resampling the kinetic parameters from their posterior distribution is to add uncertainty to the inferred
 550 thermal history models. This is perhaps not as significant as we might think given the pairs (D and E_a) of kinetic parameters are very strongly correlated (Mialhe and Khoury, 1988). Then, resampling tends to produce combinations of the kinetic parameters that provide similar predictions for a given thermal history. However, the differences are still enough to add dispersion to the distribution of the thermal histories and increase uncertainty. The degree of additional uncertainty will depend on how the data are
 555 distributed in age and accumulated radiation damage.

6. Implications

Our method to propagate ZRDAAM uncertainty highlights how variable the age-[eU] relationship
 560 might be for a given thermal history. In particular, our results suggest that the uncertainty of ZRDAAM-based thermal history inversions may be significantly underestimated. This ~~may have~~ major implications for our ability to differentiate between subtle differences in temperature at specific times. In turn, it may be challenging to resolve cooling histories sufficiently to attribute the Great Unconformity to Cryogenic Glaciations ~~or geodynamic processes using thermochronology alone~~.

565 The potential to underestimate age uncertainty for thermal modelling has been discussed by McDannell et al. (2022) and to some extent this can be accounted for in the inverse modelling software QTQt (Gallagher, 2012). For example, if two dates have the same [eU] but their measured uncertainties do not overlap, QTQt can sample additional uncertainty for the ~~measurement age to allow~~ for this excess dispersion. However, if the two ages do not have the same [eU] concentration, the situation is more difficult. ~~There are two options~~. Either additional uncertainty can be assigned to the measurements by resampling a scaling factor (> 1) that multiplies the input errors, ~~but this~~ tends to allow the predicted age-[eU] relationship to pass through the observed data+resampled uncertainty. ~~Alternatively~~, the thermal history can be adjusted to change the predicted age-[eU] relationship to try and ensure that the predictions fit the data, at least to within the error. The first option tends to produce simpler thermal histories than the second option, as the data fitting criterion is less strict. For example, McDannell et al.
 575 (2022)’s results for Pikes Peak highlight how models that ignore overdispersion appear to resolve a 700

Deleted: has

Deleted: (McDannell et al., 2022) or geodynamic processes related to the break-up of Gondwana (Flowers et al., 2020).

Deleted: measurements

Deleted: account

Deleted: . This

Deleted: Or, alternatively

585 Ma cooling signature, which is smoothed out when the overdispersion is effectively reduced by adding excess uncertainty on some of the data. Importantly, Flowers et al., (2022) argued that the data do not have the ability to measure the timing of cooling irrespective of how overdispersion is treated.

We have shown the continuous spread of ages as a function of [eU] as a probability heat map for a specific history (Figure 4). In reality, most thermochronometric studies typically analyse 5-30 crystals for each sample. We can illustrate the effect of model uncertainty by comparing two simulated datasets of 15 ages generated from the same thermal history. We produce these datasets by sampling our age probability distribution randomly (Figure 6). For this example, each crystal in the dataset might have a different combination of model parameters drawn from the parameters we have estimated. Although these two datasets display overall similarities, there are subtle differences between their age-[eU] relationship over specific [eU] values. To accurately capture the spread in age for a single radiation damage value, many more thermochronometric samples would need to be collected. To illustrate this point, we draw random samples from the probability distribution in figure 4, for [eU] values ranging from 1500 to 2000 ppm, and investigate the spread in age (Figure 6). The dispersion of the ages varies greatly with increasing sample size, converging to the predicted frequency distribution of figure 4. In our specific example, the distribution stabilises for sample sizes of 40 crystals. The need to accurately capture spread are especially important if ages need to be averaged within [eU] bins to find acceptable paths as the uncertainty for the mean age is determined by the standard deviation (Flowers et al., 2020; Peak et al., 2021; Thurston et al., 2022). Ault et al. (2018) showed that simple visual identification under the microscope of the degree of metamictization is useful for obtaining good [eU] coverage and this approach could be adopted to ensure that multiple ages for the same [eU] are measured to get an idea of the spread in age. Zonation of eU concentration can also complicate the interpretation of thermochronometric data (Anderson et al., 2017; Weisberg et al., 2018; Anderson et al., 2018), potentially because different zones can evolve to have completely different diffusion kinetics (Fox et al., 2014). One promising approach to estimate the amount of zonation and anomalous diffusion behaviour could be to screen crystals using ramped degassing experiments (RAMPED).

The large uncertainties on the parameters controlling helium diffusion in zircon and the potentially dramatic impact this has on temperature sensitivity highlights that this aspect is important to consider. Currently, it may not be practical to incorporate diffusion kinetic uncertainties in inverse models directly because this dramatically increases the volume of the parameter space that needs to be searched and could lead to significantly longer run times with current software. However, with the development

Deleted: can

Commented [CA1]: sentence could be clearer

Deleted: is

Deleted: would

Deleted: long

Deleted: .

of faster computers and parallelized inverse methods, this ~~will not be a major obstacle. More im-~~
~~portantly~~, to ensure that we sample crystals with the same [eU] values to resolve diffusion kinetic pa-
 625 rameters, many more ~~grain~~ ages per sample need to be analysed. For example, to accurately capture the
 spread in age for a relatively narrow [eU] range of 1500-2000 ppm, 40 crystals from this interval were
 required (Figure 6), ~~and therefore for a sample with a wide range of eU values many multiples of this~~
~~number would be required~~. A practical solution to avoid measuring so many crystals per sample and
 running millions of simulations in an inversion is to use forward modelling. To do this, a single thermal
 630 history that is close to what might be expected for a specific area given prior ~~geological~~ knowledge
 could be used to assess expected age spread. This expected age spread could then be added to the age
 uncertainties used for inverse modelling. This procedure could be iterated to produce realistic uncer-
 tainties.

635 ZRDAAM has been calibrated using a limited number of diffusion experiments. Additional work is re-
 quired to develop this dataset to capture diffusion kinetics at different radiation damage values. These
 experiments could also aim to replicate diffusion kinetics at previously measured radiation damage val-
 ues to quantify the degree of dispersion. In addition, natural laboratories could be utilised to resolve
 diffusion parameters: areas with known thermal histories can be exploited to predict ZHe ages by vary-
 640 ing diffusion parameters; complementary thermochronometers can be leveraged to find thermal histo-
 ries and diffusion parameters that match the observed data. Ultimately, by reducing the uncertainties in
 helium diffusion kinetics using the constraints from man-made and natural laboratories, the timings of
 cooling events in the past can be resolved with ~~improved~~ accuracy and precision.

645 **Code availability.** Codes required for this analysis can be requested from the authors.

Data availability. No new data were generated for this analysis.

Author contributions. MF designed the analysis, carried out the uncertainty modelling, ~~on the kinetic~~
~~parameters~~. AS, PV and AC contributed to the concepts discussed. KG carried out the QTQt modelling.
 All authors contributed to writing.

650 **Competing interests.** The authors declare no competing interests.

Acknowledgements

We thank ~~Willy~~ Guenther for sharing data used for the calibration and for comments on ~~our~~ manu-
 script. We thank Alexis Ault, Rebecca Flowers, Kip Hodges, Brenhin Keller and Kalin McDannell (in

Deleted: may

Deleted: possible. Crucially

Deleted:).

Deleted: more

Deleted: .

Deleted: W.

Deleted: an earlier version of the

alphabetical order) for providing details comments on our manuscript. This study was supported by NERC (NE/N015479/1).

665 References

Anderson, A. J., Hodges, K. V., and van Soest, M. C., 2017, Empirical constraints on the effects of radiation damage on helium diffusion in zircon: *Geochimica et Cosmochimica Acta*, v. 218, p. 308-322, <https://doi.org/10.1016/j.gca.2017.09.006>.

Anderson, A. J., Hodges, K. V., and van Soest, M. C., 2018, Comment on 'Distinguishing slow cooling versus multiphase cooling and heating in zircon and apatite (U-Th)/He datasets: The case of the McClure Mountain syenite standard' by Weisberg, Metcalf, and Flowers: *Chemical Geology*, v. 498, p. 150-152, <https://doi.org/10.1016/j.chemgeo.2018.07.006>.

Ault, A.K., Guenther, W.R., Moser, A.C., Miller, G.H. and Refsnider, K.A.: Zircon grain selection reveals (de) coupled metamictization, radiation damage, and He diffusivity. *Chemical Geology*, 490, 1-12, <https://doi.org/10.1016/j.chemgeo.2018.04.023>, 2018

Barr, D.R. and Slezak, N.L., 1972. A comparison of multivariate normal generators. *Communications of the Association of Computing Machinery*, 15(12), pp.1048-1049, <https://doi.org/10.1145/361598.361620>.

Cooperdock, E.H., Ketcham, R.A. and Stockli, D.F., 2019. Resolving the effects of 2-D versus 3-D grain measurements on apatite (U-Th)/He age data and reproducibility. *Geochronology*, 1(1), pp.17-41, <https://doi.org/10.5194/gchron-1-17-2019>.

Formatted: Font: Not Bold

Deleted: <https://doi.org/10.1016/j.chemgeo.2018.04.023>

690

DeLucia, M. S., Guenther, W. R., Marshak, S., Thomson, S. N., and Ault, A. K.: Thermochronology links denudation of the Great Unconformity surface to the supercontinent cycle and snowball Earth, *Geology*, 46, 167–170, <https://doi.org/10.1130/G39525.1>, 2018.

Dodson, M. H.: Closure temperature in cooling geochronological and petrological systems, *Contrib.*

695 *Mineral. Petrol.*, 40, 259–274, <https://doi.org/10.1007/BF00373790>, 1973.

~~Fechtig, H. and Kalbitzer, S., 1966, The diffusion of argon in potassium-bearing solids. In *Potassium argon dating* (pp. 68-107). Berlin, Heidelberg: Springer Berlin Heidelberg.~~
~~https://doi.org/10.1007/978-3-642-87895-4_4~~

700 Flowers, R. M., Macdonald, F. A., Siddoway, C. S., and Havranek, R.: Diachronous development of Great Unconformities before Neoproterozoic Snowball Earth, *Proc. Natl. Acad. Sci.*, 117, 10172–10180, <https://doi.org/10.1073/pnas.1913131117>, 2020.

Flowers, R.M., Ketcham, R.A., Macdonald, F.A., Siddoway, C.S. and Havranek, R.E., 2022. Existing thermochronologic data do not constrain Snowball glacial erosion below the Great Unconformities.

705 *Proceedings of the National Academy of Sciences*, 119(38), p.e2208451119, <https://doi.org/10.1073/pnas.2208451119>.

Fox, M., McKeon, R.E. and Shuster, D.L., 2014. Incorporating 3-D parent nuclide zonation for apatite $4\text{He}/3\text{He}$ thermochronometry: An example from the Appalachian Mountains. *Geochemistry, Geophysics, Geosystems*, 15(11), pp.4217-4229.

710 Fox, M. and Shuster, D. L.: The influence of burial heating on the (U–Th)/He system in apatite: Grand Canyon case study, *Earth Planet. Sci. Lett.*, 397, 174–183, <https://doi.org/10.1016/j.epsl.2014.04.041>, 2014.

Fox, M. and Carter, A., 2020. Heated topics in thermochronology and paths towards resolution. *Geosciences*, 10(9), p.375, <https://doi.org/10.3390/geosciences10090375>.

715 Fox, M., Dai, J., and Carter, A.: Badly Behaved Detrital (U–Th)/He Ages: Problems With He Diffusion Models or Geological Models?, *Geochem. Geophys. Geosystems*, 2018GC008102, <https://doi.org/10.1029/2018GC008102>, 2019.

Deleted: .:

Deleted: Diffusion

Deleted: Argon

Deleted: -Bearing Solids, in: Potassium Argon Dating,

Formatted: Font colour: Auto

Formatted: Font colour: Auto

Formatted: Font colour: Auto

Formatted: Font colour: Auto

Formatted: Font: Italic, Font colour: Auto

Deleted: , Berlin, Heidelberg, 68–107,

Formatted: Font colour: Auto

Formatted: Font: 12 pt, Font colour: Auto

Deleted: , 1966.

Formatted: Font colour: Auto

Formatted: Font: 12 pt, Font colour: Auto

Formatted: Normal

- Gallagher, K.: Transdimensional inverse thermal history modeling for quantitative thermochronology, 725 J. Geophys. Res. Solid Earth, 117, n/a-n/a, <https://doi.org/10.1029/2011JB008825>, 2012.
- Gautheron, C., Djimbi, D.M., Roques, J., Balout, H., Ketcham, R.A., Simoni, E., Pik, R., Seydoux-Guillaume, A.M. and Tassan-Got, L.: A multi-method, multi-scale theoretical study of He and Ne diffusion in zircon. *Geochimica et Cosmochimica Acta*, 268, pp.348-367. <https://doi.org/10.1016/j.gca.2019.10.007>, 2020
- 730 Gordon Gastil, R., DeLisle, M., and Morgan, J.: Some Effects of Progressive Metamorphism on Zircons, *Geol. Soc. Am. Bull.*, 78, 879, [https://doi.org/10.1130/0016-7606\(1967\)78\[879:SEOPMO\]2.0.CO;2](https://doi.org/10.1130/0016-7606(1967)78[879:SEOPMO]2.0.CO;2), 1967.
- Guenther, W. R., Reiners, P. W., Ketcham, R. A., Nasdala, L., and Giester, G.: Helium diffusion in natural zircon: Radiation damage, anisotropy, and the interpretation of zircon (U-Th)/He thermochronology, *Am. J. Sci.*, 313, 145–198, <https://doi.org/10.2475/03.2013.01>, 2013. 735
- Keller, C. B., Husson, J. M., Mitchell, R. N., Bottke, W. F., Gernon, T. M., Boehnke, P., Bell, E. A., Swanson-Hysell, N. L., and Peters, S. E.: Neoproterozoic glacial origin of the Great Unconformity, *Proc. Natl. Acad. Sci.*, 116, 1136–1145, <https://doi.org/10.1073/pnas.1804350116>, 2019.
- Ketcham, R.A., Guenther, W.R. and Reiners, P.W.: Geometric analysis of radiation damage connectivity in zircon, and its implications for helium diffusion. *American Mineralogist*, 98(2-3), pp.350-360. <https://doi.org/10.2138/am.2013.4249>, 2013 740
- McDannell, K. T., Keller, C. B., Guenther, W. R., Zeitler, P. K., and Shuster, D. L.: Thermochronologic constraints on the origin of the Great Unconformity, *Proc. Natl. Acad. Sci.*, 119, e2118682119, <https://doi.org/10.1073/pnas.2118682119>, 2022.
- 745 McDannell, K.T., Zeitler, P.K., Janes, D.G., Idleman, B.D. and Fayon, A.K., 2018. Screening apatites for (U-Th)/He thermochronometry via continuous ramped heating: He age components and implications for age dispersion. *Geochimica et Cosmochimica Acta*, 223, pp.90-106, <https://doi.org/10.1016/j.gca.2017.11.031>
- 750 Mialhe, P., Charles, J.P. and Khoury, A., 1988. The thermodynamic compensation law. *Journal of Physics D: Applied Physics*, 21(3), p.383. <https://doi.org/10.1088/0022-3727/21/3/001>

Miltich, L., “Low temperature cooling history of Archean Gneisses and Paleo- proterozoic Granites of southwestern Minnesota,” Undergraduate thesis, Carleton College, Northfield, MN (2005).

755 Peak, B. A., Flowers, R. M., Macdonald, F. A., and Cottle, J. M.: Zircon (U-Th)/He thermochronology reveals pre-Great Unconformity paleotopography in the Grand Canyon region, USA, *Geology*, 49, 1462–1466, <https://doi.org/10.1130/G49116.1>, 2021.

Rahn, M. K., Brandon, M. T., Batt, G. E., and Garver, J. I.: A zero-damage model for fission-track annealing in zircon, *Am. Mineral.*, 89, 473–484, <https://doi.org/10.2138/am-2004-0401>, 2004.

760 Reiners, P. W.: Zircon (U-Th)/He Thermochronometry, *Rev. Mineral. Geochem.*, 58, 151–179, <https://doi.org/10.2138/rmg.2005.58.6>, 2005.

Tagami, T.: Zircon Fission-Track Thermochronology and Applications to Fault Studies, *Rev. Mineral. Geochem.*, 58, 95–122, <https://doi.org/10.2138/rmg.2005.58.4>, 2005.

Thurston, O. G., Guenther, W. R., Karlstrom, K. E., Ricketts, J. W., Heizler, M. T., and Timmons, J.
765 M.: Zircon (U-Th)/He thermochronology of Grand Canyon resolves 1250 Ma unroofing at the Great Unconformity and <20 Ma canyon carving, *Geology*, 50, 222–226, <https://doi.org/10.1130/G48699.1>, 2022.

Tripathy-Lang, A., Fox, M., and Shuster, D. L.: Zircon 4He/3He thermochronometry, *Geochim. Cosmochim. Acta*, 166, 1–14, <https://doi.org/10.1016/j.gca.2015.05.027>, 2015.

770 Weisberg, W. R., Metcalf, J. R., and Flowers, R. M., 2018, Distinguishing slow cooling versus multi-phase cooling and heating in zircon and apatite (U-Th)/He datasets: the case of the McClure Mountain syenite standard: *Chemical Geology*, v. 485, p. 90-99, <https://doi.org/10.1016/j.chemgeo.2018.03.038>.

Whipp, D.M., Kellett, D.A., Coutand, I. and Ketcham, R.A.: Modeling competing effects of cooling
775 rate, grain size, and radiation damage in low-temperature thermochronometers. *Geochronology*, 4(1), 143–152. <https://doi.org/10.5194/gchron-4-143-2022>, 2022.

Yamada, R., Tagami, T., Nishimura, S., and Ito, H.: Annealing kinetics of fission tracks in zircon: an experimental study, *Chem. Geol.*, 122, 249–258, [https://doi.org/10.1016/0009-2541\(95\)00006-8](https://doi.org/10.1016/0009-2541(95)00006-8), 1995.

Yamada, R., Murakami, M., and Tagami, T.: Statistical modelling of annealing kinetics of fission tracks in zircon; Reassessment of laboratory experiments, *Chem. Geol.*, 236, 75–91, <https://doi.org/10.1016/j.chemgeo.2006.09.002>, 2007.

Zeitler, P. K., Herczeg, A. L., McDougall, I., and Honda, M.: U-Th-He dating of apatite: A potential thermochronometer, *Geochim. Cosmochim. Acta*, 51, 2865–2868, [https://doi.org/10.1016/0016-7037\(87\)90164-5](https://doi.org/10.1016/0016-7037(87)90164-5), 1987.

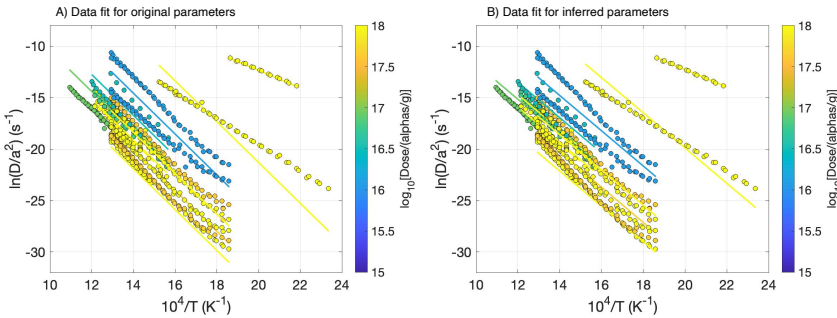


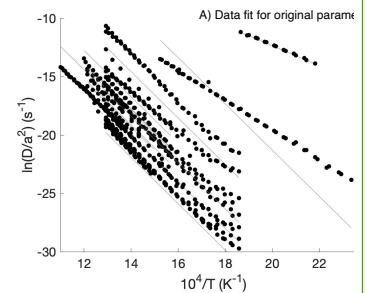
Figure 1. Key parameters controlling how radiation damage controls diffusivity have been inferred from fitting Arrhenius relationships from step-degassing experiments. A) The fit to the step-degassing experiments for the model parameters inferred by Guenther et al., (2013). B) The fit to the step-degassing experiments using model parameters extracted from the data using a Markov Chain Monte Carlo analysis. The model parameters are the minimum misfit model parameters and represent a single realization of the parameter set we infer. The point of this figure is not to show that one model predicts the diffusion data better than the other. It is simply to show the fits to the data and the data that are used to predict the model parameters, and to highlight that both models do not predict all the data.

Formatted: Normal

Formatted: Font: Cambria Math, 12 pt, Font colour: Auto

Formatted: Font colour: Auto

Formatted: Don't hyphenate



Deleted:

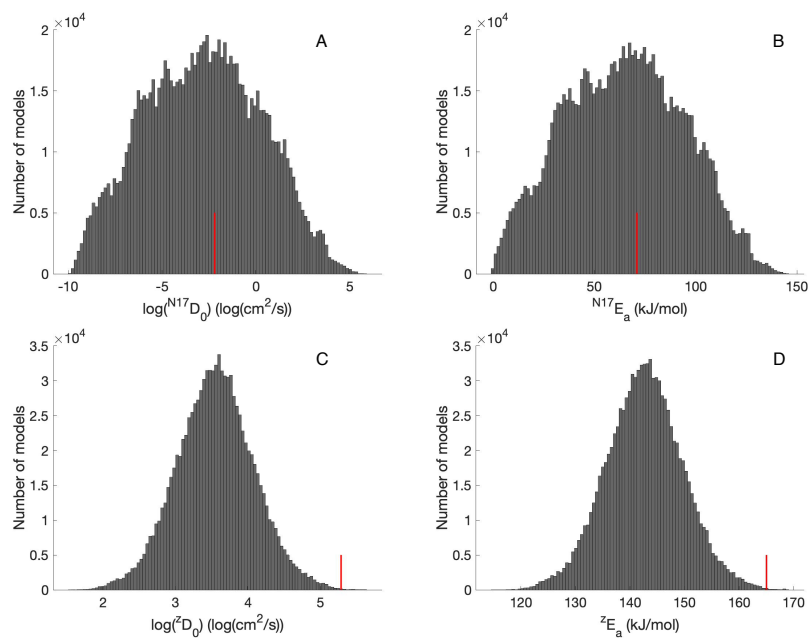


Figure 2. The sampled marginal posterior distributions for the four diffusion parameters representing the two hypothetical crystals. A) and B) are the frequency factor ($^{N17}D_0$) and activation energy ($^{N17}E_a$), respectively, for a damaged, amorphous crystal. C) and D) are the frequency factor ($^E D_0$) and activation energy ($^E E_a$), respectively, for an undamaged crystal. The red lines show the values of these parameters used by

810 Guenther et al., (2013).

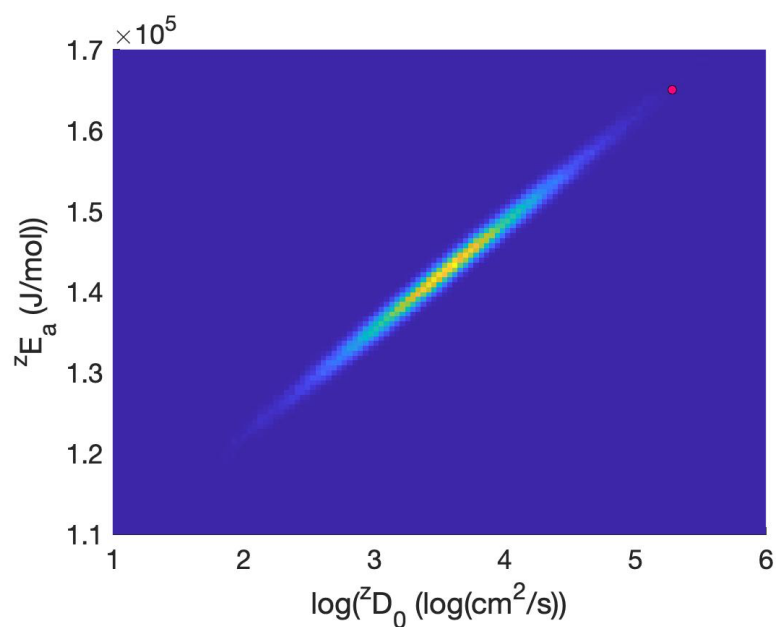


Figure 3. Inferred model parameters from diffusion data are strongly correlated. Our approach to infer the diffusion kinetics of the hypothetical crystals using the radiation damage model maintains this correlation. This is clearly illustrated with the diffusion parameters for the undamaged crystal. The pink spot shows the diffusion kinetics inferred by Guenther et al., (2013) and while it is reasonably far from the center of our inferred distribution, it still falls on the clear correlation trend defined by the sampling. The colours are proportional to posterior probability with oranges/yellows reflecting highest probabilities.

Deleted: shows that

Deleted: but

Deleted: we define.

Deleted: probability

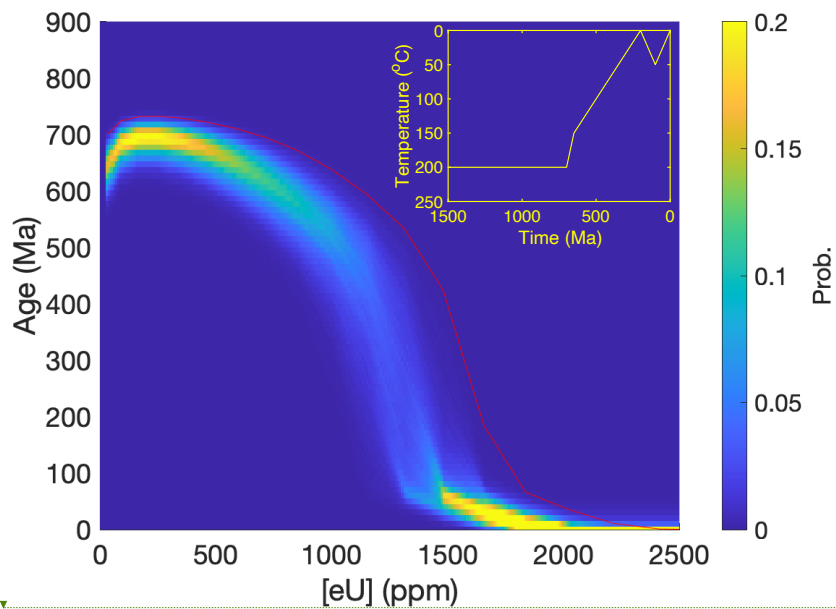
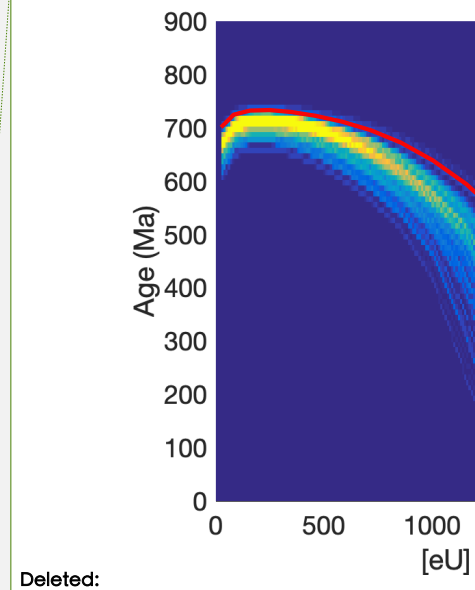
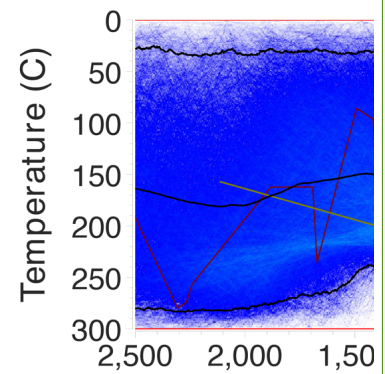


Figure 4. Propagating uncertainties in the radiation damage model produces a wide range of ZHe ages for a specific amount of damage. The red line shows the predicted age-[eU] relationship using the canonical values of the radiation damage and annealing model. The continuous thermal history used to produce the result is shown in the inset.

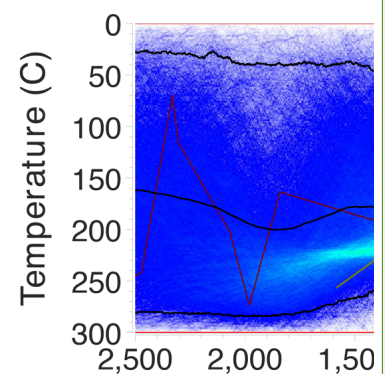


Deleted:

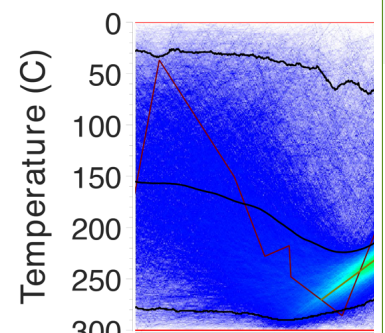
A) ZRDAAM Model Parameters



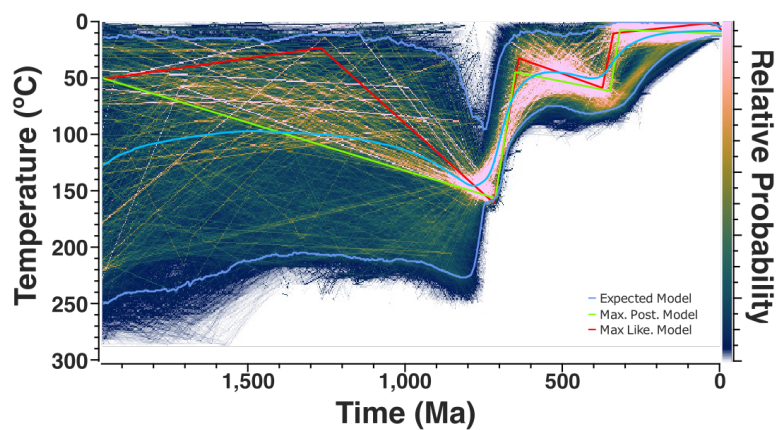
B) High amorphous frequency



C) Low amorphous frequency



A) Standard diffusion parameters



B) Accounting for uncertainty in diffusion parameters

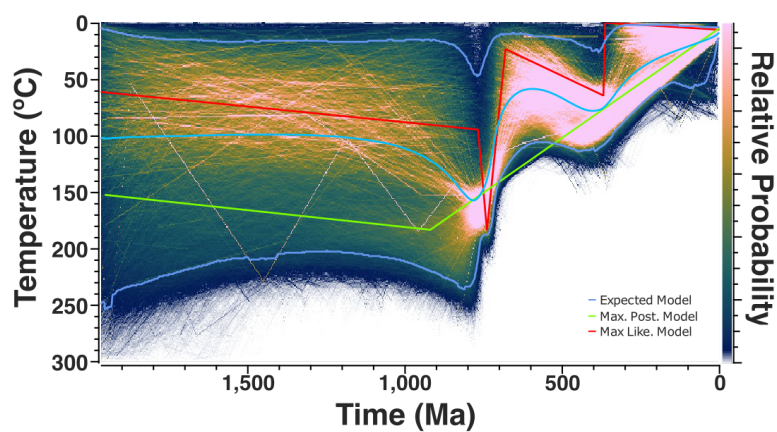


Figure 5. Thermal histories inferred using QTQt with (A) the Guenther diffusion parameters and (B) resampling of the 4 diffusion kinetics parameters. The resampling leads to a wider distribution of possible thermal histories, wider credible intervals, and also identifies 2 different forms, or modes in the posterior distribution (as seen by the Max Posterior and Max. Likelihood models). Note in this case, the Expected, or average, model falls between the two modes and so tends not to make good predictions relative to the other two models.

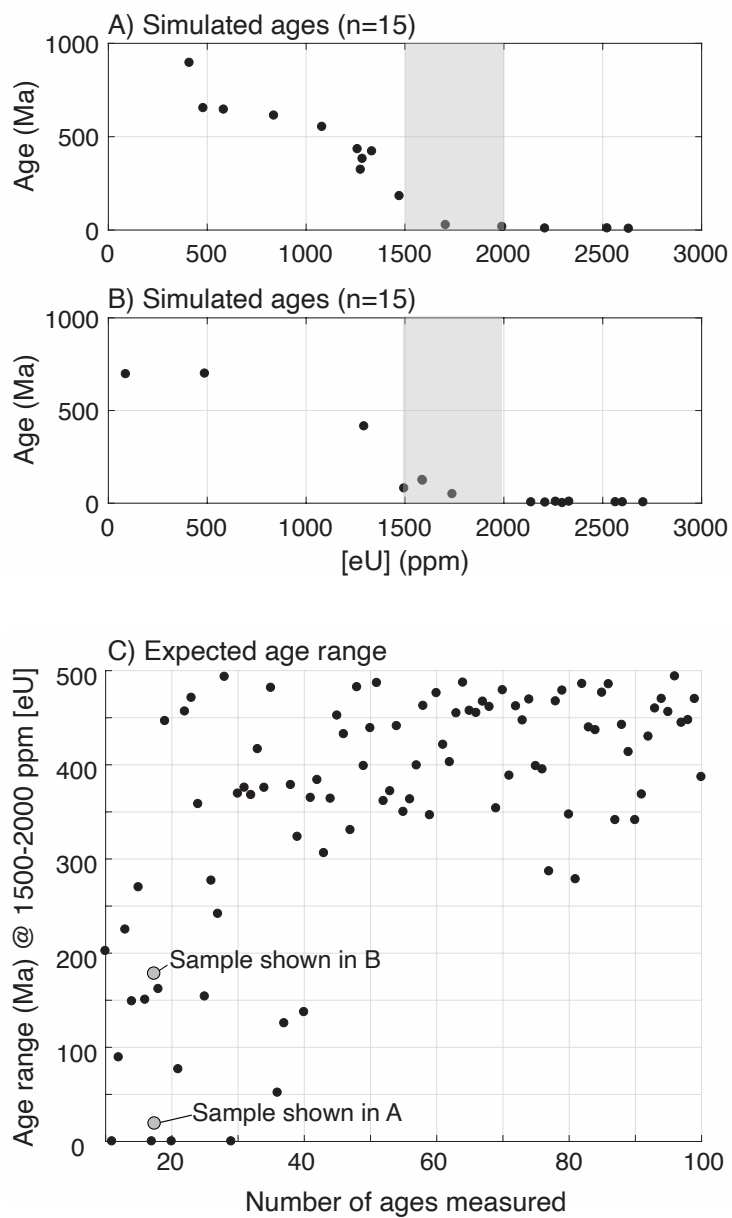


Figure 6. Model realizations and expected ranges of ages. A & B) random samples are drawn from the probability distribution in Figure 4 to highlight the sorts of datasets that are expected given the typical number of ages measured on a single sample. The simulated ages are different between the two realizations of a typical dataset. C) Many ages need to be sampled in order to accurately capture the spread in ages over a

875 specific [eU] bin. It is challenging to measure the spread because there is no easy way to estimate the amount
of radiation damage a crystal has accumulated until after the helium concentration has been measured.

880

885

890

Page 1: [1] Deleted	changes	14/01/2026 20:02:00
---------------------	---------	---------------------

▼

Page 1: [2] Deleted	changes	14/01/2026 20:02:00
---------------------	---------	---------------------

▼

Page 3: [3] Deleted	changes	14/01/2026 20:02:00
---------------------	---------	---------------------

▼

Page 3: [4] Deleted	changes	14/01/2026 20:02:00
---------------------	---------	---------------------

▼

Page 3: [5] Deleted	changes	14/01/2026 20:02:00
---------------------	---------	---------------------

▼

Page 3: [6] Deleted	changes	14/01/2026 20:02:00
---------------------	---------	---------------------

▼

Page 10: [7] Deleted	changes	14/01/2026 20:02:00
----------------------	---------	---------------------

▼

Page 21: [8] Deleted	changes	14/01/2026 20:02:00
----------------------	---------	---------------------

▼